

Local-scale snow accumulation variability on the Greenland ice sheet from ground-penetrating radar (GPR)

John Maurer
Konrad Steffen

Cooperative Institute for Research in Environmental Sciences (CIRES)
University of Colorado at Boulder

john.maurer@colorado.edu

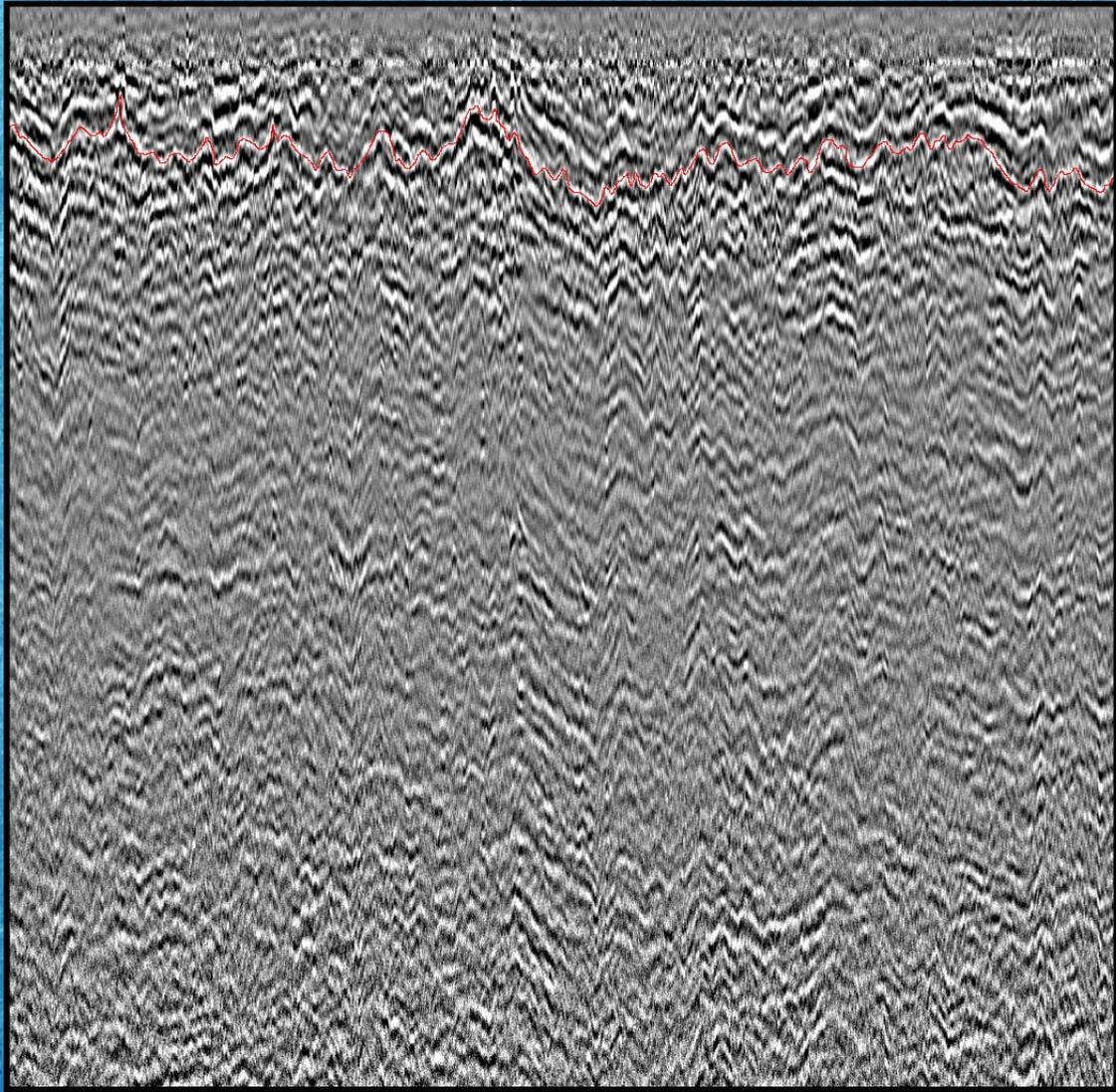
http://cires.colorado.edu/~maurerj/gpr/gpr_cryosphere.html



- **256** snow pits and ice cores
- collected 1913-1999
- 158 pre-1981, 98 post-1981
- **17** coastal met stations

Bales et al., *JGR*, 2001



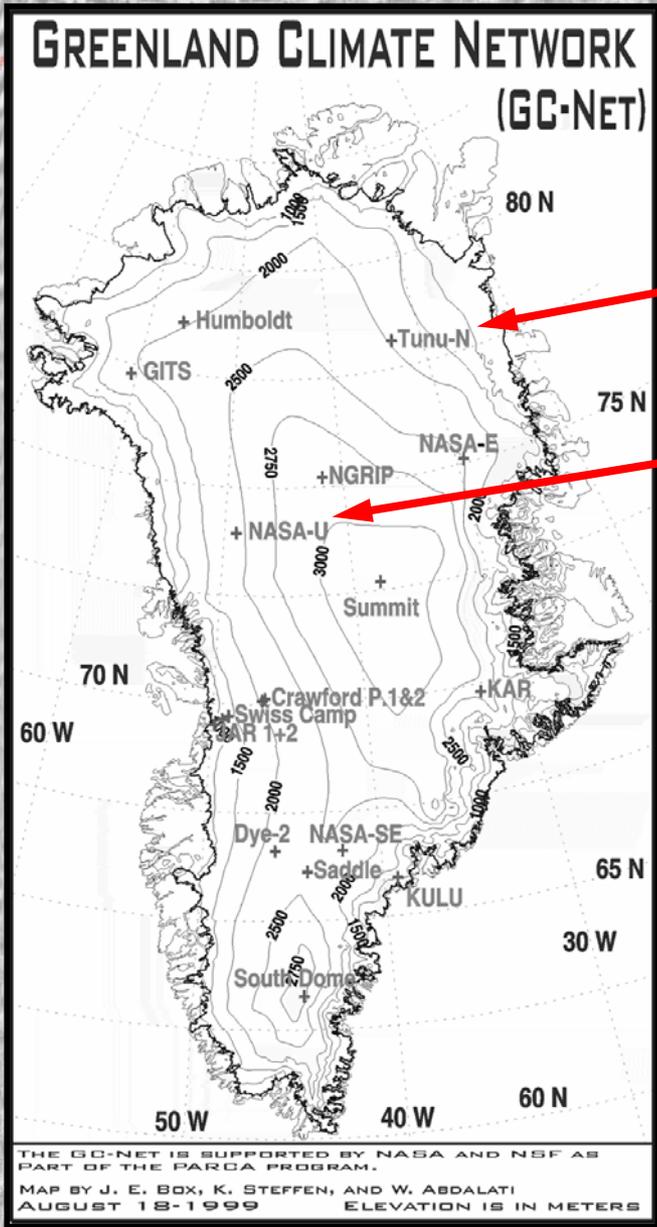


5 m

100 m



Data: 2003

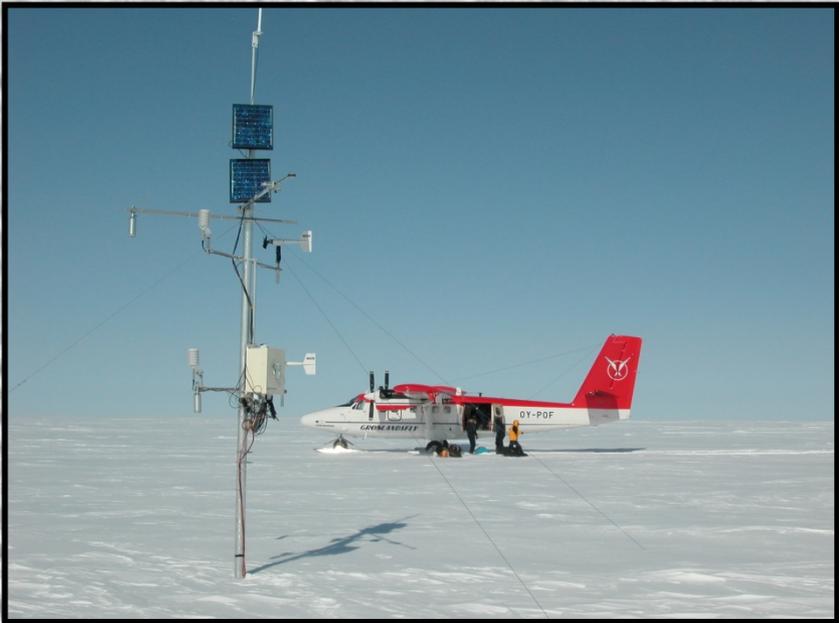


1. Tunu-N

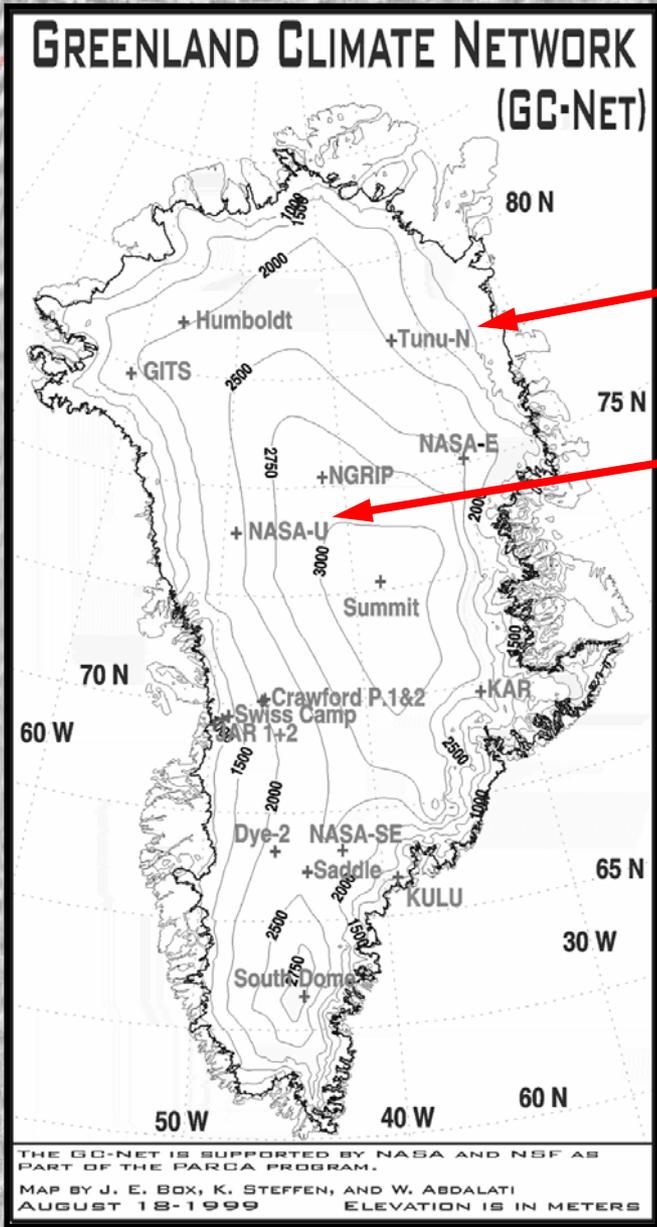
100-150 mm-yr⁻¹ SWE, ~40 cm snow depth

2. NASA-U

200-300 mm-yr⁻¹ SWE, ~75 cm snow depth



Data: 2003

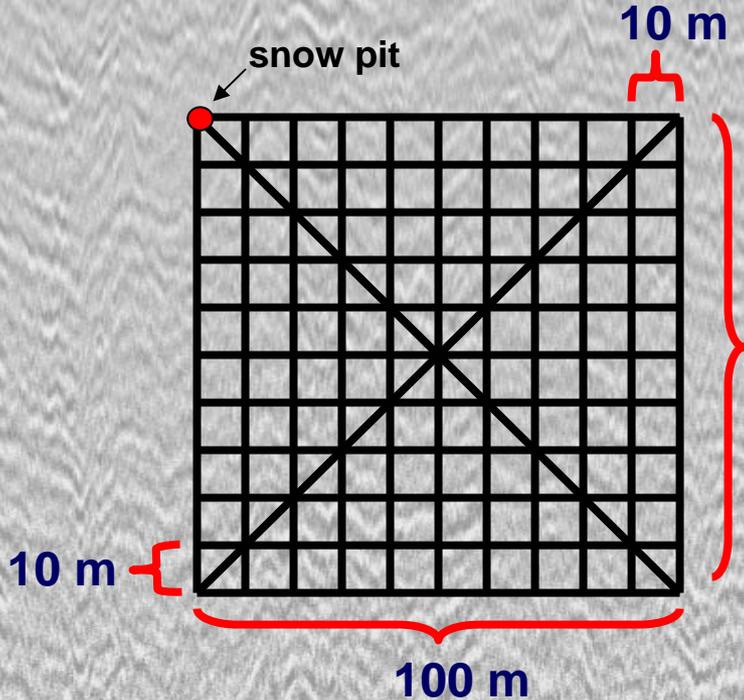


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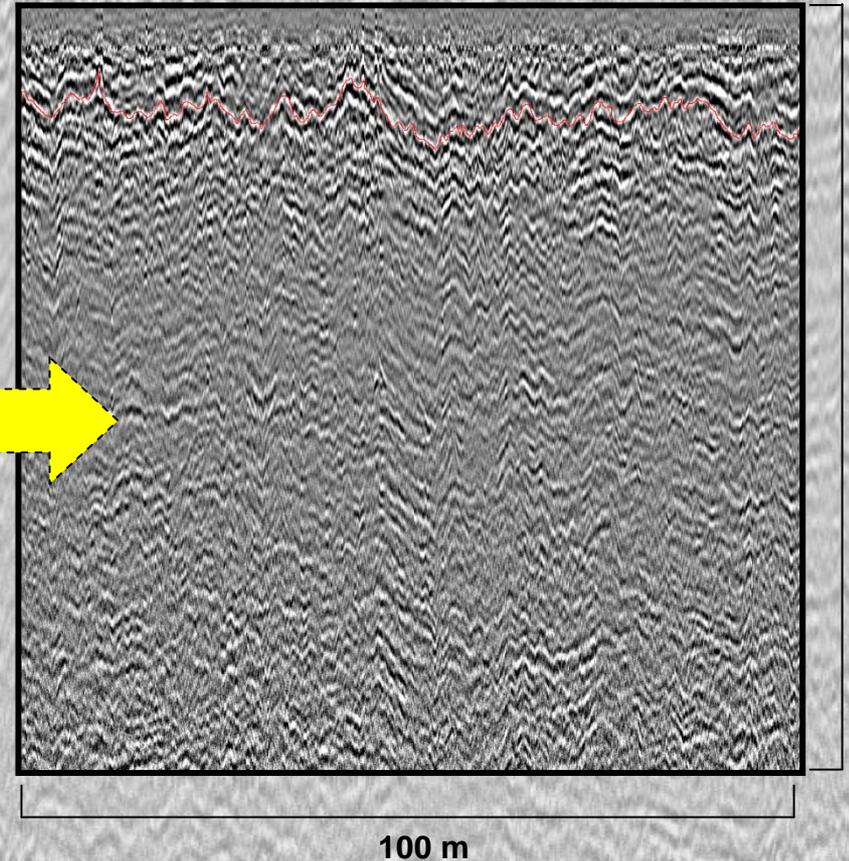
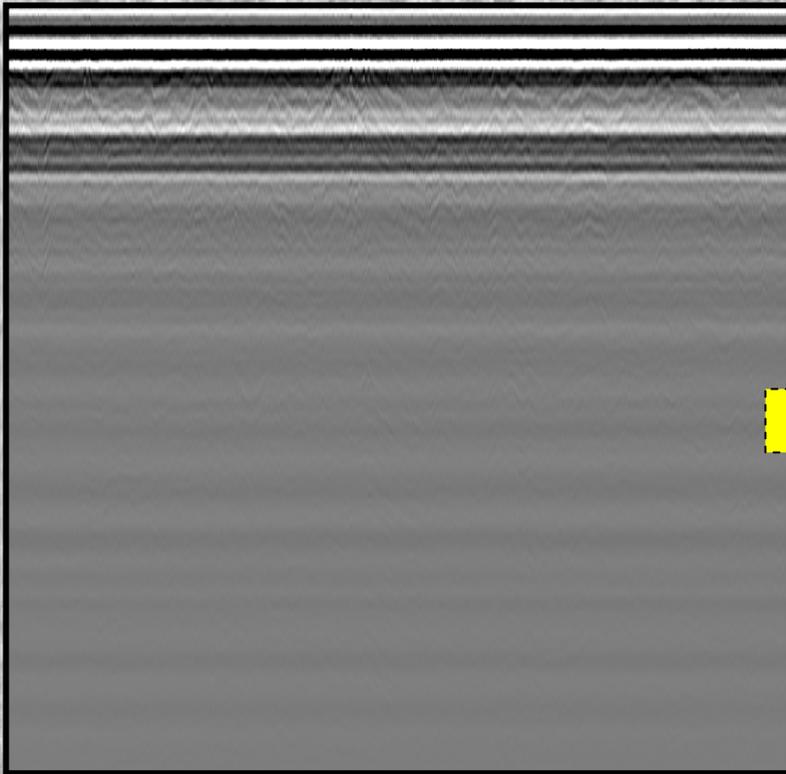
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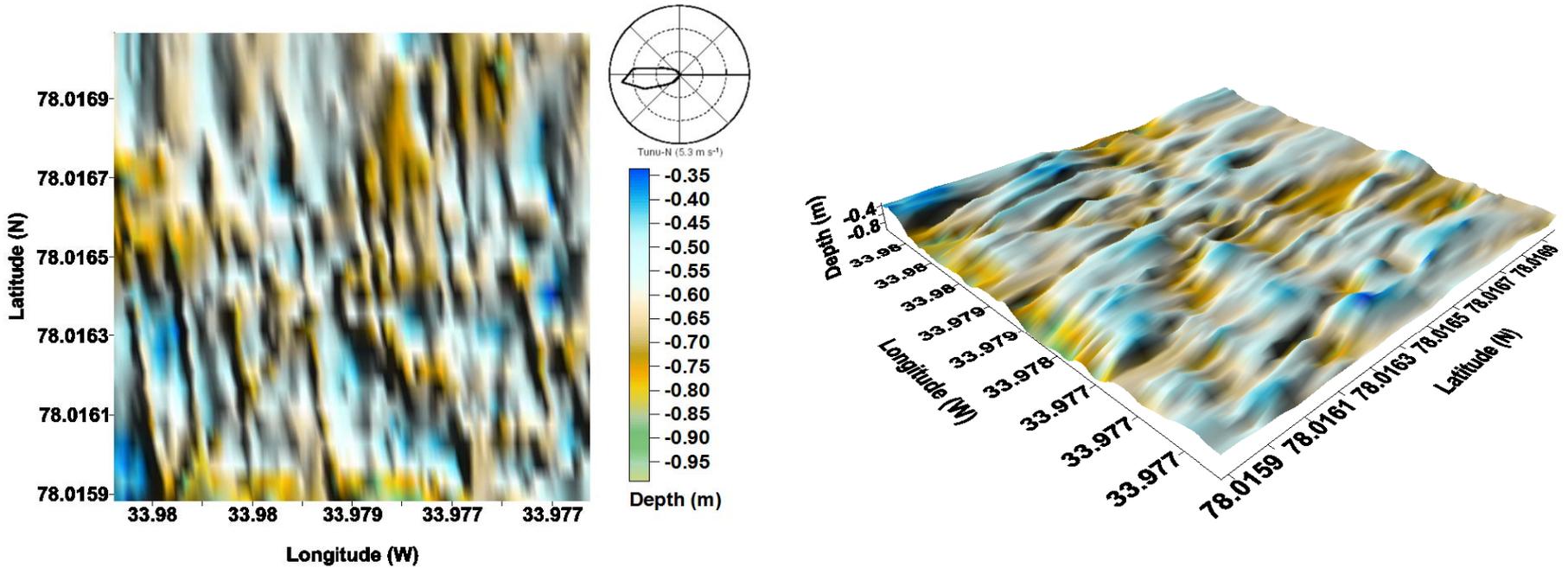
software developed for data processing



Issues: (1) antenna ringing, (2) spreading loss, (3) layer tracing, (4) depth calculation, (5) 3-D surface rendering



Tunu-N



Mean depth: 66 cm
Standard deviation: 12.2 cm
Minimum depth: 34 cm
Maximum depth: 105 cm



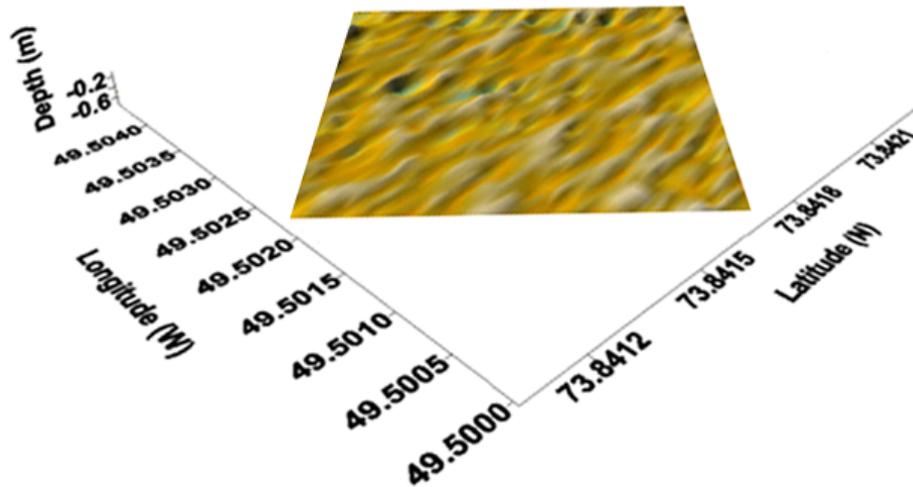
TABLE 1. Depth comparison of stratigraphic layers identified within the GPR data compared to their corresponding depths identified in a snow pit and by a nearby AWS sonic surface-height sensor, where possible.

<i>Site</i>	<i>Layer</i>	<i>Date</i>	<i>N</i>	<i>Mean Depth</i> (<i>cm ±1 S.D.</i>)	<i>Depth Range</i> (<i>cm</i>)	<i>Snow Pit</i> (<i>cm</i>)	<i>AWS</i> (<i>cm</i>)
Tunu-N	1	summer 2001	42554	65.6 (± 12.2)	33.8 – 105.5	50	64



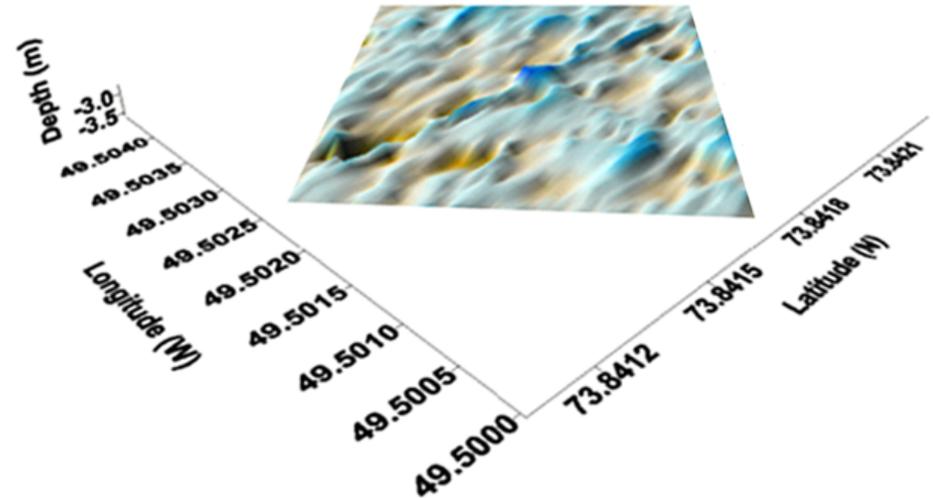
-2%

NASA-U layer #1



Mean depth: **50 cm**
Standard deviation: **7.4 cm**
Minimum depth: **35 cm**
Maximum depth: **89 cm**

NASA-U layer #2



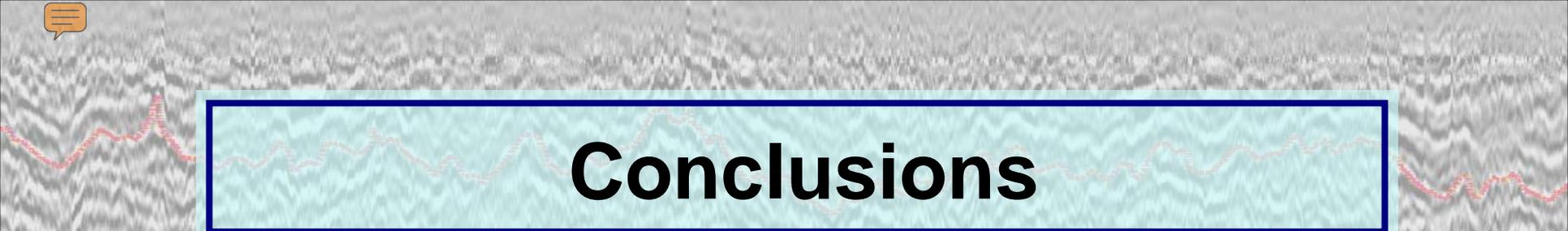
Mean depth: **302 cm**
Standard deviation: **11.3 cm**
Minimum depth: **270 cm**
Maximum depth: **356 cm**



TABLE 1. Depth comparison of stratigraphic layers identified within the GPR data compared to their corresponding depths identified in a snow pit and by a nearby AWS sonic surface-height sensor, where possible.

<i>Site</i>	<i>Layer</i>	<i>Date</i>	<i>N</i>	<i>Mean Depth</i> <i>(cm ±1 S.D.)</i>	<i>Depth Range</i> <i>(cm)</i>	<i>Snow Pit</i> <i>(cm)</i>	<i>AWS</i> <i>(cm)</i>
Tunu-N	1	summer 2001	42554	65.6 (± 12.2)	33.8 – 105.5	50	64
NASA-U	1	wind crust 2002	29291	49.9 (± 7.4)	35.2 – 88.5	54	n/a
NASA-U	2	ice layer 2000	31010	302.2 (± 11.3)	270.0 – 356.1	n/a	n/a

Average standard deviation: ±10 cm
Average range: ±35 cm



Conclusions

- **Spatial variability is important, even at local-scale (100 m x 100 m)**
- **GPR is an improvement over point measurements**
- **Aid in interpretation of ice cores & remote sensing**
- **Shallow penetration depth: cannot assess temporal variability**
- **High resolution: greatest strength & greatest weakness**
- **Complex stratigraphy can make interpretation difficult; need other *in situ* observations for validation**
- **Need more studies**



Questions?, Comments?



Thank you!

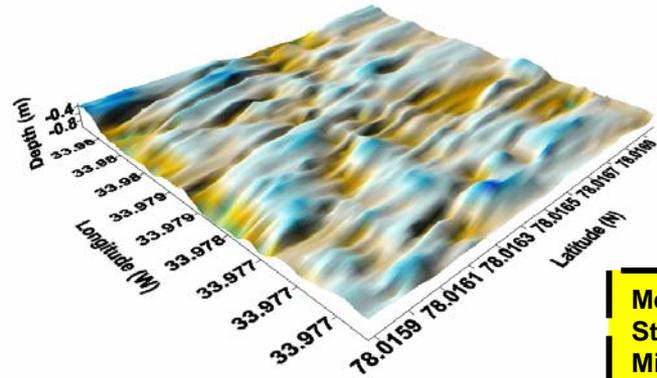
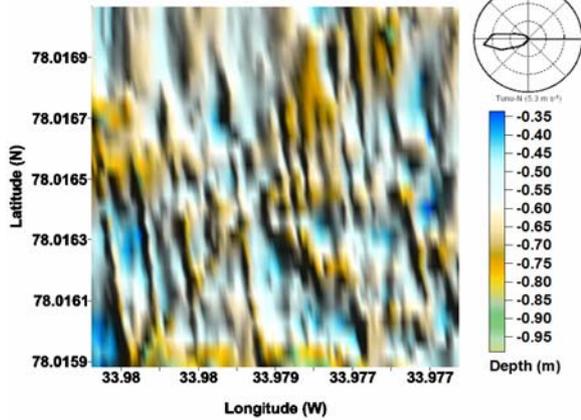
For further information:

http://cires.colorado.edu/~maurerj/gpr/gpr_cryosphere.html

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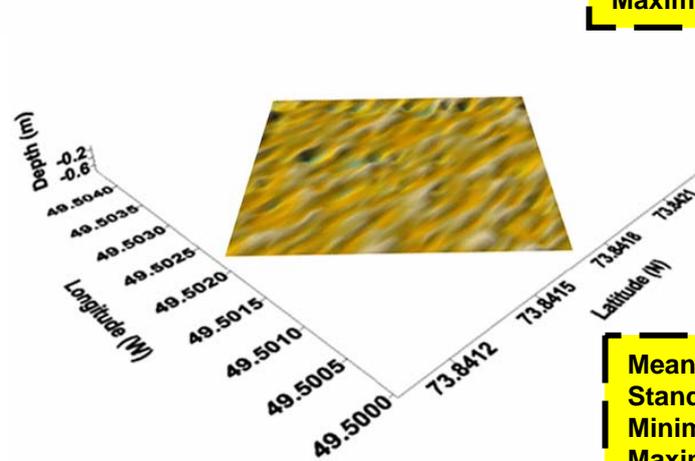
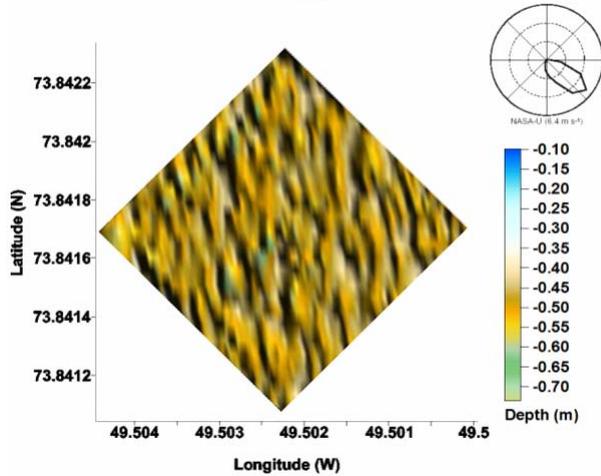


Tunu-N



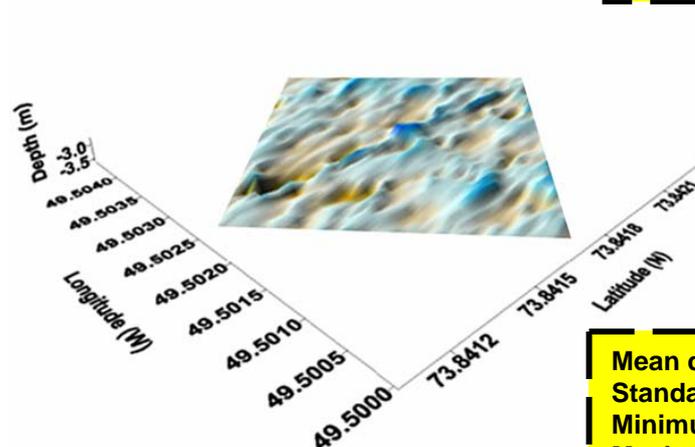
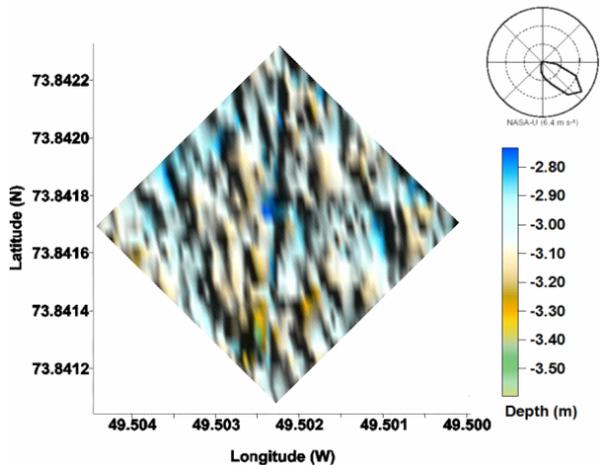
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Standard deviation: 11.3 cm
Minimum depth: 270 cm
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A Wideband Radar for High-Resolution Mapping of Near-Surface Internal Layers in Glacial Ice

Pannirselvam Kanagaratnam, Member, IEEE, Siva Prasad Gogineni, Fellow, IEEE, Vijay Ramasami, Student Member, IEEE, and David Braaten

Abstract—Snow accumulation rate is an important parameter in determining the mass balance of polar ice sheets. Accumulation rate is currently determined by analyzing ice cores and snow pits. Inadequate sampling of the spatial variations in the ice sheet accumulation has resulted in accumulation rate uncertainties as large as 24%. We designed an airborne radar system for high-resolution mapping of near-surface internal layers for estimating ice sheet mass balance. Our radar system can provide coverage by mapping a 300-km traverse over the ice sheet. During the flight, we successfully mapped the near-surface internal layers in the dry-snow zone, 120 m in the melt zone. We determined the accumulation rate at the NASA-U_1 site to be 34.6 cm/year for the year 1992. This is in close agreement with the accumulation rate of 34.6 cm/year for the year 1992.

Accumulation rate is a key variable in interpreting surface elevation measurements and in determining the mass input [5]–[7]. It is currently determined from ice cores and pits [8], and there are large errors in the estimated accumulation rate for

Index Terms—Airborne radar, internal layers, snow.

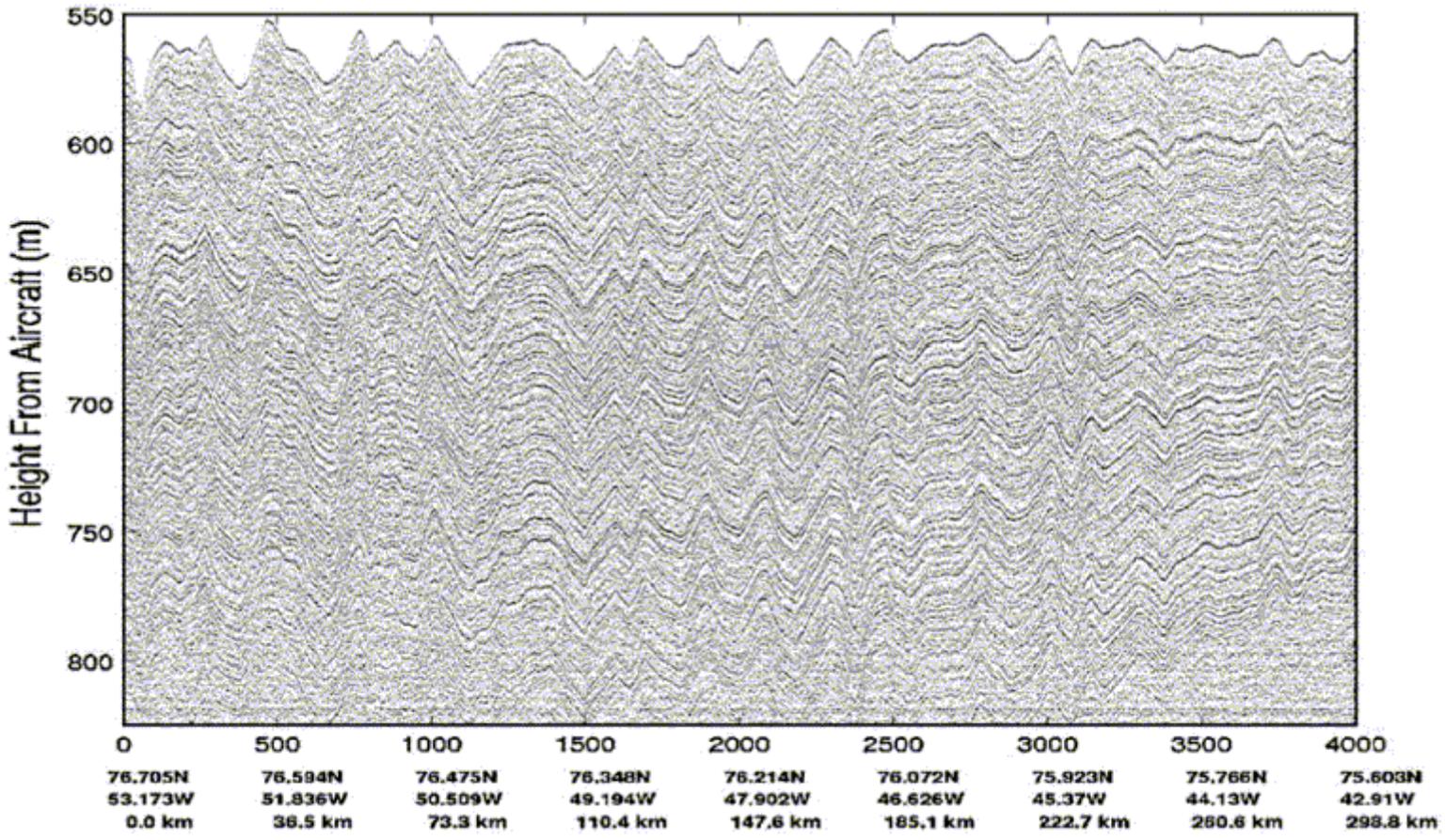


Fig. 4. (a) Internal layers observed over a 300-km traverse over the Greenland ice sheet.



sastrugi



Form parallel to predominant wind direction.

Antarctica. Source: Wikipedia



megadunes



Form perpendicular to predominant wind direction (like sand dunes).

Antarctica. Source: Ted Scambos, NSIDC