

Sand Waves Generation: A Numerical Investigation of the Infiernillo Channel in the Gulf of California

Yovani Montaña Ley¹, Noel Carbajal²

¹Instituto de Ciencias del Mar y Limnología, Universidad Nacional Autónoma de México, Unidad Académica Mazatlán, Mazatlán, México

²Instituto Potosino de Investigación Científica y Tecnológica, San Luis Potosí, México
Email: yovan@ola.icmyl.unam.mx, noelc@ipicyt.edu.mx

Received 21 April 2016; accepted 12 July 2016; published 15 July 2016

Copyright © 2016 by authors and Scientific Research Publishing Inc.
This work is licensed under the Creative Commons Attribution International License (CC BY).
<http://creativecommons.org/licenses/by/4.0/>



Open Access

Abstract

The effect of the coastal geometry on sand bed forms generation has been investigated for a tidal dominated area. Different hypothetical geometries of coastal channels with flat bottoms and unlimited sediment availability were exposed to strong oscillatory tidal currents to simulate the interaction of hydrodynamics and the bedload sediment transport. The hypothetical geometries stand for the idealization of the principal geographic features of the Infiernillo Channel, a coastal area of the Gulf of California where sandbanks and sand waves have been observed. A depth integrated hydrodynamic-numerical model and a parameterized formula to estimate the bedload sediment transport were applied coupled with a sediment conservation equation to determine the sea bottom morphodynamics. Model predictions in the Infiernillo Channel were compared to available satellite imagery. This investigation demonstrates that a vertical integrated numerical model is able to reproduce the development of incipient sand waves that exist in the Infiernillo Channel. Incipient sandbanks and shoals were also simulated. Sand waves with wavelengths of about 200 m were calculated on the same locations where sand waves actually exist. A crucial finding of this research was to show that the geometry of a shallow water basin and the presence of tidal velocity gradients associated with abrupt changes in the coastline alignment were critical in determining the sand-bed pattern generation. We demonstrate that a vertical variation of tidal currents is not necessary to generate sand waves.

Keywords

Sand Waves, Tidal Hydrodynamics, Bed Load, Sediment Transport

1. Introduction

Sand waves are rhythmic seabed features of great practical relevance. There are many economic activities interested in the presence and/or evolution of sand waves. Indeed, their migration can represent a serious hazard to pipelines that are laid in regions of intense sea-bottom activity. Dredging may be required because sand waves migrate into or along shipping channels and harbors, hence reducing the local water depth and, consequently, affecting the navigability. It has been recognized that near-shore, migrating sand waves might even affect shoreline evolution and coastal defense works [1]. The presence of sand waves in some particular areas is strongly correlated to the amplitude of the spring tidal current [2]. One of the most striking characteristics of sand waves is that they are not static bed forms. They migrate with their crests almost orthogonal to the direction of the velocity oscillations induced by the tide, at a rate that strongly depends on the tidal characteristics that can reach some tens of meters per year [3]. Their shapes are often asymmetric but the steepness of the sides does not lead to flow separation [4]. Sand waves have spatial scales smaller than other bed-forms like sandbanks [5]. The wavelengths of sand waves are of some hundreds of meters, with amplitudes of several meters, and crests perpendicular to the principal tidal current axis [3] [6]. Tidal sand banks and sand wave fields can sometimes overlap, even with heights of the same order [7].

As pointed out by Hulscher [4], Gerkema [8] and other authors, the process that gives rise to the formation of sand waves is similar to originating dunes in rivers [9] or to that causing the appearance of sea ripples under sea waves [10]. The interaction of the oscillatory tidal current with a bottom perturbation gives rise to a steady streaming in the form of re-circulating cells. When the net displacement of the sediment dragged by this steady streaming steers toward the crests of the bottom waviness, the amplitude of the perturbation grows and bed-forms are generated. Hulscher [11] noted that the model of Huthnance [12] only predicts the formation of sand ridges, but no sand waves. She attributed it to the lack of a vertical structure of tidal currents. Hulscher [4] suggests that the vertical effect produced by a three dimensional model has to be included in order to generate sand waves. Contrary to the previous suggestion, our hypothesis is that sand waves may be generated and evolve even in the absence of vertical velocity gradients. Hence, the primary goal of this work was to test this hypothesis reproducing numerically sand waves by using a depth-integrated model coupled with a parameterized bedload sediment transport and a sediment conservation equation. In this investigation, we were able to reproduce successfully sand waves. Furthermore, we demonstrate that the presence of strong current horizontal velocity gradients associated with abrupt changes in the coastline alignment is critical in determining the sand-waves generation.

The Infiernillo Channel is a complex coastal wetland that includes reef, mangrove vegetation and coastal lagoons. Due to its high productivity, endemic characteristics, and fisheries, it has been also considered as a priority zone of conservation by the OEMGC (Marine Ecological System of the Gulf of California), as well as appropriate site to develop ecotourism activities. It is also a reserved fisheries area for the indigenous people by the government of Mexico [13]. There are some efforts to study the sand waves in this area: Merrifield *et al.* [14] carried out field surveys and observations of sand waves. Marzolf and Merrifield [15] made observations of tidal currents in the Infiernillo channel. In the present research work was carried out a numerical study on sand waves formation in the Infiernillo channel.

2. Study Area

The Infiernillo Channel is a narrow, shallow and long channel, averaging 7.6 km width that separates Tiburon Island from the Sonoran mainland (Figure 1). The averaged water depth of the channel is 5.5 m and its length is about 30 km. The adjacent areas to the channel are Agua Dulce, Sargento, Kunkaak and Kino Bays as well as the Tiburon Island. There are not continuous river inflows into the channel and only few alluvial flows into the sea. Surface water temperature fluctuates between 15°C in winter and 30°C in summer. The average salinity for Infiernillo Channel is 35‰, however salinities up to 40‰ have been recorded within the Channel in winter. Benthonic organisms cover large areas of the Infiernillo channel. These biological species are of ecological importance and of great economic value for the people who live around the area. The study of the sea bottom morphodynamics where these resources find their habitats and the relevance of the issue of coastal and biotic resources conservation is evident, hence the interest of this study.

Mixed tides characterize the study area. Most of Tiburón Island presents a rocky shoreline, however el Infiernillo Channel shows a sandy bottom and shoreline [16]. Merrifield *et al.* [14] describe sand waves and tidal

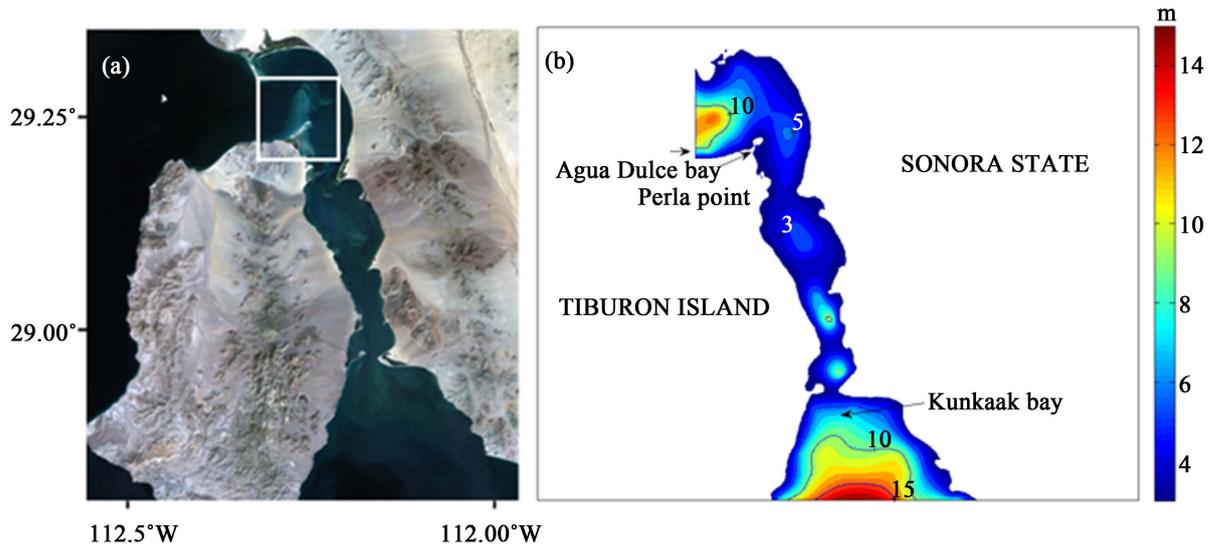


Figure 1. Geography (a) and topography (b) of the Infiernillo Channel, located in the Gulf of California. Agua Dulce and Kunkaak bays and Perla point are indicated. Satellite image obtained from U.S. *Geological Survey*. Landsat 5 TM, acquisition date 2011/10/12, color composite (3,2,1).

currents in the Infiernillo channel. Ebb currents flowing south from the channel form sand waves and tidal current ridges in Kunkaak Bay in water depths of less than 3 m. Surface ebb currents up to 1.7 knots, producing Froude Numbers of about 0.27 have been measured. The sand waves have wavelengths ranging from 12 m to 75 m and amplitudes from 0.30 m to 0.75 m. The sediment consists of well-sorted, coarse-grained sand; size-frequency distributions of the sediment samples are positively skewed.

In Agua Dulce Bay at the north end of the channel, sand waves developed in fine-grained sand. Size-frequency distributions of the sediment samples are negatively skewed. The sand on the crests is coarser in grain size and less negatively skewed than the sand in the troughs. The sand waves have wavelengths of 180 m to 250 m and amplitudes from 1.1 m to 1.5 m [14]. The sea-bottom of the Infiernillo channel consists of non-cohesive sediments, mainly sand forming shoals in the northern part of the channel as well as in the southern part. The shoals are one meter below the water surface and some of them emerge through the surface during the ebb at spring tides.

3. Methods

A two dimensional hydrodynamic-numerical model has been applied to simulate the depth average tidal currents in the Infiernillo channel, located in the archipelago region of the Gulf of California. Hence, the bedload sediment transport and the bottom morphodynamics were simulated for different geometrical configurations. Carbajal and Backhaus [17], Montañó *et al.* [18] and Montañó-Ley *et al.* [19] [20] applied the model for different sea regions. It has been successfully applied to calculate the propagation of tidal waves in regions like the North Sea, the Gulf of California and the Topolobampo coastal lagoon, Mexico, among other sites. The model is based on the vertically integrated shallow water equations:

$$\frac{\partial U}{\partial t} + \frac{U}{H + \zeta} \frac{\partial U}{\partial x} + \frac{V}{H + \zeta} \frac{\partial U}{\partial y} - fV = -g(H + \zeta) \frac{\partial \zeta}{\partial x} + A_H \left(\frac{\partial^2 U}{\partial x^2} + \frac{\partial^2 U}{\partial y^2} \right) - r \frac{U \sqrt{U^2 + V^2}}{(H + \zeta)^2} \quad (1)$$

$$\frac{\partial V}{\partial t} + \frac{U}{H + \zeta} \frac{\partial V}{\partial x} + \frac{V}{H + \zeta} \frac{\partial V}{\partial y} + fU = -g(H + \zeta) \frac{\partial \zeta}{\partial y} + A_H \left(\frac{\partial^2 V}{\partial x^2} + \frac{\partial^2 V}{\partial y^2} \right) - r \frac{V \sqrt{U^2 + V^2}}{(H + \zeta)^2} \quad (2)$$

and the continuity equation

$$\frac{\partial \zeta}{\partial t} + \frac{\partial U}{\partial x} + \frac{\partial V}{\partial y} = 0 \quad (3)$$

x and y are the horizontal space variables, t is the time, U is the water transport ($\text{m}^2\cdot\text{s}^{-1}$) in the x direction, V is the water transport ($\text{m}^2\cdot\text{s}^{-1}$) in the y direction, ζ is the sea surface elevation (m); A_H is the horizontal turbulent exchange coefficient ($\text{m}^2\cdot\text{s}^{-1}$), $f = 2\Omega \sin \phi$ is the Coriolis parameter, $\Omega = 7.29 \times 10^{-5} \text{ s}^{-1}$ is the angular velocity of the earth, ϕ is the latitude ($\sim 29.25^\circ\text{N}$ in these calculations), H is the depth (m); g is the gravitational acceleration ($\text{m}\cdot\text{s}^{-2}$); and r is the friction coefficient. The model was forced at the open boundaries with a sea surface elevation oscillation of the form

$$\zeta = \zeta_0 \cos(\omega t - \varphi) \quad (4)$$

where ζ_0 is the amplitude, φ the phase and ω the frequency of the forcing wave, *i.e.* the frequency of the tidal constituent. To force the dynamics, tidal elevations and phases were prescribed at the two open boundaries, located in the northern and southern inlets, for the tidal harmonics K_2 , S_2 , M_2 , N_2 , K_1 , P_1 and O_1 . The tidal harmonics were obtained from tidal simulations in the whole Gulf of California [17] and by interpolation of harmonics from Carbajal [21]. At the open boundaries, the gradient of the velocity was equal to zero.

$$\frac{\partial V_n}{\partial n} = 0 \quad (5)$$

where V_n and n are the component of the velocity and space variable, respectively, normal to the open boundary. More details of the model can be seen in [21].

The applied bathymetry was obtained from Lancin [22]. The depth averaged tidal velocities predicted through the 2D hydrodynamic model were used to feed the sediment transport model. The parameterized bedload transport model applied in this investigation, has been given by Huthnance [12], Hulsher *et al.* [11], Schuttelaars and De Swart [23], Van Rijn [24], Montaña-Ley [25], Montaña and Carbajal [26] and Carbajal and Montaña [27] [28]. The volumetric sediment flux S_b in the active layer, given by Van Rijn [24] is:

$$S_b = s \frac{|\mathbf{u}|^b}{u_{cr}^b} \left(\frac{\mathbf{u}}{|\mathbf{u}|} + k_* H_x \right) \quad (6)$$

$$s = \sqrt{g^1 d_{50}^3} \quad (7)$$

$$g^1 = (\rho_s - \rho) g / \rho \quad (8)$$

for some $b > 1$ and $k_* > 0$. The term measured by the coefficient k_* is a bed slope correction term which models the preferred downhill transport of sediment. Typical values for b and k_* are $b \approx 3$ and $k_* \approx 2$. The parameter s is a function of the sediment properties, ρ_s is the sediment density (2650 kg/m^3) and d_{50} the mean sediment diameter. The value of s was calculated for $d_{50} = 200$ microns, according to the Equations (7) and (8) [24] [29] [30]. The water density is given by ρ and the critical velocity by u_{cr} , *i.e.* the velocity value ($\sim 0.3 \text{ m}\cdot\text{s}^{-1}$) to initiate the bed-load transport of sediment. The bed-load sediment transport components were obtained by using the velocity components and the bottom slope components in Equation (6).

The erosion-accretion patterns of sediments, *i.e.* the changes of the sea bottom, were determined by the sediment transport divergence:

$$\frac{\partial H}{\partial t} = \nabla \cdot \mathbf{S} \quad (9)$$

An important feature of this model is that the shape of the seabed is updated at every time interval and the current and sediment transport patterns are recomputed using the new bathymetry. A review of approaches to long term modeling of coastal morphology is given by De Vriend *et al.* [31].

4. Results

The interaction between the tidal hydrodynamics and the sea bottom was investigated for the better understanding of the morphodynamics of coastal areas like the Infiernillo channel. In this research work, the morphodynamics in the Infiernillo channel has been investigated through several numerical experiments for different geometrical configurations, *i.e.* to investigate the influence of the coastal geometry on the sea-bottom morphodynamics. A depth integrated finite difference hydrodynamic model has been applied. The numerical experiments

were performed in order to test the hypothesis that the geometry of a tidal basin has a profound effect on the sea-bottom morphodynamics, specifically on the sand wave generation.

This investigation consisted of three parts: firstly, different geometries with flat bottoms and unlimited sediment availability were exposed to oscillatory tidal currents in order to simulate the interaction of hydrodynamics and the bed-load sediment transport. It is important to mention that these geometries exemplify the dominant geographical features. Secondly a realistic coastal configuration of the Infiernillo Channel was used with the assumption of a flat sandy sea-bottom. In the third case, a real bathymetry and the actual configuration of the Infiernillo channel was applied in the numerical experiments.

The predicted depth averaged instantaneous tidal currents of the infiernillo channel has have been presented in **Figure 2**. Maxima predicted tidal current of about 0.7 m/s were predicted when the water flows into Kunkaak bay (t = 6 h), located in the southern part of the infiernillo channel. **Figure 2** shows the vector fields for four different stages of the tides. Some areas presented convergence of the flow while others show flow divergence. Another feature of the velocities fields was the presence of a cyclonic eddy located right to the north of Pearl point (see **Figure 1**).

Figure 3 includes the morphodynamic of the sea bottom predicted for different geometries. The experiments started with a flat sandy sea bottoms and after 700 M_2 tidal periods (approximately one year) the vertical changes were plotted. The assumption of flat bottom has been made in order to understand the effects of coastal geometry on the bottom morphodynamics. The maximum predicted vertical change was about 0.08 m. of accreted sediment while a minimum value of -0.08 m of eroded sediment was predicted for the elapsed time. It is worthwhile to note that bed-form patterns have been generated. The scales of these patterns corresponded quite well to the typical length scales of sand waves of 100 to 500 meters [4] (**Figure 4(b)**). The same order of

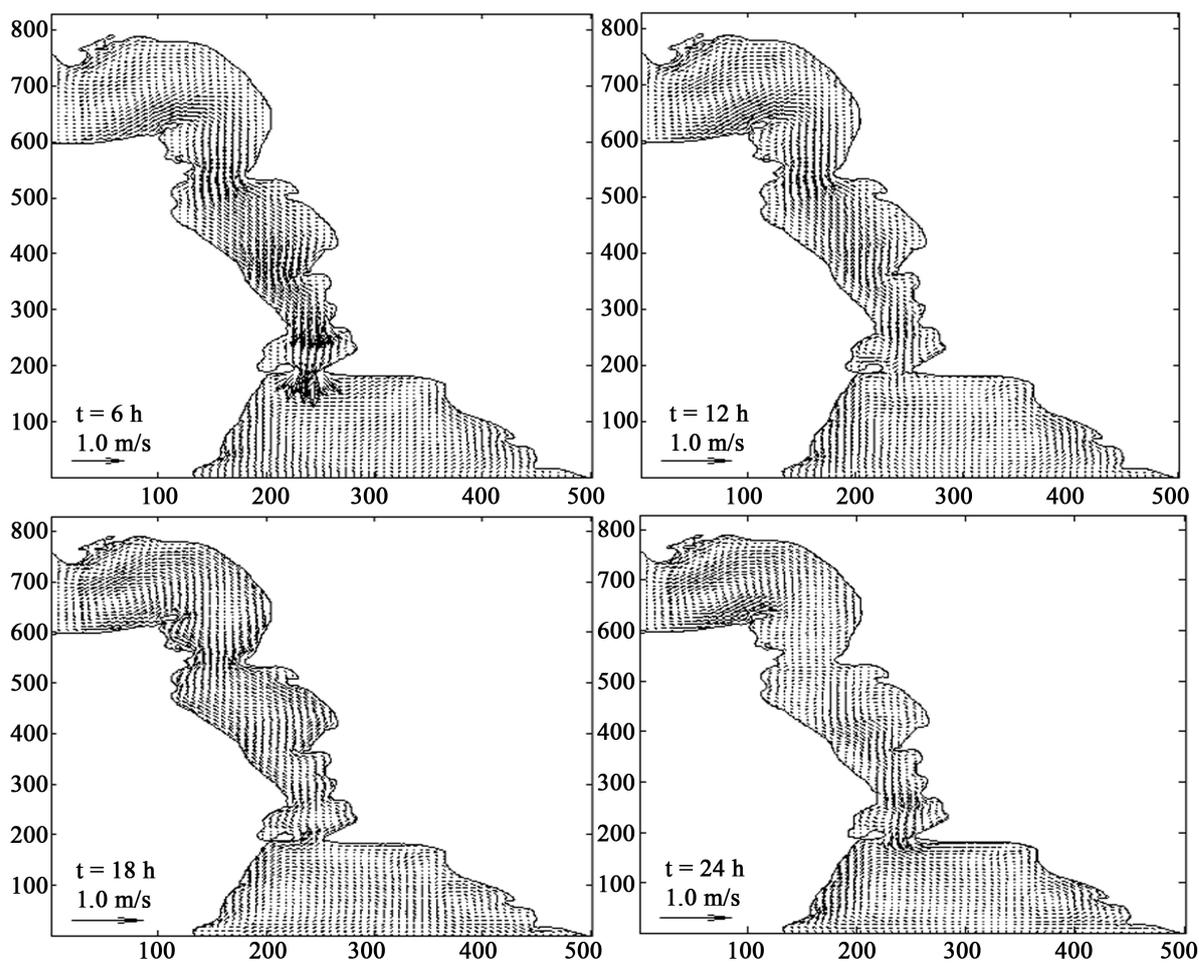


Figure 2. Tidal currents in the Infiernillo channel at different times.

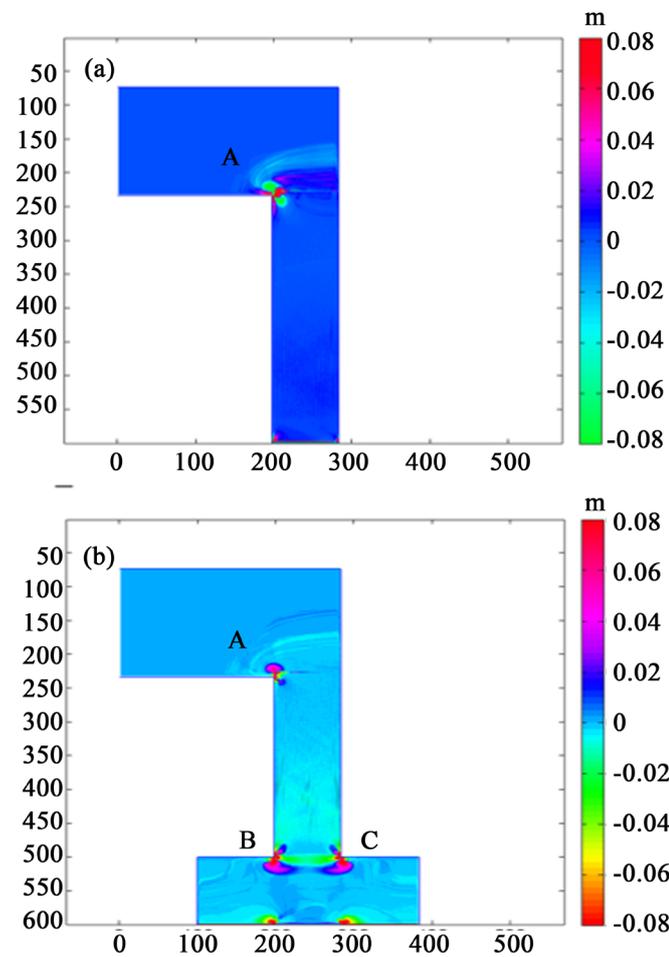


Figure 3. Bottom morphology after one year of simulation for two different geometrical configurations representing the most important geographical features of the Infiernillo channel.

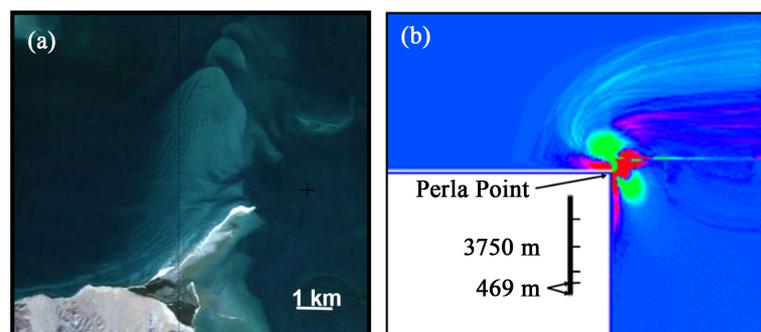


Figure 4. Comparison between observed (a) and calculated (b) sand waves in the Infiernillo channel. Satellite image obtained from U. S. *Geological Survey*. Landsat 5 TM, acquisition date 2011/10/12, color composite (3,2,1).

magnitude of the sand wavelengths were observed in satellite imagery of the Infiernillo channel (**Figure 4(a)**). It is also important to mention the presence of shoals located to the east of Perla point in **Figure 4(a)** and **Figure 4(b)**. It is also worthwhile to note that the predicted sand waves and shoals in the considered coastal geometries were probably influenced by high velocity gradients or abrupt changes (90°) in the coast alignment.

The second group of simulations includes a real coastal geometry of the Infiernillo channel but assuming a flat bottom (**Figure 5**). Under these conditions the vertical changes of the sea bottom were also computed after 700 M2 tidal periods. The results revealed the formation of tidal sand waves, shoals and probably sand banks. Once again, the sites where the sand waves were predicted coincide remarkably well with real sites where sand waves are observed in satellite imagery (see **Figure 4**). **Figure 5** shows in detail sandwaves generated near Pearl point at the northern end of the Infiernillo channel. The incipient sand waves reveal wavelengths of about 250 meters. Around Perla Point there is a tendency to the formation of shoals with separation between crests of about 1 km. The most outstanding aspect to be mentioned is that sand waves are generated in the same area in all calculations, *i.e.* the fundamental features of a sea region determines where sand waves are generated.

Similarly, **Figure 6** shows sandwaves evolving in the southern end of the infiernillo channel, *i.e.* in the Kun-kaak bay. Another interesting aspect is the incipient rhythmic bed forms with distances between crests varying

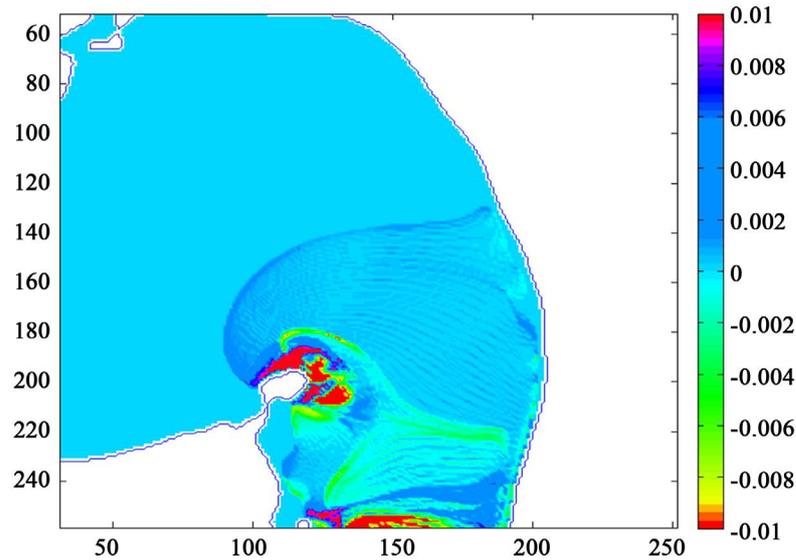


Figure 5. Incipient sand waves and shoals in the northern part of the Infiernillo channel.

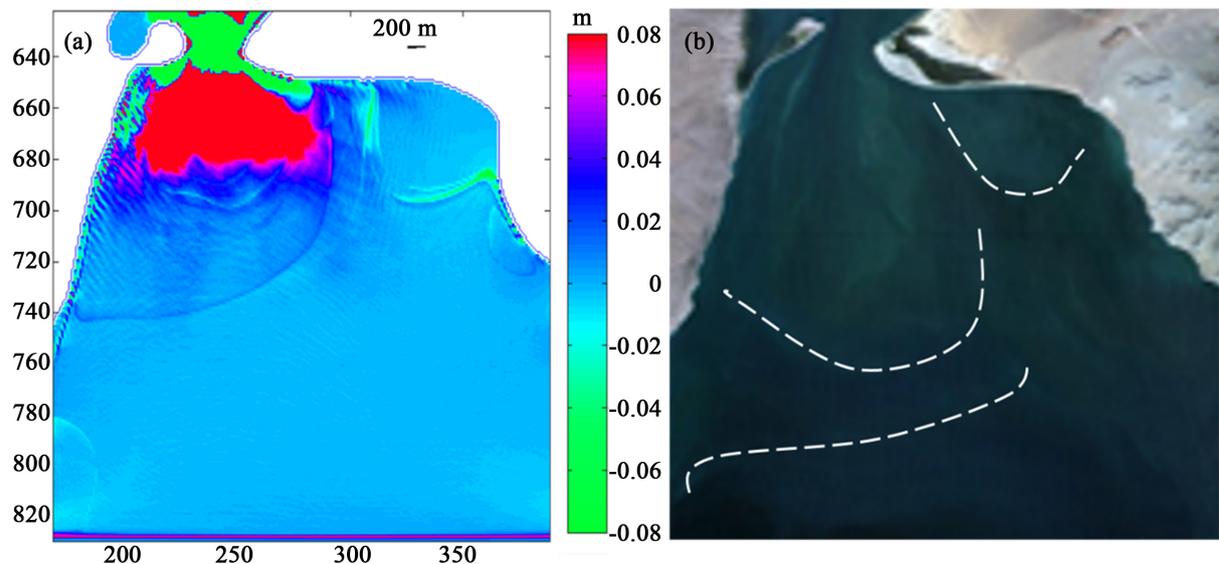


Figure 6. Calculated sand waves pattern in the southern part of the infiernillo channel (a). The konkaak bay area was taken from **Figure 1(a)**. Dashed lines indicate the position of shoals. Satellite image obtained from U.S. *Geological Survey*. Landsat 5 TM, acquisition date 2011/10/12, color composite (3,2,1).

between about 1400 and 1800 meters. Satellite imagery also shows intense morphodynamics in the same area. The shoals which appear in the calculated pattern, showed in **Figure 6(a)**, are also observed in the satellite image depicted in **Figure 6(b)** where the dashed lines indicate the position of the shoals. The comparison of calculated and observed shoals reveals a similar pattern.

The third group of numerical experiments were undertaken to simulate bedforms generation under realistic circumstances, *i.e.* a more accurate coastal geometry and bathymetry of the Infiernillo channel was applied in the calculations. Incipient wavy bed forms appear in different areas of the Infiernillo channel (**Figure 7**). Clearly, sand waves were generated near to Perla Point, in the central water body and in the southern part of the Infiernillo channel. These results agree with the previous calculations for more simple geometrical characteristics. The regions with intense morphodynamics are the same in all numerical simulations.

5. Discussion

A few decades ago, there was the perception that sand waves are generated by the vertical structure of tidal currents. Theoretical models, using vertically averaged velocities, could explain only the generation of sandbanks, *i.e.* bed forms with scales varying approximately between 2 and 14 km. The failure to explain the generation of sand waves was attributed to the lack of a vertical variation of tidal currents which is suppressed by the depth integration [11]. Later, it was demonstrated that with the consideration of a vertical variation of tidal currents, the theoretical model could predict the generation of bed forms with length scales similar to those of sand waves [4]. Several theoretical calculations of sand waves considered a vertical variation of the tidal flow [32]. Numerical models have also been able to simulate the sand wave generation [33]. Németh *et al.* [34] developed a two-dimensional numerical model describing vertical flow variations, morphological changes and the behavior of sand waves. They calculated a steady current situation induced by a wind stress applied at the sea surface. The bed forms found have wavelengths in the order of hundreds of meters when the resistance at the seabed is relatively large. The results also show that it is possible to model the initial evolution of sand waves with a numerical simulation model. They focus on sand waves in a unidirectional steady current with vertical variation.

The hypothesis that sand waves may be generated by depth averaged tidal currents and not necessarily by vertical velocity gradients has been tested. A crucial finding of this research work is then the fact that the geometry of a shallow water coastal areas and the presence of strong tidal currents (0.7 m/s) as well as abrupt changes in the coastline alignment were crucial in determining the generation of sand waves and other bed forms. It was possible to simulate the formation of sand waves applying a vertical integrated two-dimensional numerical

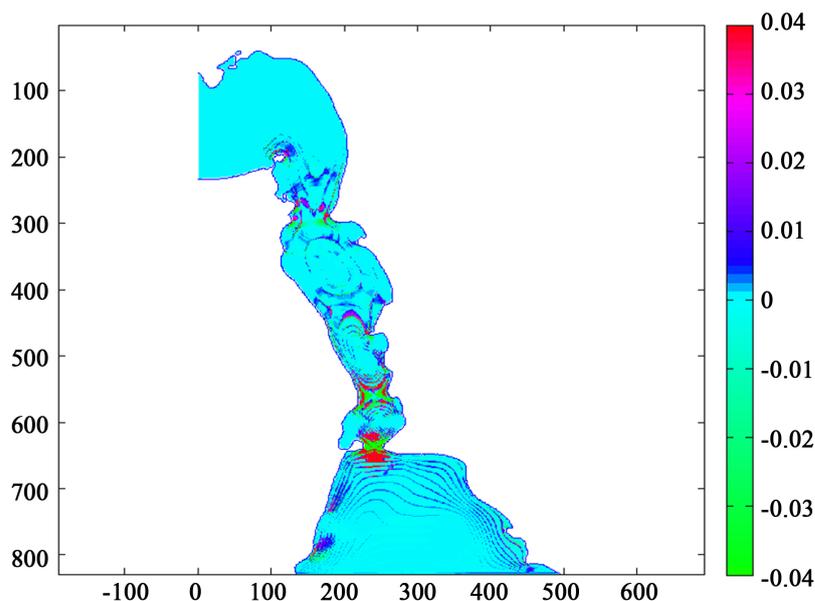


Figure 7. Calculated bed forms applying realistic geometrical conditions. accurate coastal and topographic information was used.

model. This model considers non-linear terms included in the equations of motion like the friction and the advection terms. These terms could represent the difference with respect to those results obtained in theoretical calculations. The reason for the generation of sand waves could be the non-linear interaction of Poincare waves produced by the interaction of tidal currents and the geometry of the Infiernillo channel [35]. It is important to mention that theoretical models, two or three-dimensional, consider an infinite ocean, *i.e.* no boundaries are contemplated. It represents a large difference with this 2D numerical model where a realistic interaction between tidal currents and a complex geometry is considered. Berthod and Pattiaratchi [36] studied the influence of the coastal geometry on rhythmic large-scale bed patterns. They applied a morphological model to an idealized headland situation. In the present investigation, sand waves and other bed forms were generated numerically at different locations in the main water bodies of the Infiernillo channel. The results of the numerical simulations agree with observations of sand waves made by Merrifield [14] near of Perlaspoint, in the northernmost part of the channel, and in the southern part where the Infiernillo channel widens into Kunkaak Bay. Other types of bedforms were also predicted in these numerical calculations. Shoals and sand banks were generated mainly in areas where the tidal flow interacts strongly with the geography (Figure 7). As far as we know, our investigation is the first work that simulates sand wave generation over an extensive sea bottom area applying complete vector fields velocities obtained from a 2D integrated model.

Particular features of bed forms in the Infiernillo channel have been observed. According to Merrifield [14], sand bars were detected near to Agua Dulce Bay. In the same zone, sand waves with wave lengths between 200 and 270 m and amplitudes between 1.0 and 1.5 m were also documented. These sand waves developed in very fine sand conditions. Theoretical estimates indicate that the crests of the sand waves are oriented about 10° anticlockwise from the principal tidal flow [11]. Comparison between the tidal flow showed in Figure 2 and the crest pattern depicted in Figure 7 indicates that this property is reproduced adequately since crests and principal tidal flow are almost in the same direction. In Kunkaak bay, sand waves perpendicular to tidal currents were superimposed on much larger sand ridges. The calculated sand waves had wavelengths of approximately 75 m and amplitudes of about 0.32 m. In the present investigation sand waves were numerically predicted in both locations near Agua Dulce bay and in Kunkaak bay.

Figures 5-7 reveal the connection between the presence of sand waves and shoals and geographical features which resemble headland and bays. It is worthwhile to notice that the predicted sand waves in all the coastal geometries were influenced by the presence of high velocity gradients associated with abrupt changes (90°) in the coastline alignment (Figure 3). These bed patterns will probably evolve to more defined sandbanks or sand waves. Sand waves with typical length scales of the order of hundreds of meters are also observed.

Following the discussion on the bedform generation there are more evidences of the crucial effect of the coastal geometry effects. Applying a two-dimensional morphodynamic model for a tidal channel, Schramkowski *et al.* [37] showed that estuarine bars with typical length scales on the order of the tidal excursion length can be generated as the result of a positive feedback between water motion, sediment transport and the sandy bottom in a coastal configuration. Convergences and divergences of the tidally averaged sediment fluxes result in the evolution of the bed. The bottom morphodynamics of the infiernillo channel investigated through a 2D model show good qualitative agreement with the findings of Schramkowski *et al.* [37]. In The infiernillo channel the effects were the development of sand waves and sand banks. In agreement with Schramkowski *et al.* [37], the bedform pattern is generated from flows that arise from tide and geometry interactions. Specifically, such flows always set up net sediment fluxes towards the formation of shoals and sand waves. In the present investigation we probe that coastal geometries particularly headlands and abrupt changes in the coastline orientation are associated to the generation of shoals and sand waves.

6. Conclusions

In the present investigation, sand wave and shoal patterns, resulting from the interaction of oscillatory tidal currents with the sea bottom have been predicted. The patterns were characterized by wavelengths comparable with those observed in the field or in satellite imagery.

We were able to reproduce sand waves. We also showed, through the application of a 2D depth integrated model, that sand waves are generated even in the absence of vertical velocity gradients. Sand waves formed near Perla point in the northern end of the Infiernillo Channel and in Kunkaakbay in the southern part while shoals and sand banks formed almost in the entire channel area.

The coastal geometry alignment turned out to play an important role in bedforms generation: Headlands and acute shifts in the shore alignment apparently trigger sand waves and sand bank formation. Convergence or divergence of tidal flow areas seems also to generate shoaling areas.

Another important result is that the calculated sand waves appear in the same area where they are observed in satellite imagery. The different numerical experiment yielded always the correct zone where sand waves exist. Since in these numerical experiments the main geographic properties were maintained, it is an indication that this aspect (where they occur) is determined by the coastal configuration.

Acknowledgements

Thanks to Victoria Montes-Montes for her assistance in this investigation and to the project: PAPIIT-UNAM IN-208673, for its support. We also thank to María de Jesús Guevara-Macías for her assistance in handling satellite imagery.

References

- [1] Besio G., Blondeaux, P., Brocchini, V. and Vittori, G. (2003) Migrating Sand waves. *Ocean Dynamics*, **53**, 232-238. <http://dx.doi.org/10.1007/s10236-003-0043-x>
- [2] Belderson, R.H., Johnson, M.A. and Kenyon, N.H. (1982) Offshore Tidal Sands. In: Stride, A.H., Ed., *Offshore Tidal Sands, Processes and Deposit*, Chapman Hall, London, 27-57.
- [3] Terwindt, J.H.J. (1971) Sand Waves in the Southern Bight of the Northern Sea. *Marine Geology*, **10**, 51-67. [http://dx.doi.org/10.1016/0025-3227\(71\)90076-4](http://dx.doi.org/10.1016/0025-3227(71)90076-4)
- [4] Hulscher, S.J.M.H. (1996) Tidal-Induced Large-Scale Regular Bed Form Patterns in a Three-Dimensional Shallow Water Model. *Journal of Geophysical Research: Oceans*, **101**, 20727-20744. <http://dx.doi.org/10.1029/96JC01662>
- [5] Off, T. (1963) Rythmic Linear Sand Bodies Caused by Tidal Currents. *AAPG Bulletin*, **47**, 324-341.
- [6] McCave, I.N. (1971) Sand Waves in the North Sea off the coast of Holland. *Marine Geology*, **10**, 199-225. [http://dx.doi.org/10.1016/0025-3227\(71\)90063-6](http://dx.doi.org/10.1016/0025-3227(71)90063-6)
- [7] Huntley, D.A., Huthnance, J.M., Collins, M.B., Liu, C.L., Nicholls, R.J. and Hewitson, C. (1993) Hydrodynamics and Sediment Dynamics of North Sea Sand Waves and Sand Banks. *Philosophical Transactions of the Royal Society London*, **343**, 461-474. <http://dx.doi.org/10.1098/rsta.1993.0059>
- [8] Gerkema, T. (1998) A Note on the Effect of Finite Stokes-Layer Thickness in a Morphodynamic Stability Problem. In: Dronkers, J. and Scheffers, M., Eds., *Proceedings of 8th International Biennial Conference on Physics of Estuaries and Coastal Seas*, Balkema, 387-395.
- [9] Engelund, F. (1970) Instability of Erodible Beds. *Journal of Fluid Mechanics*, **42**, 225-244. <http://dx.doi.org/10.1017/S0022112070001210>
- [10] Blondeaux, P. and Vittori, G. (1994) Wall Imperfections as a Triggering Mechanism for Stokes-Layer Transition. *Journal of Fluid Mechanics*, **264**, 107-135. <http://dx.doi.org/10.1017/S0022112094000601>
- [11] Hulscher, S.J.M.H., De Swart, H.E. and De Vriend, H.J. (1993) The Generation of Offshore Tidal Sand Banks and Sand Waves. *Continental Shelf Research*, **13**, 1183-1204. [http://dx.doi.org/10.1016/0278-4343\(93\)90048-3](http://dx.doi.org/10.1016/0278-4343(93)90048-3)
- [12] Huthnance, J. (1982) On One Mechanism Forming Linear Sand Banks. *Estuarine, Coastal and Shelf Science*, **14**, 79-99. [http://dx.doi.org/10.1016/S0302-3524\(82\)80068-6](http://dx.doi.org/10.1016/S0302-3524(82)80068-6)
- [13] Luque, D., Robles, A., Bravo, L.C., Nava, A., Rivera, M., López, A., Barnett, R., Morales, C. and Morales D. (2008) Monitoreo ambiental del Canal del Infiernillo, Territorio Comcaac (Seri). Modelo participativo comunitario basado en el conocimiento tradicional. Golfo de California. Ira Bienal del Programa de Ordenamiento Ecológico Marino del Golfo de California. La Paz, B.C.
- [14] Merifield, P.M., Marzolf, J.E. and Lamar, D.L. (1970) Marine Sand Waves in El Infiernillo Channel, Gulf of California, Final Report for the Office of Naval Research. Washington DC. Contract No N0014-69-C-0210. Task No 388093.
- [15] Marzolf, J.E. and Merrifield, P.M. (1969) Sand Waves Produced by Tidal Currents in the Gulf of California, Sonora, Mexico. Geological Society of America, Abstracts with Programs for 1969, Pallet 7 (Annual Meeting.), 143-144.
- [16] Torre Cosio, J. and Bourillón, L. (2000) Inventario y monitoreo del canal del Infiernillo para el comanejo de los recursos marinos en el territorio Seri, Golfo de California. Conservation Internacional, México, A.C. Informe Final SNIB-CONABIO Proyecto No L179, México DF.
- [17] Carbajal, N. and Backhaus, J.O. (1998) Simulation of Tides, Residual Flow and Energy Budget in the Gulf of California. *Oceanologica Acta*, **21**, 429-446. [http://dx.doi.org/10.1016/S0399-1784\(98\)80028-5](http://dx.doi.org/10.1016/S0399-1784(98)80028-5)

- [18] Montaña-Ley, Y., Peraza-Vizcarra, R. and Paez-Osuna, F. (2007) The Tidal Hydrodynamics Modeling of the Topolobampo Coastal Lagoon System and the Implications for Pollutant Dispersion. *Environmental Pollution*, **147**, 282-290. <http://dx.doi.org/10.1016/j.envpol.2006.07.007>
- [19] Montaña-Ley, Y., Carbajal, N. and Paez-Osuna, F. (2014) Bed Load Transport of Sediments and Morphodynamics in the Topolobampo Coastal Lagoon System, Mexico. *Journal of Coastal Conservation*, **18**, 55-67. <http://dx.doi.org/10.1007/s11852-013-0295-7>
- [20] Montaña-Ley, Y., Carbajal, N. and Páez Osuna, F. (2015) Sediment Dynamics in a Complex Coastal Lagoon System of the Gulf of California. *Journal of Coastal Conservation*, **19**, 295-306. <http://dx.doi.org/10.1007/s11852-015-0391-y>
- [21] Carbajal, N. (1993) Modelling of the Circulation in the Gulf of California. *Berichte aus dem Zentrum für Klima- und Meeresforschung. Reihe B: Ozeanograph*, **3**, 1-186.
- [22] Lancin, M. (1985) Geomorfología y génesis de las flechas litorales del canal del Infiernillo, Estado de Sonora. Universidad Nacional Autónoma de México. Instituto de Geología. *Revista*, **6**, 52-72.
- [23] Schuttelaars, H.M. and De Swart, H.E. (1996) An Idealized Long Term Morpho-Dynamic Model of a Tidal Embayment. *European Journal of Mechanics*, **15**, 55-80.
- [24] Van Rijn, L.C. (1993) Principles of Sediment Transport in Rivers, Estuaries and Coastal Seas. Acqua Publications, Amsterdam.
- [25] Montaña-Ley, Y. (2003) Long-Term Effects of the Bed-Load Sediment Transport on the Sea-Bottom Morphodynamics of the Colorado River Delta, México. PhD Thesis, University of Liege, Liege.
- [26] Montaña, Y. and Carbajal, N. (2008) Numerical Experiments on the Long-Term Morphodynamics of the Colorado River Delta. *Ocean Dynamics*, **58**, 19-29. <http://dx.doi.org/10.1007/s10236-007-0129-y>
- [27] Carbajal, N. and Montaña, Y. (1999) Growth Rates and Scales of Sand Banks in the Colorado River Delta. *Ciencias Marinas*, **25**, 525-540.
- [28] Carbajal, N. and Montaña, Y. (2001) Comparison between Predicted and Observed Physical Features of Sandbanks. *Estuarine Coastal and Shelf Science*, **52**, 435-443. <http://dx.doi.org/10.1006/ecss.2000.0760>
- [29] Dyer, K.R. (1986) Coastal and Estuarine Sediment Dynamics. John Wiley & Sons, Chichester, 342 p.
- [30] Fredsøe, J. and Deigaard, R. (1992) Mechanics of Coastal Sediment Transport. World Scientific.
- [31] De Vriend, H.J., Capobianco, M., Cheshire, T., De Swart, H.E., Latteux, B. and Stive, M.J.F. (1993) Approaches to Long Term Modeling of Coastal Morphology: A Review. *Coastal Engineering*, **21**, 225-269. [http://dx.doi.org/10.1016/0378-3839\(93\)90051-9](http://dx.doi.org/10.1016/0378-3839(93)90051-9)
- [32] Komarova, N.L. and Hulscher, S.J.M.H. (2000) Linear Instability Mechanism for Sand Wave Formation. *Journal of Fluid Mechanics*, **413**, 219-246. <http://dx.doi.org/10.1017/S0022112000008429>
- [33] Németh, A., Hulscher, S. and van Damme, R.M.J. (2001) Numerical Simulation of Sand Wave Evolution in Shallow Shelf Seas. *Coastal Dynamics*, **1**, 1048-1057. [http://dx.doi.org/10.1061/40566\(260\)107](http://dx.doi.org/10.1061/40566(260)107)
- [34] Németh, A.A., Hulscher, S.J.M.H. and Van Damme R.M.J. (2006) Simulating Offshore Sand Waves. *Coastal Engineering*, **53**, 265-275. <http://dx.doi.org/10.1016/j.coastaleng.2005.10.014>
- [35] Carbajal, N. and Gaviño-Rodríguez, J.H. (2013) Taylor Problem and Morphodynamics. *Estuarine Coastal and Shelf Science*, **129**, 173-179. <http://dx.doi.org/10.1016/j.ecss.2013.05.017>
- [36] Berthot, A. and Pattiaratchi, A. (2006) Mechanisms for the Formation of Headland-Associated Linear Sandbanks. *Continental Shelf Research*, **26**, 987-1004. <http://dx.doi.org/10.1016/j.csr.2006.03.004>
- [37] Schramkowski, G.P., Schuttelaars, H.M. and De Swart, H.E. (2002) The Effect of Geometry and Bottom Friction on Local Bed Forms in a Tidal Embayment. *Continental Shelf Research*, **22**, 1821-1833. [http://dx.doi.org/10.1016/S0278-4343\(02\)00040-7](http://dx.doi.org/10.1016/S0278-4343(02)00040-7)



Submit or recommend next manuscript to SCIRP and we will provide best service for you:

Accepting pre-submission inquiries through Email, Facebook, LinkedIn, Twitter, etc.

A wide selection of journals (inclusive of 9 subjects, more than 200 journals)

Providing 24-hour high-quality service

User-friendly online submission system

Fair and swift peer-review system

Efficient typesetting and proofreading procedure

Display of the result of downloads and visits, as well as the number of cited articles

Maximum dissemination of your research work

Submit your manuscript at: <http://papersubmission.scirp.org/>