

1 Article

# 2 Changes of Reference Evapotranspiration and Its 3 Relationships to Dry/Wet Conditions Based on 4 Aridity Index in Songnen Grassland, Northeast 5 China

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15 **Abstract:** Reference evapotranspiration ( $ET_0$ ) plays an irreplaceable role in regional dry/wet  
16 conditions under the background of climate change. Based on the FAO Penman-Monteith method  
17 and daily climate variables,  $ET_0$  was calculated for 22 stations in and around Songnen Grassland,  
18 northeast China, during 1960-2014. The temporal and spatial variations of  $ET_0$  and precipitation ( $P$ )  
19 were comprehensively analyzed at different time scales by using the Mann-Kendall test, Sen's slope  
20 estimator, and linear regression coupling with break trend analysis. Sensitivity analysis was used  
21 to detect the key climate parameter attributed to  $ET_0$  change. Then, the role of  $ET_0$  in regional  
22 dry/wet conditions was discussed by analyzing the relationship between  $ET_0$ ,  $P$  and aridity index  
23 (AI). Results shown a higher  $ET_0$  in the southwest and a lower in the northeast, but  $P$  was opposite  
24 to that of  $ET_0$ . Evidently decreasing trend of  $ET_0$  at different time scales was detected in almost the  
25 entire region, and the significant trend mainly distributed in the eastern, northeastern and central.  
26 For the whole region, sensitivity analysis indicated decreasing trend of  $ET_0$  was primarily attributed  
27 to relative humidity and maximum air temperature. The positive contribution of increasing  
28 temperature rising to  $ET_0$  was offset by the effect of significantly decreasing relative humidity, wind  
29 speed and sunshine duration. In addition, the value of  $ET_0$  shown higher in drought years and lower  
30 in wet years.

31 **Keywords:** reference evapotranspiration; climatic change; drought/wet; Songnen Grassland

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## 33 1. Introduction

34 Climate change becomes an indisputable fact, and it may accelerate hydrological cycle and  
35 redistribute global water resources [1]. Researches on climate change were identified not only in  
36 isolated temperature or precipitation, but also on integrated parameters [2], like reference  
37 evapotranspiration ( $ET_0$ ).  $ET_0$  is one of the vital components of hydrological cycle and controls energy  
38 and mass exchange between terrestrial ecosystems and atmosphere [2, 3], influenced by many factors  
39 including climate factors, crop factors, environmental conditions and management [4]. Changes in  
40  $ET_0$  would affect agricultural production, water resource programming and irrigation scheme. Under  
41 present global warming and climate change conditions, to identify the spatial and temporal variation  
42 and to determine the dominant climatic variables affecting  $ET_0$  trends are significant for revealing the  
43 impacts of climate change on hydrologic cycle. In addition, it can be helpful in determining  
44 appropriate adaptation measures for mitigating the potential damage from climate change bad effects

45 [5-7]. However, relationship between changing  $ET_0$  and dry/wet tendency is not quite clear yet, but  
46 it is crucial for water resource management.

47 Studies that investigate changes in climate factors have yielded mixture of results and  
48 conclusions about the trends of  $ET_0$  for specific locations during last decades. Contrary to the general  
49 expectations that increase in temperature would lead to an increase in evapotranspiration, some  
50 previous studies concluded that evaporation have diminished in the last decades [2]. In the upper  
51 and mid-lower Yangtze River basin, Wang, et al. [1] reported decreasing trend in  $ET_0$  during 1961-  
52 2000 based on daily data of 115 meteorological stations. Irmak, et al. [8] found over 116-year period  
53 there was a significant decreasing in  $ET_0$  in the Platter river basin, USA. The same decline trends also  
54 found in other regions throughout the world such as Canadian, New Zealand and India [9-11].  $ET_0$   
55 has identified an increasing trend in some regions for the same period, however, such as Iran,  
56 Northern Eurasia, and parts of Romania [12-14].

57 Being a typical farming-pastoral ecotone, located in the central part of northeast China, Songnen  
58 Grassland have experienced highly spatial and temporal variability in  $ET_0$  and precipitation. This  
59 makes the management of water resources difficult in the region. However, to the best of our  
60 knowledge, there was no comprehensive study of the relationships between changes in  $ET_0$  and  
61 dry/wet conditions has been done, and especially, the sensitivity of  $ET_0$  to climatic variables has not  
62 been done in this region. Water is the lifeline for the socioeconomic development of Songnen  
63 Grassland, because agriculture and animal husbandry, which heavily depend on precipitation and  
64 irrigation, are the pillar industry in the area. Therefore, understanding  $ET_0$  trend and its role in  
65 regional dry/wet conditions are important to address water shortage in this region and give a  
66 scientific basis for regional water resources management.

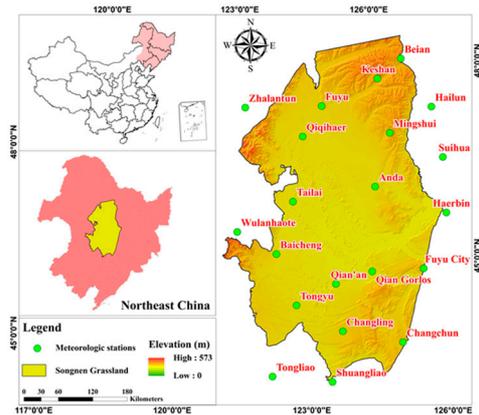
67 In this case, the objectives of this study were (1) to evaluate spatial distribution of the  $ET_0$  and  
68 precipitation at different time scales in Songnen Grassland over the period of 1960-2014; (2) to  
69 investigate temporal variability of  $ET_0$  by using the Mann-Kendall test and liner regression coupling  
70 with break trend analysis and the slopes of trend lines using the Sen's slope estimator; (3) to analyze  
71 temporal trends of the climatic parameters needed to calculate  $ET_0$ , and the sensitivity and its trends  
72 of  $ET_0$  to the climatic parameters by sensitivity analysis method; (4) to explore the role of changing  
73  $ET_0$  in regional drying or wetting conditions based on AI index. Results of this study will improve  
74 our understanding the impact of climate change on hydrological process and agriculture irrigation  
75 management.

## 76 2. Materials and Methods

### 77 2.1. Study area

78 Songnen Grassland is located in central northeast China. It lies from 43°30' N to 48°05' N and  
79 from 122°12' E to 126°20' E, and covers an area of approximately 22,350,000 km<sup>2</sup> (Fig. 1). Generally, it  
80 is distributed in the meadow steppe belt of China and is the important grassland in Eurasian steppe  
81 zone. The region is more dominated by temperate continental monsoon climate, with four distinct  
82 seasons: quite dry, windy springs, warm, rainy summers, sunny, mild autumns, and often long,  
83 freezing and dry winters. Mean annual temperature ranges from 1.9 °C in the southwest to 6.2 °C in  
84 the northeast region, while the mean annual precipitation amount varies from 350 mm in the  
85 southwest to 500 mm in the northeast region. Meanwhile, the mean annual amount of evaporation is  
86 roughly two or three times than precipitation.

87 Songnen Grassland forms a typically agricultural area at the east and an agro-pastoral transition  
88 zone at the west region, determined by various physical geographical features and regional climatic  
89 differences. The current ecological environment tend to be deteriorated due to the recent and ongoing  
90 climate change and land use change in the region. Since the agriculture and animal husbandry  
91 development are the main priorities within the general economic strategy of the region, investigation  
92 on changes in  $ET_0$  and its role in regional dry/wet conditions are needed for agriculture managers  
93 and stakeholders.



94

95 **Figure 1.** Location of Songnen Grassland in China and meteorological stations considered.96 *2.2. Climate data and quality control*

97 Climate data from 21 meteorological observatory stations is provided by the National  
 98 Meteorological Information Centre of China, including daily observations of maximum air  
 99 temperature (Max T, °C), minimum air temperature (Min T, °C), average air temperature (Ave T, °C),  
 100 average relative humidity (Ave RH, %), wind speed (Win S, m/s), sunshine duration (Sun H, h) and  
 101 precipitation (P, mm), for the period of 1960-2014. Regional of seasonal, growing season of vegetation  
 102 (from April to October) and annual values of these climatic variables are then calculated by the  
 103 weighted average or sum method. The weight of every station is obtained by Thiessen polygon  
 104 method which assigns weight in proportion to the study area that is closest to that station. Thermic  
 105 seasons are considered as winter that contains December, January and February; spring: March, April  
 106 and May; summer: June, July and August; and autumn: September, October and November.

107 The chosen meteorological stations are showed in Figure 1. They are all distributed inside or  
 108 adjacent the study area in order to cover the entire region. All selected stations have good-quality  
 109 data and meet the QA/QC requirements, and the missing data was substituted with the  
 110 corresponding long-term mean value.

111 *2.3. Methods*112 *2.3.1. Calculation of Reference Evapotranspiration (ET<sub>0</sub>)*

113 The FAO-56 Penman-Monteith (FAO-PM) equation was recommended as the sole and global  
 114 standard method for ET<sub>0</sub> calculation [4, 15]. The method had been widely verified its accuracy and  
 115 reliability under various climatologic zones around the world [16-18]. Accordingly, the FAO-PM  
 116 method was used to estimate daily ET<sub>0</sub> in this study, and subsequently seasonal, growing season and  
 117 annual ET<sub>0</sub> values were derived from daily values. The FAO-PM equation was expressed as:

$$118 \quad ET_0 = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{(T + 273)} U_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34U_2)} \quad (1)$$

119 where ET<sub>0</sub> is reference evapotranspiration (mm/d), Δ is the slope of the saturation vapour pressure  
 120 curve at a given air temperature (kPa/°C), R<sub>n</sub> is the net radiation at the crop surface (MJ/(m<sup>2</sup>•d)), G  
 121 is the soil heat flux density (MJ/(m<sup>2</sup>•d)), γ is the psychrometric constant (kPa/°C), T is the mean  
 122 daily air temperature at 2 m height (°C), U<sub>2</sub> is wind speed at 2 m height (m/s), e<sub>s</sub> is saturation  
 123 vapour pressure (kPa), e<sub>a</sub> is actual vapour pressure (kPa), and (e<sub>s</sub> - e<sub>a</sub>) is the saturation vapour  
 124 pressure deficit (kPa).

125 R<sub>n</sub> is the difference between the incoming net shortwave radiation (R<sub>ns</sub>) and the net outgoing  
 126 longwave (R<sub>nl</sub>). R<sub>ns</sub> is calculated as

$$127 \quad R_{ns} = (1 - \lambda) R_s \quad (2)$$

128 where  $R_s$  is the incoming solar radiation (MJ/(m<sup>2</sup>•d)) and  $\lambda$  (=0.23) is the albedo of the hypothetical  
 129 grass reference crop (dimensionless). And it was estimated based on sunshine duration record  
 130 according to the calibration equation by Croitoru, et al. [2]. At the same time,  $R_{nl}$  is given by

$$131 \quad R_{nl} = \sigma \left[ \frac{T_{max,K}^4 + T_{min,K}^4}{2} \right] \left( 0.34 - 0.14 \sqrt{e_a} \right) \left( 1.35 \frac{R_s}{R_{so}} - 0.35 \right) \quad (3)$$

132 where  $\sigma$  is Stefan–Boltzmann constant (=4.903×10<sup>9</sup> MJ/(K<sup>4</sup>•m<sup>2</sup>•d));  $T_{max,K}$  is the maximum absolute  
 133 temperature during the 24-h period (K=°C+273.16);  $T_{min,K}$  is the minimum absolute temperature  
 134 during the 24-h period (K=°C+273.16); and,  $R_{so}$  is the clear - sky radiation (MJ/(m<sup>2</sup>•d)). All the  
 135 variables in Eq. (1) were calculated using the standard procedure outlined by Allen, et al. [15].  
 136 Calculation of  $ET_0$  was based on CROPWAT 8.0 software developed by FAO during 1960-2014.

### 137 2.3.2. Trend analysis

138 To calculate trends, six data sets including one annual, four seasonal and one growing season  
 139 were conducted for each meteorological station. In this study, the Mann-Kendall (MK) test, a non-  
 140 parametric test, recommended by the World Meteorological Organization, was used to detect trends  
 141 [19, 20]. The MK test had been widely used for trend detecting in hydrologic and climatic research  
 142 owing to it does not need to conform any distribution form for the data and can allow the missing  
 143 data [21-23]. Besides, Sen's slope estimator was used to measure the magnitude of the trend [24]. The  
 144 MK test and Sen's slope estimator calculations for various time series of SPI and  $ET_0$  were performed  
 145 using the Excel-based template MAKESENS 2.0 beta, developed by researchers of Finnish  
 146 Meteorological Institute [17].

### 147 2.3.3. Sensitivity analysis

148 Sensitivity analysis was a quantitative description method of the important degree of input  
 149 variables to the output [25]. In present study, it was performed to evaluate the effect of climatic  
 150 variables on  $ET_0$ . Because of the different approaches used in conducting  $ET_0$ , there were no standard  
 151 or common procedure for carrying out the sensitivity analyses on  $ET_0$  [16, 26]. This study applied the  
 152 sensitivity analysis method which has the advantages of simple procedures and clear outcome was  
 153 developed from Li and T J [27]. The measure  $S_x$  is the sensitivity of the FAO-PM method to a  
 154 meteorological parameter, defined as:

$$155 \quad S_{xij} = \left| \frac{ET_0 \langle 1.1x_{ij} \rangle - ET_0 \langle 0.9x_{ij} \rangle}{ET_0 \langle x_{ij} \rangle} \right| \quad (4)$$

156 where  $x_i$  is one of a meteorological parameters needed calculation of  $ET_0$ ,  $j$  is the parameter at year,  
 157 and  $x_{ij}$ ,  $\langle 1.1x_{ij} \rangle$  and  $\langle 0.9x_{ij} \rangle$  are the estimated  $ET_0$  when the parameter  $x_i$  equals its reference  
 158 value or is  $1.1x_{ij}$  and  $0.9x_{ij}$  at  $j$  year, respectively.

### 159 2.3.4. Aridity index (AI index)

160 Aridity was usually expressed as a comprehensive function of precipitation, temperature,  
 161 and/or potential evapotranspiration, and reflects the level of meteorological drought [28]. There are  
 162 many AI indexes have been proposed. Among these indexes, following Thornthwaite [29] and Huo,  
 163 et al. [3] defined the AI index as the ratio of difference between  $P$  and  $ET_0$  to  $ET_0$ . The definition can  
 164 express the arid degree in any arid or semiarid region, and can be understood as the dearth of water  
 165 availability at the surface and subsurface levels.

166 In this study, based on the previous study, the AI value was calculated as:

$$167 \quad AI = (ET_0 - P) / ET_0 \quad (5)$$

168 where  $ET_0$  is reference evapotranspiration, and  $P$  is precipitation. If AI be equal to or close to 1, it  
 169 indicate that there is no precipitation and aridity in the highest. In contrast, if the precipitation be  
 170 equal to or higher than  $ET_0$ , the AI will be equal to 0 or be negative. The AI values of growing season

171 at each station were calculated, and the average AI in the whole Songnen Grassland was expressed  
 172 by the arithmetic average from stations in and around the region.

### 173 2.3.5. Spatial interpolation

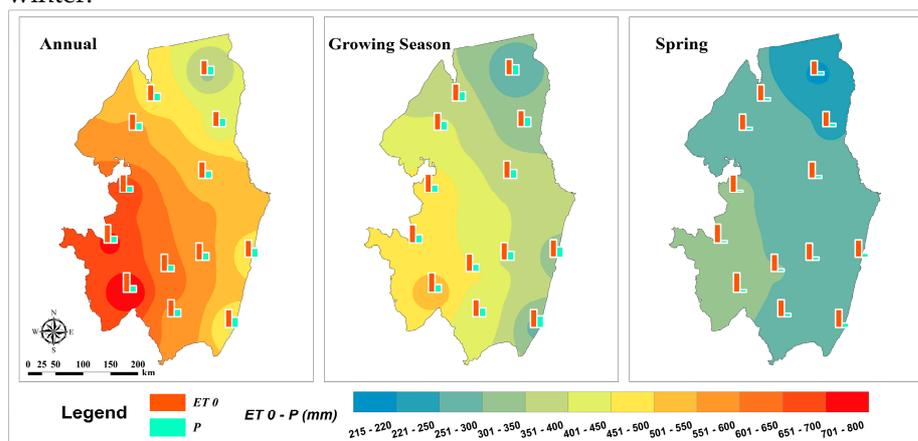
174 For analyzing of spatial patterns of trends and its magnitudes, the method of IDW was used to  
 175 the whole region. This method was a simple deterministic interpolation, which was extensively used  
 176 for mapping the spatial extent of climatic and hydrological point data [30]. All the procedures were  
 177 made using ArcGIS 10.2 software.

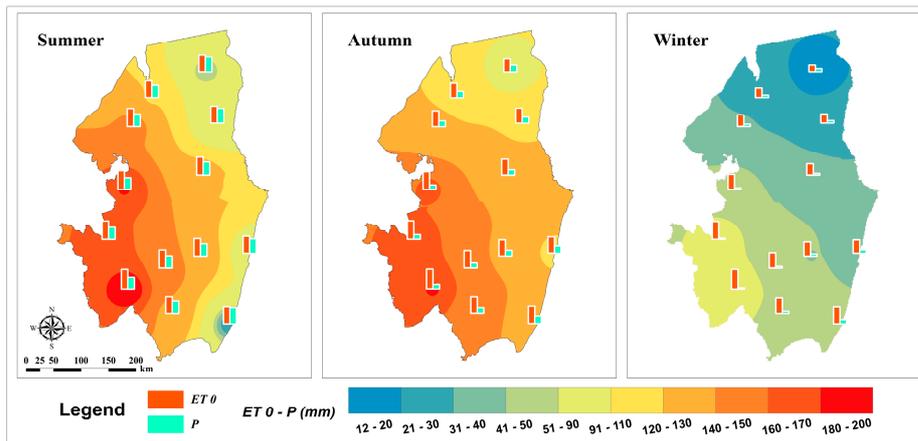
## 178 3. Results

### 179 3.1. Spatial distribution of $ET_0$ , $P$ and their difference over the period 1960-2014

180 The spatial distributions of  $ET_0$  and  $P$  at annual, seasonal and growing season from 1960 to 2014  
 181 were shown in Fig. 2, and their difference was displayed using the IDW interpolation as well. Fig.2  
 182 shows strong variability and marked difference between the northeastern ranges and the  
 183 southwestern region. From the visual inspection, it was observed that the spatial pattern of high and  
 184 low values of annual  $ET_0$ ,  $P$  and the difference all shows similarities to other time scales. The spatial  
 185 pattern of  $ET_0$  and the difference indicated higher values corresponding to the southwest and lower  
 186 values was distributed in the northeast areas, but precipitation was opposite, which the minimum  
 187 value was in western and the highest value was in eastern. Generally,  $ET_0$  value was more than  
 188 double the precipitation amounts at all of the time scales considered, especially during annual and  
 189 growing season.

190 For  $ET_0$ , the average of annual and growing season increased from northeast to southwest from  
 191 less than 850 and 750 mm respectively, to more than 1100 and 980mm respectively. And seasonally,  
 192 it presented that all of stations showed higher  $ET_0$  in summer, followed by spring, autumn and  
 193 winter. For  $P$ , the value of annual and growing season increased from western to eastern, ranges from  
 194 381 to 577 mm and from 352 to 511 mm respectively, and both the minimum was found in Tongyu  
 195 and the maximum was in Changchun. Seasonally, higher the amount of precipitation was exhibited  
 196 during summer, followed by autumn, spring and winter. For their difference, the maximum was  
 197 distributed in the southwest at every time scales (Fig.2), especially at annual, that the difference has  
 198 reached 700-800 mm in southwestern region. From seasonal perspective, the spring was experienced  
 199 the maximum value of the difference ranged from 215 to 350 mm, followed by summer, autumn and  
 200 winter.





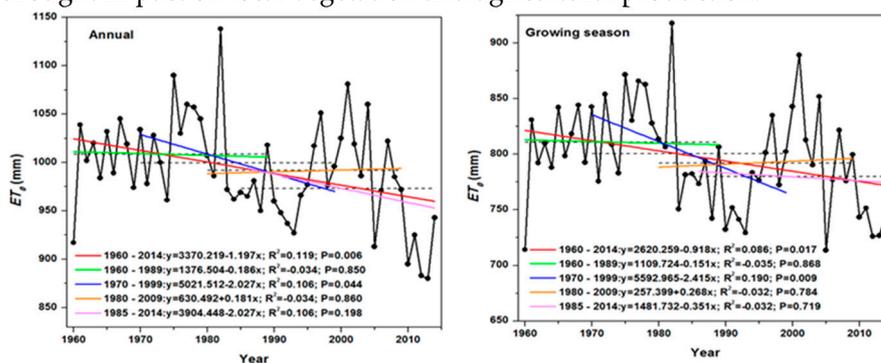
202  
203 **Figure 2.** Spatial distribution of  $ET_0$ ,  $P$  and their difference for the period 1960-2014.

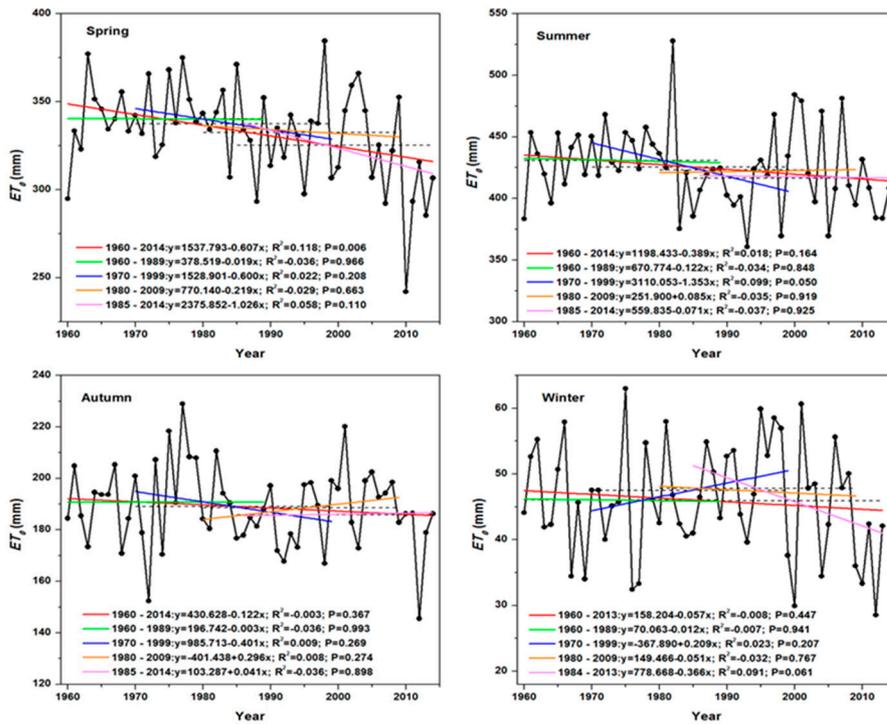
### 204 3.2. Temporal variations of $ET_0$

205 To calculate temporal trends of  $ET_0$  for different time scales, the linear regression coupling with  
206 break trend analysis as well as the MK test and the Sen's slope estimator were applied for the period  
207 1960-2014. The period of break trend analysis was chosen every 30 years, which was considered as  
208 the climatological normal according to the recommendation of WMO [31]. In this paper, the  
209 climatological normal ranged from 1960 to 1989, from 1970 to 1999, from 1980 to 2009 and from 1985  
210 to 2014, respectively. The values of the linear regression coupling with break trend analysis at 21  
211 were shown in Fig. 3, and those of the MK test and Sen's slope estimator were shown in Table 1.

212 Fig. 3 indicated an evidently decreasing trend in  $ET_0$  at every time scales, especially in annual,  
213 growing season and spring which all had passed the significance test at the 0.05 level. Besides, it was  
214 shown that almost all considered time scales had experienced decreasing trend of  $ET_0$  at every  
215 climatological normal with a rate ranging from -2.415 to -0.003 mm/a. The dropping at the sharpest  
216 rate was detected from 1970 and 1999, and the maximum was in growing season (-2.415 mm/a).  
217 However, it was noteworthy that the region was experienced an increasing trend of  $ET_0$  in all time  
218 scales during 1980-2009, with a rate ranging from 0.085 to 0.296 mm/a, except in spring and winter,  
219 although the increasing trend was not significant.

220 The results of trend analysis were almost identical with the nonparametric methods exhibited in  
221 Table 1. It was showed that  $ET_0$  had a decreasing trend in different time scales at the 21 stations inside  
222 and around the region. More than 90 % for the total number of analyzed stations at every time scales  
223 were reported decreasing trends except during autumn, and among them 50 % were statistically  
224 significant during annual, growing season and spring. Increasing trends were found in less than 10  
225 % of data sets, especially during autumn of approximately 38 %, but none of them were significant.  
226 In terms of magnitude of the trend, the highest decrease of  $ET_0$  was recorded in Haerbin (around the  
227 region) and Fuyu City (inside the region) at annual, with an average decrease rate of -4.45 and -  
228 2.86mm/a, respectively. On the basis of this results, there is no doubt that Songnen Grassland  
229 experienced an evidently decrease of  $ET_0$  in the last 55 years. The variations of  $ET_0$  rates may mitigate  
230 drought impact on local vegetation and agricultural production.





232

233

234

**Figure 3.** Linear trend and break trend analysis (for every 30 years) of  $ET_0$  from 1960 to 2014.

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**Table 1.** Results of the MK test and Sen's slope estimator for  $ET_0$  at different time scales.

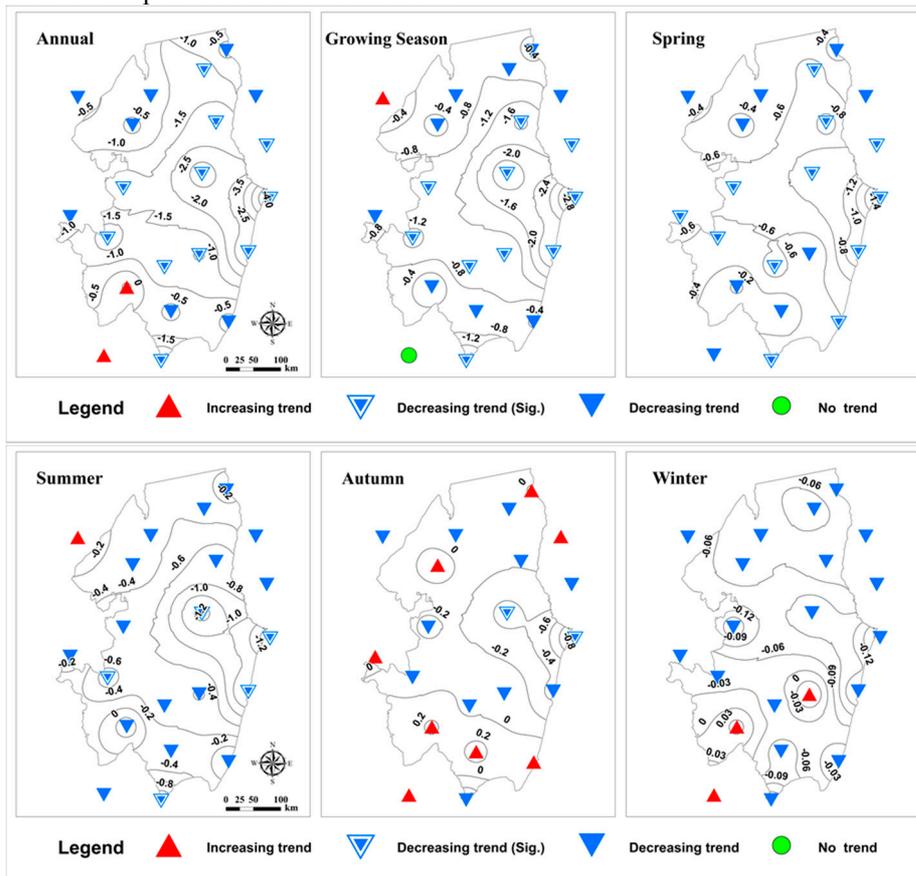
Station	Annual		Growing season		Spring		Summer		Autumn		Winter	
	Z	$\beta$	Z	$\beta$	Z	$\beta$	Z	$\beta$	Z	$\beta$	Z	$\beta$
Keshan <sup>a</sup>	-2.19*	-1.11	-1.89	-0.78	-2.37*	-0.65	-1.52	-0.38	-0.86	-0.11	-1.89	-0.08
Fuyu <sup>a</sup>	-1.47	-0.86	-1.15	-0.61	-1.48	-0.49	-0.94	-0.35	-0.44	-0.10	-0.66	-0.05
Qiqihaer <sup>a</sup>	-0.83	-0.41	-0.52	-0.30	-1.28	-0.35	-0.83	-0.27	0.60	0.10	-0.39	-0.04
Mingshui <sup>a</sup>	-2.82*	-2.01	-2.82*	-1.65	-2.52*	-0.83	-1.90	-0.74	-1.21	-0.20	-0.66	-0.04
Tailai <sup>a</sup>	-2.77*	-1.77	-2.19*	-1.14	-2.48*	-0.74	-1.45	-0.58	-1.73	-0.24	-1.22	-0.13
Anda <sup>a</sup>	-4.30*	-2.70	-4.17*	-2.23	-2.89*	-0.87	-3.22*	-1.25	-2.69*	-0.46	-1.42	-0.11
Baicheng <sup>a</sup>	-2.61*	-1.71	-2.24*	-1.29	-2.13*	-0.58	-2.08*	-0.66	-0.28	-0.06	-0.54	-0.03
Qian'an <sup>a</sup>	-3.27*	-1.50	-2.94*	-1.19	-2.86*	-0.67	-1.95	-0.56	-0.16	-0.02	-0.45	-0.04
QianGorlos <sup>a</sup>	-2.26*	-0.95	-2.27*	-0.83	-1.87	-0.42	-1.33	-0.38	-0.23	-0.04	0.32	0.03
Tongyu <sup>a</sup>	0.09	0.05	-0.12	-0.05	-0.65	-0.17	0.38	0.10	1.47	0.22	0.34	0.04
Changling <sup>a</sup>	-0.83	-0.40	-1.05	-0.41	-1.21	-0.22	-0.83	-0.23	1.81	0.26	-0.53	-0.08
FuyuCity <sup>a</sup>	-4.98*	-2.86	-4.52*	-2.33	-3.52*	-1.12	-3.51*	-1.21	-2.19*	-0.43	-1.58	-0.15
Changchun <sup>a</sup>	-0.74	-0.40	-0.68	-0.37	-2.04*	-0.60	-0.06	-0.03	1.02	0.17	-0.15	-0.01
Zhalantun <sup>b</sup>	-0.44	-0.14	0.29	0.15	-1.31	-0.30	0.83	0.18	-0.58	-0.10	-0.90	-0.07
Wulanhaote <sup>b</sup>	-1.71	-0.83	-1.23	-0.62	-2.24*	-0.70	-0.29	-0.10	0.12	0.02	-0.56	-0.05
Tongliao <sup>b</sup>	0.83	0.66	0.00	0.00	-0.83	-0.26	-0.16	-0.05	1.93	0.39	0.89	0.15
Beian <sup>b</sup>	-1.03	-0.40	-0.63	-0.24	-1.64	-0.33	-0.38	-0.08	0.04	0.01	-1.01	-0.04
Hailun <sup>b</sup>	-0.89	-0.53	-1.12	-0.58	-1.33	-0.39	-0.95	-0.29	0.41	0.08	-0.37	-0.02
Suihua <sup>b</sup>	-2.50*	-1.47	-2.69*	-1.30	-2.47*	-0.79	-1.78	-0.59	-1.00	-0.20	-0.61	-0.02
Haerbin <sup>b</sup>	-5.36*	-4.45	-5.11*	-3.39	-4.59*	-1.64	-3.48*	-1.50	-4.68*	-1.04	-1.72	-0.19
Shuangliao <sup>b</sup>	-2.86*	-1.91	-2.82*	-1.54	-2.19*	-0.52	-2.89*	-0.89	-2.03*	-0.32	-0.89	-0.13

Region <sup>a</sup>	-2.23*	-1.28	-2.04*	-1.01	-2.10*	-0.59	-1.48	-0.50	-0.37	-0.07	-0.68	-0.05
Region <sup>b</sup>	-1.74	-1.13	-1.66	-0.94	-2.07*	-0.62	-1.14	-0.41	-0.73	-0.14	-0.65	-0.05
Region <sup>a,b</sup>	-2.83*	-1.32	-2.57*	-1.04	-2.41*	-0.59	-1.77	-0.53	-0.65	-0.10	-0.45	-0.06

236 The Z value of more than 1.96 represent significant upward trend, while values less than -1.96 show a decreasing  
 237 trend at  $\alpha < 0.05$ . a represents station in Songnen Grassland; b represents station around Songnen Grassland;  
 238 a,b represents station in and around Songnen Grassland; \* represents significant trend at the 0.05 level.

### 239 3.3. Spatial variations of $ET_0$

240 The spatial variations of  $ET_0$  at different time scales from 1960 to 2014 were displayed in Fig. 4.  
 241 All considered time scales had experienced similar changing pattern, which presented a majority of  
 242 time series of  $ET_0$  have witnessed downward trends in the whole region, except during autumn. The  
 243 most stations with significant decreasing trend of  $ET_0$  were mainly distributed in the eastern,  
 244 northeastern and central region frequently occurred at annual, spring and growing season. These  
 245 region were distributed the highest decrease rate of  $ET_0$  ranged from -0.2 to -4.0 mm/a as well.  
 246 However, the positive changes were mainly recorded in the southwestern and southern regions at  
 247 autumn and winter, but they were statistically insignificant with the increasing rate ranging from  
 248 0.01 to 0.40 mm/a (Table 1). Overall, the spatial changing pattern of  $ET_0$  in Songnen Grassland is that  
 249 a significant downward trend across the central part region with a gradually reduced intensity from  
 250 the eastern parts to the western areas.



251

252

253 **Figure 4.** Spatial distribution of trends and its magnitude in  $ET_0$  over the period 1960-2014.

### 254 3.4. Changes in the climatic parameters

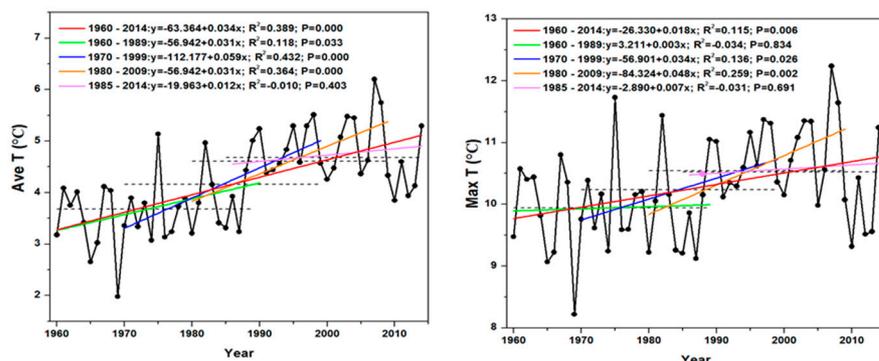
255 The changes in basic climatic parameters that played a considerable role in  $ET_0$  change have been  
 256 investigated on 21 stations across Songnen Grassland during 1960-2014. With the increase of air  
 257 temperature (Ave T, Max T and Min T), region averaged annual of Ave RH, Sun H and Win S all  
 258 significantly decreased during the study period (Fig. 5). All the climatic parameters experienced

259 consonant changes at every climatological normal. All the increase trends of Ave T, Max T and Min  
 260 T and the decrease trends of Sun H and Win S had passed the significance test at 0.01 level, while the  
 261 Ave RH had passed the significance test at 0.05 level.

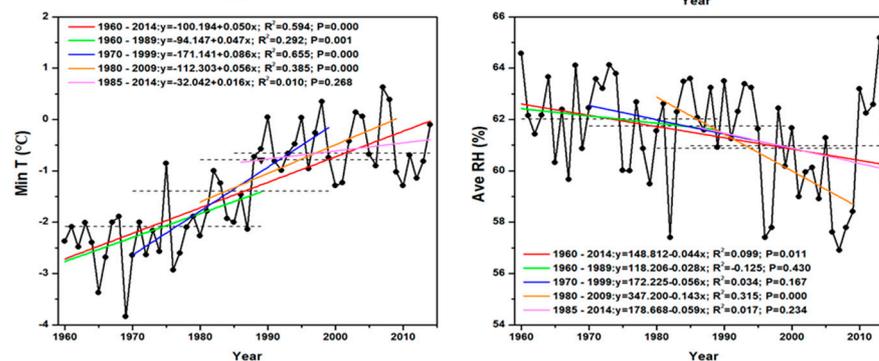
262 The MK test, then, was used for detecting the statistical significant trend at 21 stations in various  
 263 time scales, and the results were shown in Fig. 6. The percentage value of the statistical significant  
 264 increasing ( $\alpha = 0.05$ ) was 100 % for Ave T and Min T at all considered time scales, except during  
 265 winter 95 % in Min T and more than 70 % in Ave T. In the case of Sun H, all of the series illustrated  
 266 downward trends, especially at annual, among them between 52 % and 71 % were statistically  
 267 significant. As for Win S, significant decreasing trends were found in roughly 100 % at all the  
 268 analyzed locations. For Ave RH, an overwhelming majority of data series have seen negative trends,  
 269 but in spring the positive trends were detected more than 50 %.

270

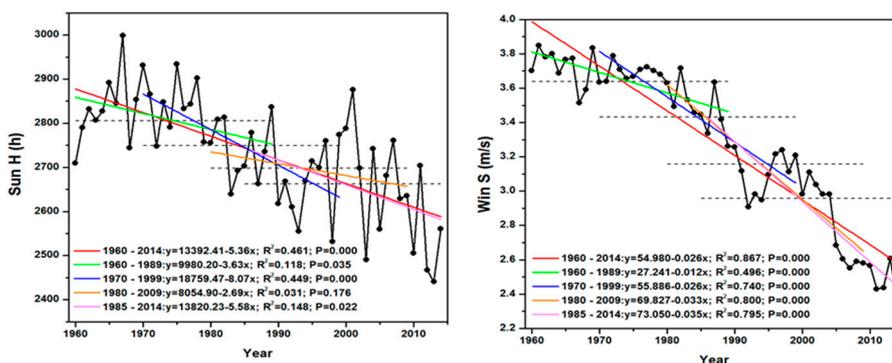
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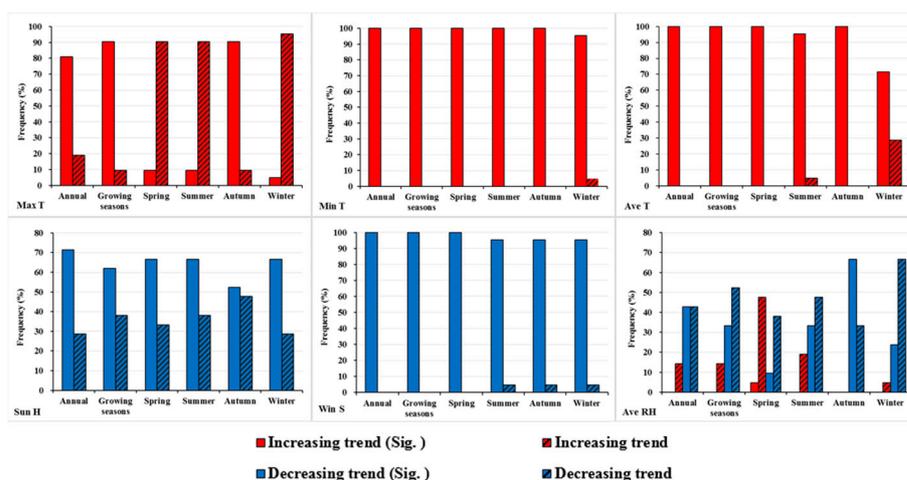
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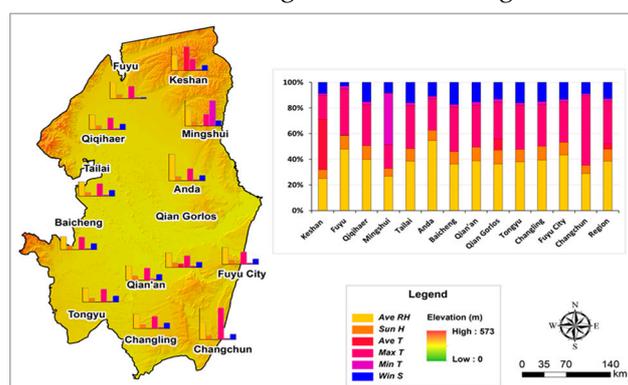
274 **Figure 5.** Linear trend and break trend analysis (for every 30 years) of climatic parameters from 1960 to  
 275 2014.



276

277 **Figure 6.** Frequency of climatic parameters trends in Songnen Grassland over the period 1960-2014 (%).278 *3.5. Attribution of climatic variables*

279 Results of the sensitivity analysis for annual  $ET_0$  to the climatic variables were performed in Fig.  
 280 7 and Table 2 for the 13 stations (inside the Songnen Grassland). It can be seen from Fig. 7 that the  
 281 Ave RH and Max T were the most sensitive variables in change of  $ET_0$  over the whole region, followed  
 282 by Win S, Sun H, Min T and Ave T. However, there was a little difference in the northeast region at  
 283 Keshan and Mingshui stations, mainly because there have a higher elevation and distributed small  
 284 hills. The most sensitive climatic variables at this two stations was Ave T and Min T, respectively.  
 285 The summary of dominant climatic variables and the temporal trends of their sensitivities were given  
 286 in Table 2. Results indicated that Ave RH shown a significant decreasing trend except in Anda station  
 287 which shown a significant increasing trend, and trend of Max T was the same as Ave RH that  
 288 experienced significant downward trend at almost the whole region besides significant upward  
 289 trends in Changchun. For Win S and Min T, a majority of stations witnessed a significant increasing  
 290 trend. The Sun H shown the same temporal trend with Win S and Min T, but in the southwest region,  
 291 Tongyu and Changling, shown no significant decreasing trend. As for the Ave T, results shown it  
 292 had slight impact on changing  $ET_0$  in Songnen Grassland, except in Keshan. In short, the sensitivity  
 293 of Ave RH, Max T, Ave T and Min T experienced a significant decreasing trends, whereas that of Win  
 294 S and Sun H shown a significant increasing trends in the whole region.



295

296 **Figure 7.** Spatial distribution of average sensitivity in annual comparison of various climatic variables.297 **Table 2.** Summary of dominant climatic variables that caused changes in  $ET_0$  at the inside stations.

Station	Order and trend of importance for the climatic variables in $ET_0$											
Keshan <sup>a</sup>	Ave T	-5.55*	Ave RH	-3.54*	Max T	-3.59*	Win S	-0.81	Sun H	3.85*	Min T	0.94
Fuyu <sup>a</sup>	Ave RH	-2.38*	Max T	-3.17*	Sun H	2.06*	Min T	1.56	Win S	0.07	Ave T	1.10

Qiqihaer <sup>a</sup>	Ave RH	-4.56*	Max T	-4.15*	Win S	1.31	Sun H	2.06*	Min T	4.24*	Ave T	0.00
Mingshui <sup>a</sup>	Min T	-6.90*	Ave RH	-3.96*	Max T	-4.85*	Win S	0.64	Sun H	1.51	Ave T	0.00
Tailai <sup>a</sup>	Ave RH	-3.89*	Max T	-3.12*	Win S	4.02*	Sun H	1.29	Min T	3.66*	Ave T	0.00
Anda <sup>a</sup>	Ave RH	2.83*	Max T	-3.56*	Win S	0.20	Sun H	4.07*	Min T	0.40	Ave T	0.00
Baicheng <sup>a</sup>	Ave RH	-3.05*	Max T	-3.03*	Win S	1.19	Sun H	1.80	Min T	2.08*	Ave T	0.00
Qian'an <sup>a</sup>	Ave RH	-2.90*	Max T	-2.87*	Win S	2.53*	Sun H	0.28	Min T	1.68	Ave T	0.00
Qian Gorlos <sup>a</sup>	Ave RH	-6.27*	Max T	-4.66*	Win S	0.96	Sun H	3.53*	Ave T	-4.72*	Min T	2.31*
Tongyu <sup>a</sup>	Ave RH	-2.95*	Max T	-1.74	Win S	3.03*	Sun H	-0.15	Min T	1.06	Ave T	0.00
Changling <sup>a</sup>	Ave RH	-3.27*	Max T	-2.63*	Win S	4.01*	Sun H	-1.26	Min T	2.42*	Ave T	0.00
Fuyu City <sup>a</sup>	Ave RH	-5.62*	Max T	-4.37*	Win S	2.42*	Sun H	1.28	Min T	3.14*	Ave T	0.00
Changchun <sup>a</sup>	Max T	3.11*	Ave RH	-6.11*	Win S	1.97*	Sun H	3.47*	Min T	2.24*	Ave T	0.00
Region <sup>a</sup>	Ave RH	-4.07*	Max T	-4.18*	Win S	2.90*	Sun H	3.92*	Ave T	-6.94*	Min T	-7.93*

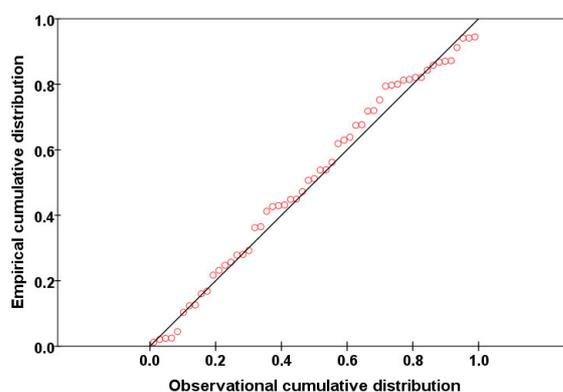
### 298 3.6. The role of $ET_0$ in regional dry/wet conditions

299  $ET_0$  accounts for more than 80 % of the total annual amount in growing season within a year,  
 300 besides, this period is critical for crops and vegetation growing which effect directly regional  
 301 socioeconomic. Therefore, this paper only considered what role  $ET_0$  may play in regional dry/wet  
 302 conditions during growing season. Regional trend and its magnitude of  $ET_0$ , AI and P were analyzed  
 303 on both the inside 13 stations and the total 21 stations, but here have presented only the results of the  
 304 former in Fig. 8, owing to the similarity in results.

305 Fig. 8 shown the interannual variations of growing season's  $ET_0$ , AI and P over Songnen  
 306 Grassland for the period 1960 to 2014. From the visual inspection,  $ET_0$ , AI and P have all experienced  
 307 the decreasing trends, but only the trend of  $ET_0$  passed the significance test. As for the fluctuation of  
 308 temporal evolution, results indicated that the fluctuating pattern of  $ET_0$  was similar to AI, but  
 309 opposite to that of P. Then, the simple regression equation is established as follows:

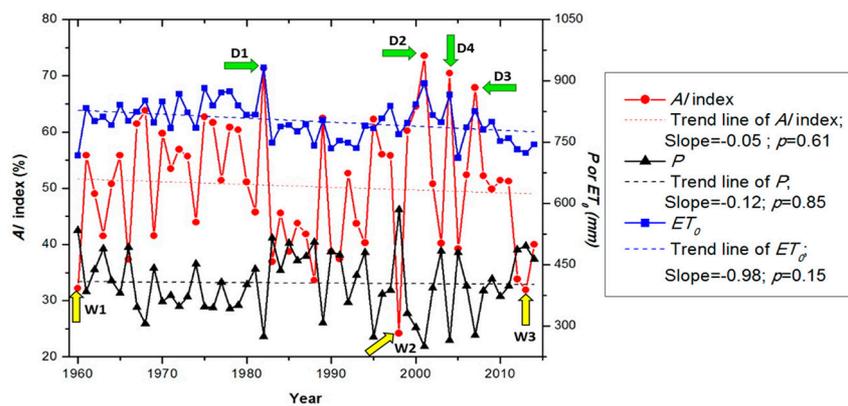
$$310 \quad AI = 63.81 - 0.125P + 0.046ET_0 \quad (R^2 = 0.98, P = 0.00) \quad (6)$$

311 Eq. (6) shown there was a positive correlation between AI and  $ET_0$  and a negative correlation between  
 312 AI and P. The regression residual standard error displayed in Fig.8. Thus, the upward trends of  $ET_0$   
 313 and AI were associated with downward trends of P and vice versa. This character was highly explicit  
 314 during drought and wet years in particular. As shown in Fig. 8, during the severe drought years D1  
 315 (1982), D2 (2001), D3 (2004) and D4 (2007), the values of  $ET_0$  and AI had reached or were about to the  
 316 maximum in the last decades, however, the values of P had nearly attained the minimum at the same  
 317 time. Similarly, the wet years such as W1 (1960), W2 (1998), and W3 (2013) had witnessed lower  
 318 values of  $ET_0$  and AI (Fig. 9).



319

320 **Figure 8.** P-P plot of regression residual standard error.



321

322 **Figure 9.** Comparison of temporal trend of  $ET_0$ , AI and P during 1960-2014. D, the drought year; W, the wet  
 323 year.

#### 324 4. Discussion

325 Songnen Grassland shows obviously spatial variations of  $ET_0$  rates from northeast to southwest  
 326 at different time scales (Fig. 2). It was influenced by both of the different regional climate and local  
 327 topography. In the north region, there was distributed some valleys and the stations located in  
 328 relatively high latitude area, while in the south, the topography was relatively flat with lower latitude  
 329 (Fig. 1). Moreover,  $ET_0$  rates over Songnen Grassland have been obviously decreasing over the last  
 330 decades (Fig. 3 and 4). The results were consistent with a decreasing trends in the lower reaches in  
 331 Taoer River basin of Northeast China investigated by Liang, et al. [32]. Huo, et al. also reported that  
 332 a decreasing trend in the arid area of northwest China during 1955-2008 [3]. However, as discussed  
 333 in introduction, some studies has identified an increasing trends in some regions for the last decades.  
 334 One of the reasons for inconsistent findings in  $ET_0$  trends is due to the fact that some studies on  
 335 climate change utilize divergent climatic parameters, potentially providing incomplete or artificial  
 336 trends and magnitudes in  $ET_0$  [8].

337 During the study period, observed variations in Max T, Min T and Ave T were similar to the  
 338 global pattern of increasing minimum, mean and maximum temperature [9], and the results of  
 339 analyzing on other climatic factors were consistent with the finding reported from Yunnan Province  
 340 in China [4]. Then, attribution analyses for changes in annual  $ET_0$  shown that Ave RH and Max T  
 341 were the most sensitive climate variables over the whole Songnen Grassland, which was in agreement  
 342 with Yin, et al. [33] studied at the whole China. However, Liu, et al. has been reported that Win S and  
 343 Sun H were the most sensitive factors in other part of China [11]. Thus, it can be concluded that  $ET_0$   
 344 has different responses to climate variables in different regions and climate conditions. What's more,  
 345 this study indicated the increase of  $ET_0$  induced by rising air temperature can be compensated by  
 346 reduced  $ET_0$  as significant decrease of Ave RH, Win S and Sun H. As a result, regional  $ET_0$  appeared  
 347 to show a declining trend. Remarkably, human activity should take upon some of the responsibility  
 348 about local climate change [28], such as GHG emissions and rapid urbanization, which should be  
 349 further studied in Songnen Grassland.

350  $ET_0$  variations and its response to regional dry/wet conditions are of great importance for crop  
 351 growing and natural vegetation [2, 34]. For this purpose, the present paper analyzed long term  
 352 variations of  $ET_0$ , AI and P at growing season during 1960-2014. Results shown that  $ET_0$  and AI  
 353 decreased as the P increased, especially during drought or wet years, and vice versa (Fig. 8). This was  
 354 consistent with Madhu, et al. [35] who reported that higher ET values were detected in the moderate  
 355 and severe droughts years. Contrary to the Eq. (5), notably, the trends of  $ET_0$ , AI and P were both  
 356 experiencing decrease trend during the study period in this paper. This could be due to the changes  
 357 of precipitation was not significant. However, the decreasing rate of  $ET_0$  and AI was higher than that  
 358 of P. Faced with this, the climate in Songnen Grassland gets slightly wetter from 1960-2014, and the  
 359 trend would be continued if the decreasing trend of  $ET_0$  persists. In short, regional climate change  
 360 brings positive influence for vegetation growth, agricultural production and ecological environment.

## 361 5. Conclusions

362 In this study, spatial distributions of  $ET_0$ , P and their difference were obtained at different time  
 363 scales. The temporal and spatial variations of  $ET_0$  were performed for 55 years of data from 22 stations  
 364 in and around Songnen Grassland, northeast China, during 1960-2014. Then, the interannual  
 365 variability of climatic variables was investigated during the study period, and sensitivity analysis  
 366 was conducted in this context. The role of  $ET_0$  in regional dry/wet conditions, ultimately, was  
 367 discussed based on analyzing relationships between  $ET_0$ , P and AI. The following conclusions can be  
 368 drawn from this study.

- 369 (1) Trend analysis of  $ET_0$  at different time scales shown an evidently decreasing trend in the last  
 370 55 years, especially in annual and spring. Break trend analysis shown that almost all  
 371 considered climatological normals had experienced the decreasing trend, with a range of -  
 372 2.415 to -0.003 mm per year. Spatial variations of  $ET_0$  indicated that most significant  
 373 decreasing trends were mainly distributed in the eastern, northeastern and central region  
 374 during annual, spring and growing season.
- 375 (2) The interannual variability of climatic parameters shown the annual Max T, Ave T and Min  
 376 T experienced significant increasing trend and significant decreasing trends were found for  
 377 Ave RH, Win S and Sun H. Ave RH was the dominant climate variable for the declining  
 378 annual  $ET_0$  over the complete region, followed by Max T, Win S, Sun H, Min T and Ave T.
- 379 (3) In general, the results of this study indicated that drought/wetness condition was getting  
 380 slightly wetter with decreasing  $ET_0$  during growing season. Regional climate drought has  
 381 been alleviated in recent several decades. The findings could be contribute to a better  
 382 planning and efficient use of agricultural water resources in Songnen Grassland.

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 387 Guo was responsible for data compilation; Feng Zhang and Mengmeng Wang was responsible for the data  
 388 processing and drawing; Qiyun Ma drafted the manuscript and all authors read and revised the final  
 389 manuscript.

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