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1 **Eco-hydrological requirements of dune slack vegetation and the implications of climate**
2 **change.**

3
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21
22
23 **Abstract**

24 Dune slacks are a seasonal coastal wetland habitat, whose plant assemblages and soil
25 properties are strongly linked to a fluctuating water table. Climate change is predicted to
26 cause major shifts in sand dune hydrological regimes, yet we know remarkably little about
27 the tolerance of these communities to change, and their precise hydrological requirements are
28 poorly quantified. Dune slack vegetation and soils were sampled within five vegetation types
29 across four west coast UK sites. Relationships between vegetation assemblages, and
30 parameters of soil development (moisture, loss on ignition, pH, KCl extractable ions) and
31 groundwater hydrological regime (annual maximum, minimum water levels and range,
32 duration of flooding) were established to define the environmental tolerances of different
33 communities. In multivariate analysis of the vegetation, the dominant gradient was
34 hydrological: dry to wet, followed by a secondary soil development gradient: young
35 calcareous organic-poor soils to acidic/neutral soils with greater organic matter contents.
36 Most measured hydrological and soil variables explained a significant proportion of observed
37 variation in species composition when tested individually, with the exception of soil nitrate
38 and soil calcium concentrations. Maximum water level was the key hydrological variable,
39 and soil moisture and soil pH were the key soil variables. All hydrological and soil
40 parameters together explained 22.5% of the total species variation. There were significant
41 differences in hydrological and soil parameters between community types, with only 40 cm
42 difference in mean annual minimum water levels (averaged over four years) separating the
43 wettest and the driest dune slack communities. Therefore, predicted declines in water level
44 exceeding 100 cm by 2080 are likely to have a major impact on the vegetation of these
45 priority conservation habitats.

46
47 **Keywords:** hydroecology; sand dunes; groundwater; global change; water table.

48 **1. Introduction:**

49 Sand dunes are rich in biodiversity due to the heterogeneity of habitat niches, and have both
50 considerable amenity value and a strategic function as coastal defence (e.g. Everard et al.
51 2010; Louisse and van der Meulen, 1991). Dune slacks are seasonally flooded humid
52 depressions between dune ridges, and are a priority habitat for rare species of conservation
53 importance, including orchids such as Liparis loeselii, Dactylorhiza praetermissa and
54 Dactylorhiza purpurella, lower plants like Petalophyllum ralfsii, a liverwort listed in Annex II
55 of the European Union Habitats Directive, and amphibians like Epidalea calamita which
56 breed in temporary pools (Smith, 2006). Many of these species are restricted to pioneer or
57 successional young vegetation communities (Davy et al., 2006; Rhind and Jones, 1999;
58 Sival et al., 1998) where competition is low.

59

60 Dune slacks form when bare sand is disconnected from seawater influence by the
61 establishment of a new dune front, or inland, where wind erosion scours bare sand down to
62 the water table or to the capillary wetted layer (Ranwell, 1960). Thus their formation and
63 subsequent plant and soil development are intimately connected to the dune groundwater
64 hydrological regimes. Large water table fluctuations are a feature of most slacks, and control
65 slack vegetation development. Variation of water levels occurs both within the year, typically
66 around 70 cm with a rapid rise in autumn and a gradual decrease from spring to summer
67 (Ranwell, 1959) and between years, depending on precipitation and evapotranspiration
68 balances (Ranwell, 1959).

69

70 Winter flooding, intensity of drought and persistence of waterlogging in the rooting zone
71 during the growing season are key environmental factors affecting vegetation, through
72 impacts on germination and productivity (Ernst, 1990; Grootjans et al., 1998). The timing and

73 duration of these events can alter inter/intraspecific competition, thus changing community
74 composition (Bossuyt et al., 2003; Bossuyt et al., 2005). Groundwater fluctuations also
75 control nutrient status: high water levels in slacks reduce the mineralisation of organic matter,
76 maintaining low nitrogen and phosphorous levels (Lammerts and Grootjans, 1997). The
77 chemistry of groundwater is important, and can vary considerably across a site (Jones et al.,
78 2006). In older de-calcified dune slacks, buffering action of carbonate-rich groundwater
79 allows the survival of basiphilous wetland plants (Grootjans et al. 1988, 1991; Sival et al.
80 1998; Van Dijk and Grootjans, 1993).

81

82 Most authors agree that species distribution and community structure across slacks are highly
83 correlated with groundwater levels (Grootjans et al., 1991, 1998; Jones and Etherington 1971;
84 Lammerts et al. 1998, 2001; Noest, 1994; Olf et al. 1993; Ranwell, 1960; Sival et al., 1998;
85 Van der Laan 1979; Willis, 1959b). The differentiation as wet or dry slacks and the
86 hydrological characteristics of different communities are well understood by ecologists (e.g.
87 Rodwell, 2000). Yet, despite this, the precise eco-hydrological requirements of these
88 communities are poorly, if at all, quantified in the UK. In the Netherlands there is a wealth of
89 ecohydrological knowledge on dune slacks (e.g. Witte et al. 2007; von Asmuth et al. 2012),
90 but slacks in the UK differ somewhat from those on the continent. In general they experience
91 higher rainfall, they are often present on narrower dune sites with greater potential for
92 groundwater influence from inland. In the majority of west coast sites, the sand parent
93 material is usually more calcareous, resulting in highly buffered systems with slower
94 decalcification rates, despite the higher rainfall. However, in north UK, slacks can be less
95 well buffered and higher rainfall leads to decalcification of surface soils and rapid organic
96 matter accumulation.

97

98 Previous attempts to define eco-hydrological requirements for UK dune slack communities
99 have been based on relatively small numbers of combined vegetation and hydrology records
100 (Ranwell 1959), and conducted at single sites (Jones, 1993; Ranwell, 1959). A further failing
101 is the short duration of hydrological records considered: two years or less, with little
102 understanding of longer term hydrological variability and whether the vegetation is in
103 equilibrium with hydrological conditions (Davy et al., 2010).

104

105 Dune slack habitats worldwide are under increasing anthropic pressure from water
106 abstraction, afforestation, urbanisation (Grootjans et al. 1998; Martinez et al., 2004; Provoost
107 et al., 2011), nitrogen deposition (Jones et al., 2004; Plassmann et al., 2008; Sival and
108 Strijkstra-Kalk, 1999) and from grass and scrub encroachment or exotic species invasion
109 (Martinez et al., 2004). In addition, changes in evapotranspiration due to vegetation change or
110 management will affect the water balance (Davy et al. 2010; Ford et al. 2012). An emerging
111 threat is climate change, which may shift the biogeographical range of dune slack species, but
112 which also alters the dune environment. Changes in precipitation and temperature affect
113 groundwater levels directly by altering the delicate balance between rainfall and
114 evapotranspiration, which controls recharge. Sea-level rise or shoreline erosion act indirectly
115 on groundwater levels by altering water table gradients (Clarke and Sanitwong Na Ayutthaya,
116 2010; Saye and Pye, 2007). Modelling of groundwater trends for dune slacks in a sand dune
117 system in north west England based on long term records, predicted a substantial lowering of
118 water levels of 1 to 3 metres over the next 90 years (Clarke and Sanitwong Na Ayutthaya,
119 2010). Physiological adaptations typical of plants growing on humid calcareous substrates
120 can make them less resilient to rapid habitat changes (Bakker et al., 2007; Grootjans et al.,
121 2004; Schat, 1984) and the rapidity of community shifts in response to variations of
122 groundwater regime is difficult to estimate (Noest, 1994; Van der Laan, 1979). The predicted

123 changes in water levels are so large that major changes in slack vegetation are probable, but
124 the outcomes will remain uncertain until we better understand the hydrological requirements
125 of these vegetation communities.

126

127 The aim of this investigation was therefore to improve our understanding of the relationships
128 between dune slack vegetation communities and the underlying hydrological and
129 biogeochemical controls, across a range of west coast UK sites, through the following steps:

130

131 1. Creating a network of co-located vegetation and hydrological monitoring locations,
132 maximising the use of previously unconnected long-term vegetation or hydrological data
133 records.

134 2. Using multivariate analysis to determine the principle environmental parameters governing
135 species assemblages in sampled dune slack communities.

136 3. Using hydrological data from a reasonably climatically stable four year period, to
137 characterise the hydrological and environmental requirements of each community type.

138 4. Interpreting those requirements in the light of predicted climate change impacts on dune
139 groundwater regimes.

140

141 Using this information, the following specific questions were postulated. Can vegetation
142 communities be distinguished according to hydrological regime? Are projected changes in
143 groundwater regime likely to have serious consequences for current dune slack assemblages?

144

145 **2. Methods:**

146 *2.1 Study sites*

147 Data were collected from four dune systems on the west coast of the UK (Table 1).
148 Newborough Warren, a dune system located on the SW corner of the Isle of Anglesey in
149 North Wales, UK, which contains all five of the UK slack communities. The sand dune area
150 has developed between two estuaries; the western part of the dunes has been forested with
151 Pinus nigra ssp. laricio commencing in 1948. Aberffraw, close to Newborough (~6km north
152 west), is a smaller system about 1 km wide and extending 3 km inland, enclosed within a
153 valley. Ainsdale Sand Dunes National Nature Reserve in the Sefton coast (Merseyside, UK),
154 the largest area of open dune landscape in England, where part of the dunes are also forested
155 (Pinus nigra ssp. laricio), and Whiteford Burrows National Nature Reserve, a spit dune
156 system on the south side of the Loughor estuary, Gower Peninsula, South Wales. Here the
157 dunes are bounded by saltmarsh on the east side and also partially afforested in the southeast.
158 All sites are grazed by rabbits and, within fenced areas, are grazed by domestic livestock
159 (Welsh mountain ponies, cattle and sheep). Climate in these locations can be defined as
160 oceanic, with mild winters and cool summers. Long-term annual rainfall (1971-2001) at the
161 sites varies from 828 to 1107 mm (Table 1); evapotranspiration in these dune slacks can vary
162 between 360 and 550 mm/year depending on slack wetness and grazing pressure (Stratford et
163 al., 2007).

164

165 [Table 1 about here]

166

167 *2.2 Vegetation communities*

168 In the UK, five dune slack vegetation communities have been described (Rodwell, 2000) with
169 a number of sub communities. These vary from younger base-rich slacks through to older,
170 decalcified communities, and range from wet to dry types. The main communities and
171 subcommunities described in this study are summarised in Table 2, together with a dry dune

172 grassland community (SD8), which forms the common dry end-point of the driest slack
173 community SD16.

174

175 *2.3 Hydrological data.*

176 Even though water table levels and vegetation are monitored in several dune systems in the
177 UK, their sampling locations rarely coincide. At Newborough there was an existing network
178 of piezometers in the dunes, with the oldest hydrological records dating to 1985. The majority
179 of wells were located in a single community type, SD14, therefore new piezometers were
180 installed in early 2010 in other vegetation types, and adjacent to permanent vegetation
181 quadrats established in 1987 (Plassmann et al. 2010). Additional transects of wells were
182 installed across natural vegetation gradients in three slacks. New piezometers were also
183 installed at Aberffraw, while existing piezometers were used at Ainsdale and Whiteford
184 Burrows (Table 1), giving a total of 58 useable piezometers across all sites.

185

186 Monthly manual measurements of groundwater levels were made in all 58 piezometers from
187 early 2010 for a period of 18 months. Older data at monthly resolution were available since
188 2006 for 16 piezometers at Newborough. Automatic logging water level recorders (DIVER®,
189 Schlumberger, The Netherlands) were located in some of these older piezometers, and
190 additional loggers installed in a selection of new piezometers, providing higher temporal
191 resolution and independent verification of water levels at these locations. However, for
192 statistical analysis, monthly resolution manual data were used from all piezometers to
193 maximise spatial coverage across the sites. A hydrological year commencing on 1st June,
194 during the summer water table decline was used as the basis for summary variables, in order
195 to reliably record minimum and maximum water table heights in the same year, since timing
196 of the rapid autumn re-wetting can vary considerably from year to year. The following

197 hydrological variables were extracted from monthly records for 2010 and for longer time
198 series where available: minimum water level, maximum water level, 10th and 90th percentiles
199 (in case of truncation of records), annual range (maximum minus minimum water level), an
200 estimated duration of flooding (assuming that where the water table was at or above the
201 ground surface flooding occurred for four weeks), annual median (only calculated from 2006
202 to 2009), seasonal (spring, summer, autumn, winter) averages and averages for spring months
203 (March, April, May and June).

204

205 [Table 2 about here]

206

207 *2.4 Vegetation survey*

208 Vegetation was surveyed in 1 m x 1 m quadrats around existing and newly established
209 piezometers. Quadrats were arranged as a cross around the piezometers, with an additional
210 quadrat placed in the most homogeneous part of the stand. Species occurrence was recorded
211 as cover values (using visual estimates of percentage cover); nomenclature follows Clapham
212 (1987) for vascular plants, since it is compatible with NVC community descriptions, and Hill
213 (1991–94) for bryophytes. Physiognomical parameters, such as vegetation height, bare
214 ground and litter cover, evidence of grazing, slope and aspect, were recorded for each
215 quadrat. The location of the centre of each quadrat was recorded using a Leica 1200
216 RTKGPS.

217

218 *2.5 Topographical resolution*

219 Elevation of the ground surface at each piezometer and each quadrat was also measured using
220 the Leica 1200 RTKGPS, with a vertical accuracy of ± 1 cm. This provided elevation relative
221 to the water table measured at nearby piezometers, used to derive adjusted hydrological

222 summary data for each quadrat. Field-testing showed that water table height varied by less
223 than 5 cm within 7 m of a piezometer and this was deemed acceptable variation without
224 adjustment. At greater distances groundwater depth was measured using an auger, or quadrats
225 were excluded from analysis.

226

227 *2.6 Soil survey and laboratory measurements*

228 At each quadrat a soil sample was collected, 5 cm in diameter to a depth of 15 cm. Samples
229 were stored in the dark at 5° C until processed. After measuring depth of the organic horizon,
230 large roots were removed and the samples thoroughly mixed. Ten grams of mixed fresh
231 sample were weighed and dried overnight at 105° C to measure moisture content. Loss On
232 Ignition was calculated by re-weighing the samples after 16 hours at 375° C. pH was
233 measured after 30 min equilibration of 10 grams dry soil stirred in 25 ml deionised water,
234 with pH electrode calibrated with standard solutions. Water-extractable soil nutrients were
235 sampled using a solution prepared with 5 grams of dry soil in 45 ml of ultra high purity water
236 (1:10 wt/vol), shaken for 24 hrs at 70 rpm in the dark followed by centrifugation at 5000 rpm
237 for 10 minutes, and finally filtered through 0.45 µm cellulose nitrate filters. Extracts were
238 analysed using a Metrohm ion chromatograph. Reference soils and duplicates were used for
239 quality assurance.

240

241 *2.7 Statistical analysis*

242 Groups of quadrats with similar vegetation were identified using cluster analysis in
243 TWINSpan (Two-Way INdicator SPecies ANalysis, Hill, 1979b). The UK National
244 Vegetation Classification (NVC) unit (Rodwell, 2000) for each of these clusters was assigned
245 using MATCH software (Malloch, 1998). In order to determine whether quadrats represented
246 ‘core’ good quality vegetation or ‘transitional’ vegetation, they were also assigned
247 individually to an NVC unit using MATCH. If the quadrat assignment matched that of its

248 parent cluster, it was classed as 'core', if it did not match, or was intermediate between two or
249 more communities it was classed as 'transitional'. Mean unweighted Ellenberg scores for
250 light (L), moisture (F), reaction/pH (R) and nutrients (N) (Ellenberg et al., 1991) were
251 calculated for each quadrat using UK Ellenberg indicator values for vascular plants and
252 bryophytes (Hill et al., 1999).

253

254 The relationship between the vegetation and environmental variables was explored through
255 indirect gradient methods using Principal Component Analysis (PCA) after testing the length
256 of the first gradient for linearity of response gradients (CANOCO 4.0, ter Braak and
257 Smilauer, 1998). The main hydrological and soil parameters were then used in a direct
258 gradient method using Canonical Correspondence Analysis (CCA) for a subset of 189
259 quadrats, excluding those too far from the piezometers or with missing data. Significance of
260 environmental variables in the CCA was assessed using Monte Carlo methods within
261 CANOCO.

262

263 Since the majority of piezometers had only one year of hydrological records, it was explored
264 whether piezometers with longer time series could be used to extrapolate longer-term
265 hydrological trends. Average annual values for the hydrological parameters minimum water
266 level, maximum water level and range were calculated for a four year sequence (2006 –
267 2009) for twelve piezometers located in the main dune slack community types. Relationships
268 between the longer-term data and the data for 2010 were established using linear regression,
269 and were all significant (Maximum water level, $R^2 = 89.2\%$, $p < 0.001$; Minimum water level,
270 $R^2 = 96.6\%$, $p < 0.001$; Range, $R^2 = 37.4\%$, $p = 0.009$). These equations were used to calculate
271 four-year averages for the wider set of piezometers and associated quadrats. Frequency
272 distributions of water levels for each community were calculated from mean and standard

273 deviation of annual median water levels calculated for all quadrats within each vegetation
274 community type, assuming a normal distribution.

275

276 Differences between vegetation clusters (or NVC units) for the main hydrological (minimum
277 and maximum water levels, and annual range) and other environmental variables (soil pH,
278 soil moisture and percentage Loss On Ignition, percentage bare ground and litter) were
279 assessed by Analysis Of Variance, performed on the 2010 variables using Minitab v. 16.

280 Summary hydrological data for each community were calculated from the four-year averages
281 predicted from regression analysis. Data transformation was necessary for almost all soil
282 elements, since data were not normally distributed (Kolmogorov-Smirnov test). Chloride,
283 phosphate, sulphate, nitrate, potassium and sodium were natural log transformed, ammonium
284 was square root transformed. Nitrite was excluded from the analysis due to excessive missing
285 values.

286

287

288 **3. Results:**

289 The initial cluster analysis allocated the 245 quadrats to seven dune slack and grassland units
290 of the UK National Vegetation Classification (Rodwell, 2000). To illustrate differences in the
291 hydrological regimes underlying the main communities, Figure 1 shows typical hydrographs
292 over four years from piezometers located in each of the main community types at
293 Newborough Warren. The Potentilla anserina-Carex nigra (SD17) and Salix repens-
294 Calliergon cuspidatum (SD15) communities experienced higher winter water tables and more
295 frequent surface flooding of 10/20 cm compared to the other units. The Salix repens-
296 Campylium stellatum community (SD14, only Bryum pseudotriquetrum-Aneura pinguis
297 SD14c subcommunity shown) showed less annual fluctuation with lower winter but higher
298 summer water table levels. The driest slack community Salix repens-Holcus lanatus (SD16)

299 had generally lower water tables throughout the year. The groundwater levels underlying dry
300 dune grassland Festuca rubra-Galium verum (SD8) were considerably lower throughout and
301 are shown for comparison.

302

303 [Figure 1 about here]

304

305 The PCA shows the pattern of variation in the quadrats, coded according to vegetation
306 community (Figure 2), with environmental variables (bottom right) overlain for
307 interpretation. The primary axis of variation corresponds to a wetness gradient: Ellenberg F,
308 all the hydrological variables and to a lesser degree soil moisture, were positively correlated
309 with axis 1 scores. The second axis corresponds to soil development, with Ellenberg N,
310 %LOI and NH_4^+ positively correlated whilst Ellenberg L and R and soil pH were negatively
311 correlated with axis 2 scores. The vegetation communities are arranged along these gradients
312 as follows. The wettest units are SD15b and SD14b, the former extending to older (higher
313 organic matter) and wetter soil conditions, the latter having lower organic matter
314 concentration indicative of younger sites. Quadrats clustering as SD14c have high scores for
315 Ellenberg L and R reflecting a species composition characteristic of an open, early
316 successional habitat with abundant heliophilous and basiphilous species growing on young
317 calcareous soils with a thinly developed organic layer. The dry slack community SD16 is
318 concentrated at the right hand side of the diagram indicative of low soil moisture whilst
319 variability along the second axis indicates that it can develop on substrates of varying ages.
320 The dry dune grassland SD8 quadrats are next to those of SD16, associated with older and
321 usually drier soils, although the two communities clearly have overlapping hydrological
322 tolerances on this part of the hydrological gradient. SD14d occupies a position intermediate
323 between the wet SD15b and dry SD16. SD17 is not as well defined as the other communities;

324 with only seven quadrats, however, it appears to occupy damper and more mature systems
325 associated with SD14d and SD15b.

326

327 [Figure 2 about here]

328 [Figure 3 about here]

329

330 The CCA (Figure 3) shows the influence of the measured environmental variables (Figure 3a)
331 on the position of all vegetation quadrats (Figure 3b) and of groups consisting only of core
332 vegetation (Figure 3c). The dominant axis relates to hydrology and the second axis to soil
333 development, as in the unconstrained gradient analysis (Figure 2). Communities occupy
334 broadly the same niches as seen in PCA, with some distinctions. The SD8 dry dune grassland
335 and the wet slack SD16 are more clearly separated along the wetness gradient, and
336 communities are more clearly distinguished along the soil development (secondary) axis.
337 However, a key component of Figures 3b and 3c is the very high degree of overlap of the
338 communities along the hydrological gradient represented by Axis 1.

339

340 Significance testing using Monte Carlo permutation showed that the model with all variables
341 was highly significant ($p < 0.001$). The first four axes explained 61.6 % of the species-
342 environment relationships, but explained a smaller proportion of the total species variance
343 (first four axes 22.5%). Most variables tested singly were highly significant (Table 3), with
344 annual maximum water level (6.3%) and soil moisture (6.3.%) explaining the greatest
345 proportion of the total species variance. Only soil available nitrate and calcium were not
346 significant explanatory variables. However, there was a high degree of correlation amongst
347 the hydrological variables and all hydrological variables together explained only 8.9%.
348 Hydrological and soil variables were largely orthogonal, with annual maximum water level,
349 soil moisture and soil pH explaining 12.9% of species variance. An extended model

350 combining key hydrological (maximum water level), soil (%moisture, soil pH) and soil
351 chemistry parameters (NH_4^+ , PO_4^-) increased the cumulative species variance to 14.6%.
352 Ellenberg indicator values were separately tested as explanatory variables. Ellenberg F
353 explained 10%, Ellenberg N an additional 7%, and a further 3% for Ellenberg R; when
354 Ellenberg L was included in a combined model, together they explained 21.2% of the
355 variance, similar to the proportion explained by the measured environmental variables.

356

357 Water extractable nutrients, presented as average ion concentration in six vegetation
358 communities (combined data from core and transitional vegetation quadrats), showed
359 significant differences between communities (Figure 4): calcium $p=0.05$, potassium $p=0.01$,
360 all other ions $p<0.001$. Sulphate and phosphate had higher concentrations in SD17 as did
361 nitrate. The other ions presented higher values in SD15b, except calcium, which was more
362 concentrated in SD14 subcommunities. The SD14 subcommunities all had lower
363 concentrations of nitrate, magnesium and potassium than the other communities.

364

365 [Figure 4 about here]

366 [Table 4 about here]

367

368 Hydrological parameters, the principal soil parameters and vegetation physiognomy were
369 summarised for the main vegetation communities (Table 4). Minimum and maximum water
370 levels differ significantly between the wetter slack communities (SD15b, SD14b), and the
371 drier communities (SD16, SD14d). Winter water levels of the two wettest units (SD15b and
372 SD14b) differ by around 20 cm from the slightly drier SD14c and SD14d whilst their summer
373 minimum is at least 10 cm higher. The dry slack community (SD16) has mean winter and
374 summer levels around 20-25 cm lower than the other slack types. The younger SD14c

375 subcommunity appears to be intermediate in character regarding water levels, but its annual
376 range (49 cm) is much lower than the others (~70 cm). Winter water tables in the dry dune
377 grasslands (SD8) are a further 25 cm lower on average, although the wettest examples in this
378 study have a winter maximum which overlaps much of the hydrological range of the dry
379 slack community SD16 (Figure 5). %LOI and pH reflect soil age, with the youngest (SD14c)
380 and older/wetter (SD15) communities differentiated, with the rest intermediate. The older,
381 slightly decalcified, SD17 community, although poorly represented in this study, was not
382 significantly different from SD15 in terms of hydrological and soil parameters, but diverged
383 in bare ground and litter, with >25% plant litter compared with <4% in the other
384 communities. Figure 5 shows the 67 and 90 percentile ranges for maximum winter water
385 levels and minimum summer water levels for quadrats with core vegetation assemblages
386 only. This shows clearly that the majority of core quadrats in the three main slack
387 assemblages in this study have hydrologically distinct summer and winter water levels.

388

389 [Figure 5 about here]

390

391 **4. Discussion:**

392 This study has shown that dune slack vegetation assemblages are associated with distinct
393 hydrological regimes, and with differing soil physical and chemical properties, in broad
394 agreement with studies in the Netherlands (e.g. Lammerts et al. 2001). The cluster analysis
395 and PCA axis scores demonstrated a clear separation of vegetation communities based on
396 their species composition with Ellenberg F scores suggesting axis 1 was strongly linked to
397 soil moisture. The majority of environmental variables tested through direct gradient methods
398 (CCA) were significant in explaining species variation. However, although the proportion
399 explained by the species-environment relationships is similar to published studies, e.g. 58 %

400 for the first two axes in Lammerts et al. (2001), surprisingly little of the overall species
401 variation was explained by the measured environmental parameters. There are a number of
402 potential reasons for this. Firstly, vegetation assemblages are the result of many factors, not
403 just hydrological regime and soil development (Bakker et al., 2006). Inter-specific
404 competition, migration rates and chance colonisation (Bossuyt et al., 2003; Bossuyt et al.,
405 2005) introduce a high degree of heterogeneity to slack vegetation, even to adjoining slacks
406 with similar hydrological regimes on the same site. A similar effect of site overriding
407 hydrological parameters has been shown on floodplain meadows (Kalusova et al., 2009).
408 Other factors such as secondary disturbance, and seed bank longevity also influence
409 composition (Studer-Ehrensberger et al., 1993) and increase heterogeneity. Lastly, grazing
410 pressure varies within and between sites and this may be a factor. However, Plassmann et al.
411 (2010) showed that introduction of grazing had no major effect on slack vegetation at
412 Newborough, with grazing effects secondary to both moisture and soil parameters. In these
413 relatively young (20 – 150 years) well-buffered slacks prior to decalcification, although
414 organic matter accumulates rapidly with age (Jones et al., 2008), slack age is a poor
415 determinant of community type (Jones et al., 2010), nonetheless, it may be an important
416 factor influencing vegetation composition. In contrast to some Dutch systems, salt spray,
417 seawater influence and decalcification (Lammerts and Grootjans, 1998) are not dominant
418 influences here; soil pH drops to a mean of 6.3 in the older SD17 community. Lastly, this
419 study covered a narrower environmental gradient compared with e.g. Lammerts et al. (2001)
420 who also included primary slacks and those with periodic saline influence, and included
421 samples from the transitional areas between communities rather than focusing on
422 homogenous stands of vegetation (e.g. Rodwell et al. 2000).

423

424 A potential issue for interpretation with respect to hydrological regimes is the temporal
425 resolution of the hydrological data. Water level data available at monthly resolution in this
426 study did not allow calculation of sum exceedance values, (i.e. the duration of time the water
427 table is above or below thresholds for drought stress or water logging stress and the
428 cumulative summation of that exceedance) which have been used to interpret plant species
429 hydrological niches in other wetland ecosystems (Gowing et al., 1997). The water-holding
430 capacity and soil permeability alter with organic matter content. As a result, the same water
431 table depth below ground surface may differentially affect communities with different
432 organic profiles and different mineralogy. It is also possible that there is a disconnect
433 between water tables and plant responses in winter and summer, with groundwater levels in
434 summer being less important than rainfall events which may re-wet surface layers and the
435 rooting zone but do not contribute to recharge and therefore changes in water table.
436 Nonetheless, despite these constraints, this study has for the first time identified the average
437 hydrological regimes of a suite of Atlantic wet slack vegetation communities.

438

439 Relationships are based on hydrological data over a four-year period. This is an improvement
440 on previous ecohydrological studies in the UK based on one or, at the most, two years
441 hydrological data (Jones 1993; Ranwell, 1959), and is consistent with other studies. In Dutch
442 dune slacks Noest (1994) found that 5-yearly means of a range of hydrological parameters
443 provided better explanatory power on species' distribution than the same parameters
444 measured for the year of vegetation recording, or the previous year. In UK wet meadows, the
445 best explanatory power for changes in vegetation is found for hydrological variables over the
446 preceding three to seven year period: Sum Exceedance for waterlogging stress over a 3-year
447 period proved effective for predicting shifts between vegetation communities (Gowing et al.,
448 2005) whilst 5-7 year values for aeration and drought stress provide good explanatory power

449 for stable community distributions (Gowing et al., 2002). In this study we make the
450 assumption that the vegetation communities were in equilibrium with those hydrological
451 conditions, but further work is required to establish variability in dune slack species
452 assemblages between years and in response to changing hydrological regimes.

453

454 Dune groundwater levels are closely tied to climatic patterns and water levels vary
455 considerably from one year to the next. The few long-running records available for the UK
456 spanning several decades show large changes in average water tables, often over decadal
457 scales (Davy et al., 2010; Robins and Jones, 2012). The net recharge to groundwater is
458 strongly dependent on a fine balance between rainfall and evapotranspiration, and is as
459 dependent on the timing of rainfall during the year as it is on total annual rainfall (Clarke and
460 Sanitwong Na Ayutthaya, 2010). This makes dune slacks very sensitive to climatic changes.
461 Studies on the impact of climate change on recharge to aquifers in the eastern UK suggest
462 increasing recharge and water levels over the next few decades but becoming drier thereafter
463 (Younger et al., 2002; Yusoff et al., 2002), while in Ireland, an overall decrease in effective
464 run-off is predicted (Charlton, 2001). Clarke and Sanitwong Na Ayutthaya (2010) used the
465 Darcian groundwater flow equation to model the movement of water through the dunes,
466 which was then used to predict the effects of climate change. Their results suggested a
467 drastic, long-term decline in water levels of more than 100 cm was highly likely at a west
468 coast UK site. They also showed that climate change impacts on water tables were greater
469 than impacts of afforestation or coastal change. Climate change impacts were a factor of two
470 greater than the effects of sea level rise or afforestation, and an order of magnitude greater
471 than the effects of rapid coastal erosion or accretion.

472

473 Although studies in the Netherlands have developed models to predict responses of Dutch
474 dune species to groundwater change (von Asmuth et al. 2012; Witte et al. 2007; Geelen et al.
475 2004), no studies to date have quantified how changing water levels may affect dune
476 vegetation communities in the UK. Findings from this study allow, for the first time, a
477 projection of likely changes in vegetation communities under projected water level decline.
478 Figure 6 shows frequency distributions of median water level for the main slack and dry dune
479 grassland communities in this study plotted against water level projections to the 2080s for
480 what is currently a wet slack type at Ainsdale (after Clarke and Sanitwong Na Ayutthaya,
481 2010). This suggests that towards the end of the 2030s conditions are no longer favourable
482 for wet slacks and only dry slack community can persist. More worryingly, by the end of the
483 2050s it is likely that even dry slack communities will be replaced by dry dune grassland.
484 Although in reality these communities are unlikely to replace each other sequentially (e.g. see
485 Edmondson 1993 for a review of UK dune slack succession), it illustrates the severe shift in
486 hydrological regimes that may occur and shows the implications for vegetation assemblages.
487 These effects only take account of changes in water level. However changes in groundwater
488 chemistry and the interaction with surface soil layers will also occur. Surface soil
489 acidification, which is detrimental for the growth of basiphilous slack species, is a
490 consequence of lowering water levels, due to a reduction in the seasonal replenishment of
491 buffering capacity from base-rich groundwater (Sival and Grootjans, 1996). The diminution
492 of winter flooding is likely to have direct impact on soil chemistry too, changing nutrient
493 accumulation rates (Grootjans et al., 1991), redox processes (Sival and Grootjans, 1996),
494 microbial cycles (Grootjans et al., 1996) and will consequently affect species interactions.

495

496 [Figure 6 about here]

497

498 The effects of climate change may be exacerbated by drainage or groundwater abstraction,
499 and any form of water abstraction should be discouraged (Bakker et al., 2006; Davy et al.,
500 2010; Grootjans et al., 1996; Van Dijk and Grootjans, 1993). In contrast, management
501 techniques which encourage natural sand mobility may guarantee natural renovation of young
502 successional stages, allowing the formation of new blowouts or creation of new secondary
503 dune slack habitat through natural dune dynamics (Davy et al. 2010; Stratford et al., 2007).
504 Other management methods used to improve mobilisation in sand dunes systems (such as sod
505 cutting, removal of invasive scrubs, etc.) may be useful to alleviate detrimental effects of
506 climate change in the absence of natural mobility (Kooijman, 2004).

507

508 **5. Conclusions:**

509 Atlantic dune slack assemblages can be characterised by their hydrological regime and soil
510 parameters. However, there remains unexplained heterogeneity in the vegetation and the
511 measured hydrological and soil parameters explain only 22.5% of observed variation.
512 Assemblages separate broadly into wet slack, intermediate and dry slack types, with average
513 winter water levels differing by around 20 cm between categories. Given the magnitude of
514 these differences between community types, projected changes in water levels of over 100 cm
515 due to climate change over the next few decades are a cause for concern. Hydrological
516 regimes may shift completely from those currently associated with wet slacks to regimes
517 currently found under dry dune grassland over a period of 50 years, and management options
518 for responding to these changes should be explored urgently.

519

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527

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713

714 **Figure Legends:**

715

716 **Figure 1.** Typical hydrographs showing monthly water levels at Newborough Warren over a
717 four year period (2006-2009) for four dune slack communities and one dry grassland
718 community (SD8) for comparison. y axis shows metres above ground level, zero represents
719 ground surface.

720

721 **Figure 2.** PCA scores of the 245 quadrats, coded by vegetation community. The first three
722 axes (Eigenvalues 0.13, 0.11, 0.06) explain 29.6% of the variance. See table 4 for full list of
723 environmental variables.

724

725 **Figure 3.** CCA ordination analysis (a) distribution of environmental variables along first two
726 axes, (b) distribution of quadrats grouped by NVC -core and transitional vegetation-, (c)
727 distribution of quadrats representing core vegetation only, and distinguishing between SD14
728 subcommunities. See table 4 for a complete list of variables.

729

730 **Figure 4.** Average ion concentrations in soil, by vegetation community (core plus transitional
731 vegetation) -bars represent SE-. a) anions; b) cations. Calcium is downscaled 10 times to be
732 visualised with the other cations (e.g.: 1=10mg/100g dry soil).

733

734 **Figure 5.** Plot of hydrological niche occupied by three dune slack vegetation communities
735 and dry dune grassland showing mean annual maximum (i.e. winter) and minimum (i.e.
736 summer) water levels (metres above ground level, zero = ground surface), based on four-year
737 hydrological regime (2006-2009). Thick inner lines indicate bounds of 67%ile of their
738 distribution, thinner outer lines represent bounds for 90%ile. Dashed lines denote mean

739 annual maximum water levels for dry dune grassland based on quadrats surveyed in this
740 study.

741

742 **Figure 6.** Frequency distribution curves of median water level for different slack types (data
743 from this study, rescaled on 2nd y axis to max = 1), plotted against predicted changes in April
744 water level (metres above ground level, 1st y axis) of a currently wet dune slack (SD15) at
745 Ainsdale in north west England (after Clarke & Sanitwong Na Ayutthaya, 2010). Predicted
746 April water levels show central estimate (thin black line), 75% confidence interval (dark grey
747 zone) and 90% confidence interval (light grey zone).

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750 **Table 1: Locations and details of survey sites.**

Site	Latitude, Longitude	Rainfall 1971-2000 l.t.a. (mm)	Dune area (ha)	Piezometers	Slacks surveyed	Quadrats
Newborough Warren	53°08' N, 4°21' W	830	529	49	46	212
Aberffraw	53°11' N, 4°27' W	830	248	2	2	12
Ainsdale	53°35' N, 3°4' W	870	508	3	3	9
Whiteford Burrows	51°38' N, 4°14' W	1110	142	4	4	12

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Table 2: Summary of National Vegetation Classification communities and subcommunities identified in this study (adapted from Rodwell, 2000). Affinities with European vegetation associations in brackets.

NVC code	Name	Description
SD13	<u>Sagina nodosa-Bryum pseudotriquetrum</u> community (<u>Saginion maritimae</u> Westhoff, van Leeuwen & Adriani 1962, <u>Samolo-Littorelletum</u> Westhoff 1943)	Early successional stage , rich in bryophytes and liverworts. Usually with bare sand. Fairly drought tolerant.
SD14	<u>Salix repens-Campylium stellatum</u> Community (<u>Junco baltici-Shoenetum nigricantis</u> Westhoff 1946, <u>Pyrolo-Salicetum</u> Meltzer 1941)	Frequently species rich and associated with persistently humid soils and base-rich groundwater. Several rare species occur in this vegetation.
SD14b	<u>Rubus caesius-Galium palustre</u> subcommunity	Some of its constant species (<u>Ranunculus flammula</u> , <u>Carex nigra</u>) can indicate tolerance to very wet periods.
SD14c	<u>Bryum pseudotriquetrum-Aneura pinguis</u> subcommunity	Young successional stage , mosses have sparse cover, heliophilous and pioneer species can be present.
SD14d	<u>Festuca rubra</u> subcommunity	Characteristic of drier substrates , it can be an intermediate stage towards grass encroachment.
SD15b	<u>Salix repens-Calliergon cuspidatum</u> community <u>Equisetum variegatum</u> subcommunity (<u>Equiseto variegati-Salicetum repentis</u> Westhoff & Schaminée 1995)	Late successional stage , generally species poor. Less dependent on base-richness of water, but strongly related with flooding.
SD16	<u>Salix repens-Holcus lanatus</u> Community (<u>Salicion repentis arenariae</u> Tüxen 1952)	Late successional stage in dry slacks. Dominated by fescue and grasses, forbs are still indicative of calcicolous substrate
SD17	<u>Potentilla anserina-Carex nigra</u> Community (<u>Elymo-Rumicion crisp</u> Nordhagen 1940)	Species composition reflects damp habitat , recalling fen meadows. Forbs-rich, with a sparse shrub cover.
SD8b	<u>Festuca rubra-Galium verum</u> Community <u>Luzula campestris</u> subcommunity (<u>Galio-Koelerion</u> , Westhoff and den Held 1969).	Dune grassland rich in dicotyledons characteristic of fixed sands. This vegetation has some affinities with a calcicolous sward.

(Caricion davallianae Kilka 1934)

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Table 3. Environmental variables used in vegetation analysis, showing percentage of total species variation explained within CCA and significance, when tested singly or in combination. *: significant at 0.05 level; **: significant at 0.01 level; *: significant at 0.001 level.**

	Variables	Variance explained (%)	Significance
Hydrological variables	Annual maximum water level (Max) (m)	6.3%	***
	Annual minimum water level (Min) (m)	3.8%	***
	Annual range (Range) (m)	2.0%	***
	Time flooding (TiFlood) weeks	0.8%	*
Soil variables	Soil moisture % (Moist)	6.3%	***
	LOI% (LOI)	4.1%	***
	pH	3.8%	***
	sulphate (SO₄²⁻) ppm	3.3%	***
	ammonium (NH₄⁺) ppm	3.1%	***
	sodium (Na⁺) ppm	2.4%	***
	chloride (Cl⁻) ppm	2.3%	***
	fluoride (F) ppm	2.3%	***
	phosphate (PO₄³⁻) ppm	1.4%	***
	magnesium (Mg²⁺) ppm	1.3%	***
	potassium (K⁺) ppm	1.1%	**
	nitrate (NO₃) ppm	0.6%	n.s.
calcium (Ca²⁺) ppm	0.5%	n.s.	
Ellenberg indicators	Ellenberg F, N, R, L combined	21.2%	***
Combined models	Hydrological parameters (Max + Min + Range + TiFlood)	8.9%	***
	Key hydrological + soil parameters (Max + pH + Moist)	12.9%	***
	Extended hydrological + soil parameters (Max + Moist + pH + NH₄ + PO₄)	14.6%	***
	All variables	22.5%	***

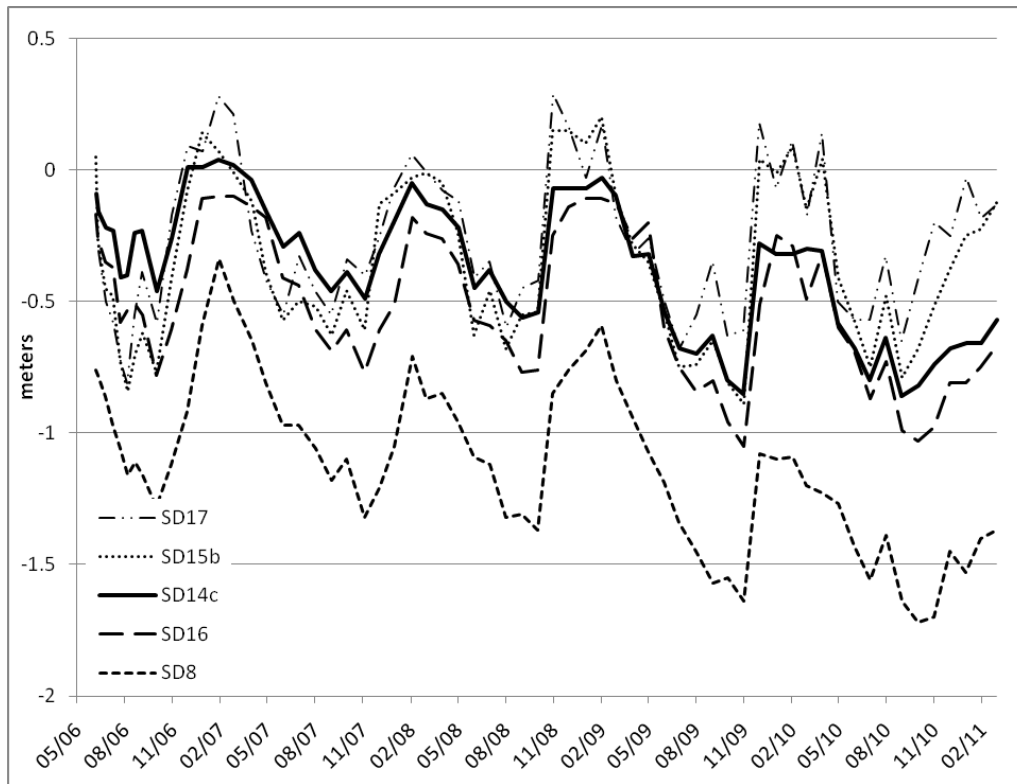
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Table 4: Summary environmental parameters for core quadrats (n=139) and core plus transitional quadrats (n=233), SD8 excluded. Values for each variable are expressed as mean \pm standard error, brackets show minimum and maximum values for each community. Hydrological parameters are calculated averages over 4 years. For each variable, lower case letters denote significant differences between community types using core quadrats only, upper case letters for tests using core plus transitional quadrats.

		CORE GOOD QUALITY VEGETATION					CORE AND TRANSITIONAL VEGETATION			
NVC		SD15b	SD14b	SD14c	SD14d	SD16	SD17	SD15	SD14	SD16
VARIABLE		n=15	n=28	n=47	n=35	n=14	n=7	n=32	n=165	n=29
WATER LEVEL (m)	average Minimum	-0.61 \pm 0.020 ^a (-0.69 , -0.51)	-0.62 \pm 0.016 ^a (-0.78 , -0.53)	-0.70 \pm 0.020 ^{ab} (-0.90 , -0.41)	-0.75 \pm 0.014 ^b (-0.88 , -0.62)	-0.98 \pm 0.059 ^c (-1.57 , -0.81)	-0.75 \pm 0.034 ^A (-0.85 , -0.58)	-0.64 \pm 0.015 ^A (-0.79 , -0.51)	-0.73 \pm 0.014 ^A (-1.41 , -0.41)	-0.98 \pm 0.050 ^B (-1.57 , -0.77)
	average Maximum	0.21 \pm 0.018 ^a (0.12 , 0.29)	0.17 \pm 0.027 ^a (-0.07 , 0.32)	0.00 \pm 0.022 ^b (-0.21 , 0.22)	0.02 \pm 0.016 ^b (-0.12 , 0.18)	-0.25 \pm 0.061 ^c (-0.82 , -0.02)	0.09 \pm 0.032 ^{AB} (-0.01 , 0.25)	0.16 \pm 0.020 ^A (-0.09 , 0.29)	0.02 \pm 0.014 ^B (-0.48 , 0.32)	-0.22 \pm 0.043 ^C (-0.82 , -0.02)
	average Range	0.81 \pm 0.008 ^{ab} (0.77 , 0.85)	0.80 \pm 0.013 ^a (0.70 , 0.85)	0.69 \pm 0.015 ^b (0.49 , 0.85)	0.77 \pm 0.007 ^a (0.70 , 0.83)	0.72 \pm 0.016 ^{ab} (0.63 , 0.80)	0.84 \pm 0.005 ^A (0.83 , 0.85)	0.79 \pm 0.009 ^{AB} (0.70 , 0.85)	0.74 \pm 0.008 ^B (0.49 , 1.02)	0.76 \pm 0.023 ^{BC} (0.63 , 1.02)
SOIL	Moisture (%)	33.6 \pm 1.68 ^a (19.9 , 43.7)	29.8 \pm 1.06 ^{ab} (17.9 , 41.6)	18.1 \pm 0.76 ^c (7.2 , 33.8)	25.4 \pm 1.42 ^b (12.4 , 60.7)	15.4 \pm 1.52 ^c (7.5 , 25.0)	34.8 \pm 2.86 ^A (19.8 , 44.3)	36.4 \pm 1.46 ^A (19.9 , 63.7)	23.6 \pm 0.61 ^B (7.2 , 60.7)	17.4 \pm 1.05 ^C (5.7 , 28.5)
	LOI (%)	7.2 \pm 0.37 ^a (3.8 , 9.5)	5.1 \pm 0.38 ^{bc} (0.9 , 10.8)	2.6 \pm 0.17 ^d (0.6 , 5.5)	5.8 \pm 0.42 ^{ab} (3.0 , 18.4)	3.6 \pm 0.30 ^{cd} (1.5 , 6.0)	8.9 \pm 0.77 ^A (5.8 , 11.9)	8.2 \pm 0.55 ^A (3.8 , 17.9)	4.38 \pm 0.18 ^B (0.6 , 18.4)	4.3 \pm 0.24 ^B (1.4 , 6.6)
	pH	7.1 \pm 0.17 ^c (6.2 , 8.1)	7.8 \pm 0.10 ^{ab} (6.3 , 8.6)	8.2 \pm 0.04 ^a (6.9 , 8.7)	7.1 \pm 0.11 ^c (6.0 , 8.1)	7.5 \pm 0.15 ^{bc} (6.5 , 8.3)	6.3 \pm 0.17 ^C (5.8 , 7.0)	7.0 \pm 0.12 ^{BC} (5.9 , 8.2)	7.7 \pm 0.05 ^A (6.0 , 8.7)	7.4 \pm 0.12 ^B (6.0 , 8.3)
% cover	Bare ground	0.1 \pm 0.13 ^{ab} (0 , 2)	0.4 \pm 0.25 ^b (0 , 6)	4.2 \pm 1.18 ^a (0 , 40)	0.0 \pm 0.00 ^b (0 , 0.1)	2.2 \pm 1.8 ^{ab} (0 , 25)	2.6 \pm 1.19 ^A (0 , 8)	0.1 \pm 0.06 ^B (0 , 2)	2.0 \pm 0.46 ^A (0 , 40)	1.4 \pm 0.93 ^{AB} (0 , 25)
	Litter	1.9 \pm 0.73 ^a (0 , 10)	0.4 \pm 0.22 ^b (0 , 5)	0.4 \pm 0.12 ^b (0 , 4)	0.4 \pm 0.20 ^b (0 , 6)	0.7 \pm 0.48 ^{ab} (0 , 5)	28.6 \pm 14.2 ^A (0 , 90)	3.5 \pm 1.88 ^B (0 , 60)	1.3 \pm 0.49 ^B (0 , 60)	2.6 \pm 1.43 ^B (0 , 40)

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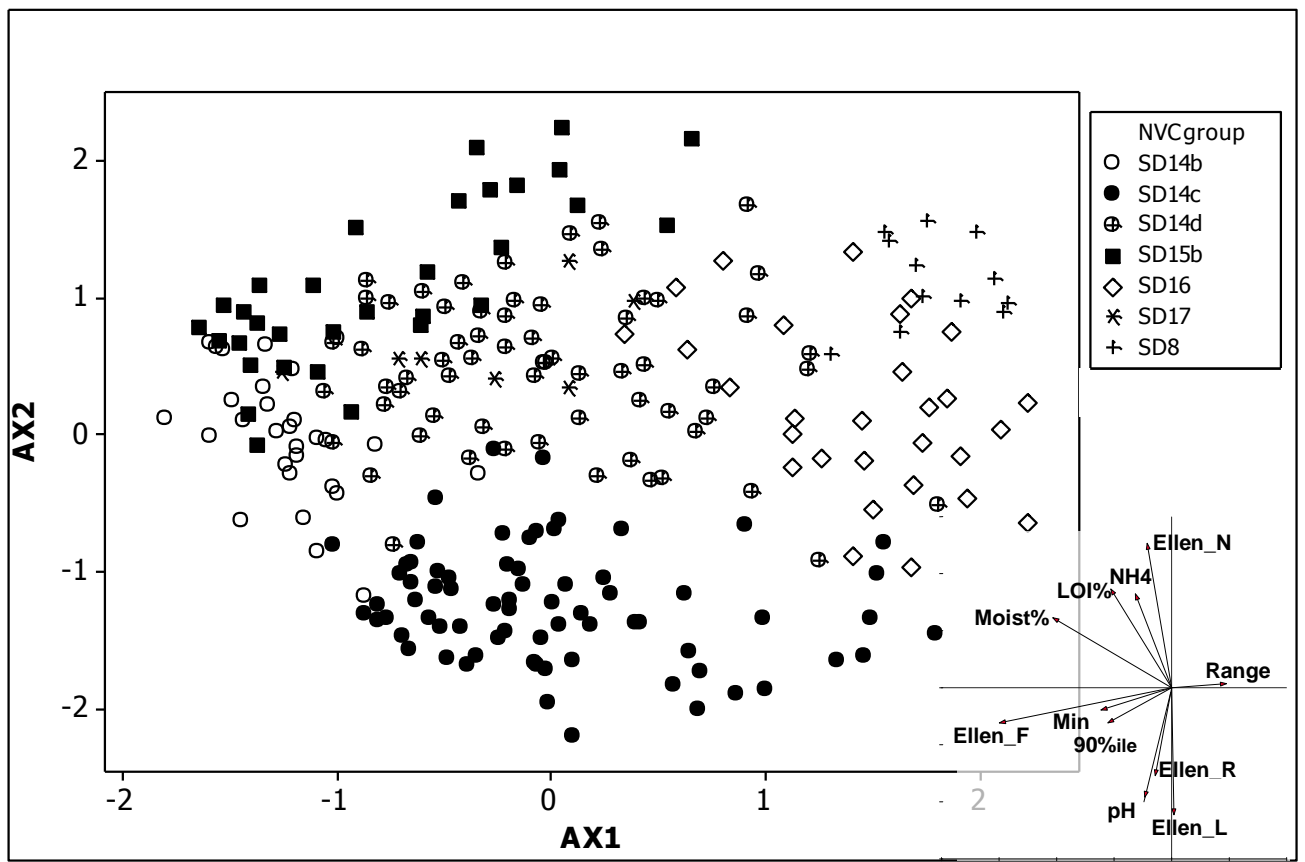
Figure 1.



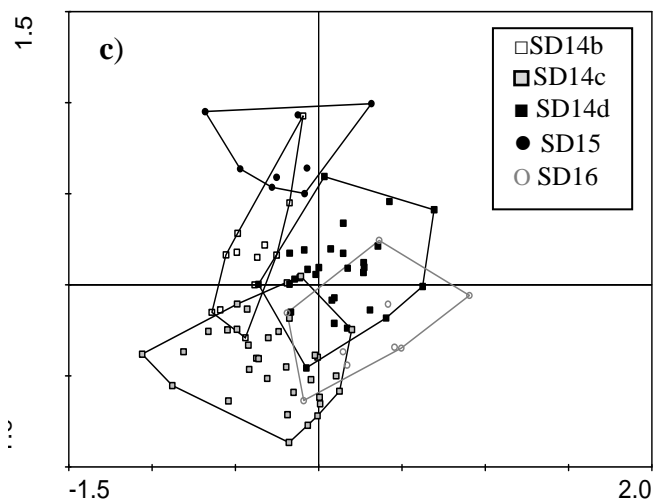
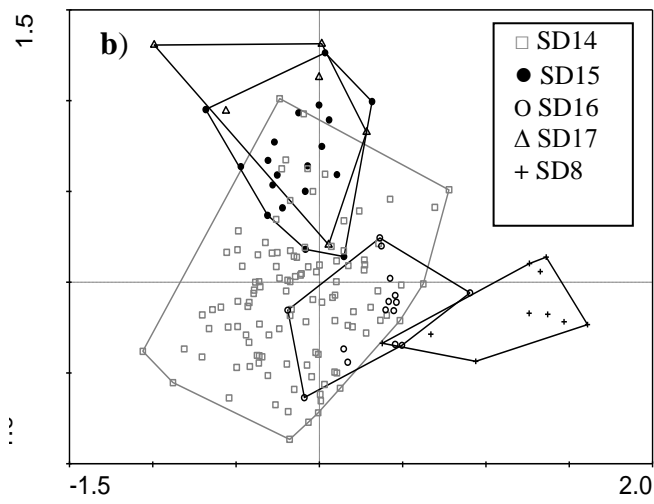
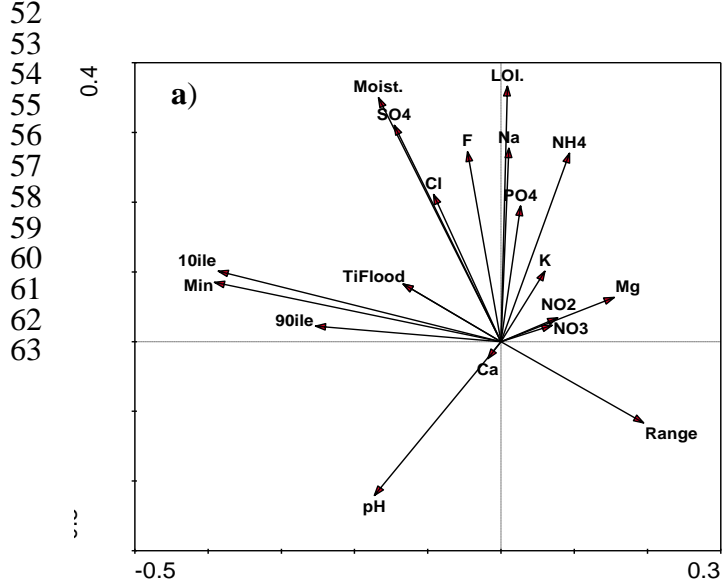
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Figure 2.

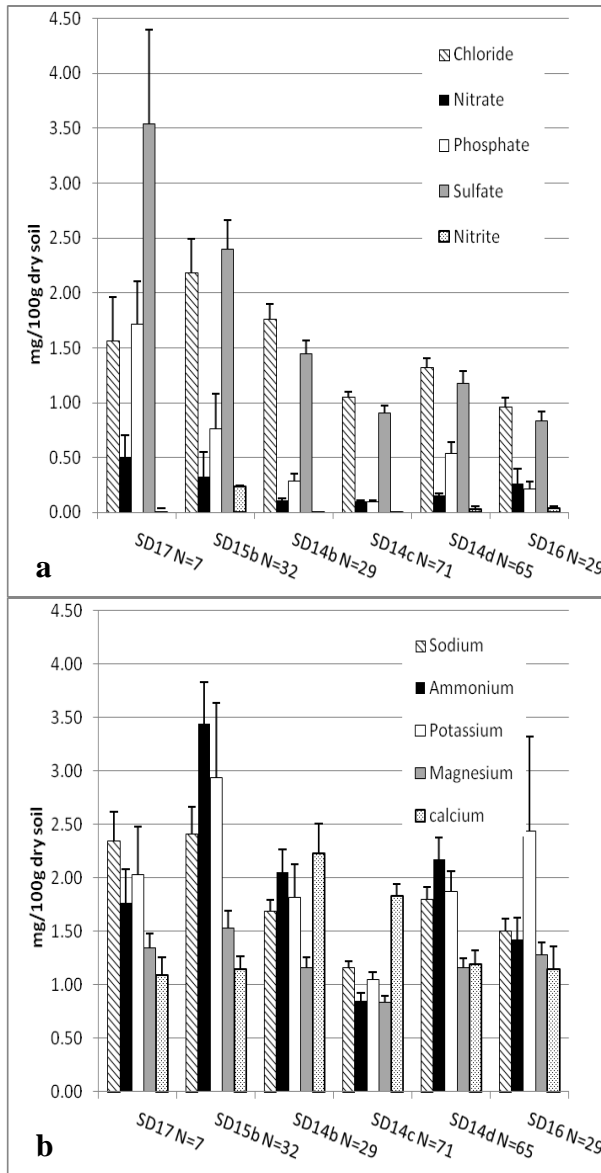


51 **Figure 3.**

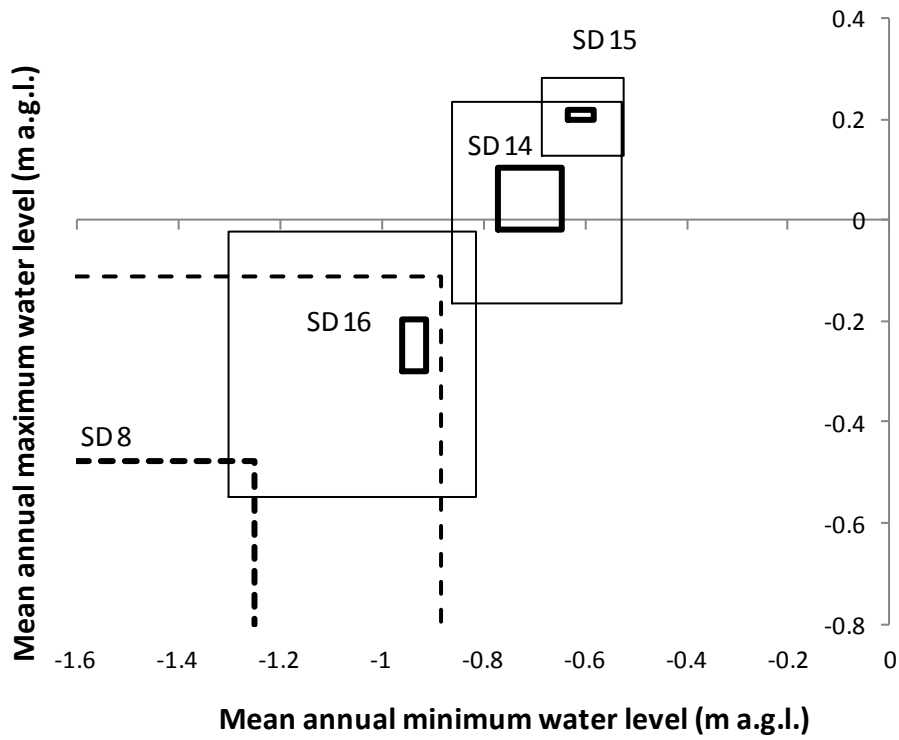


64 **Figure 4.**

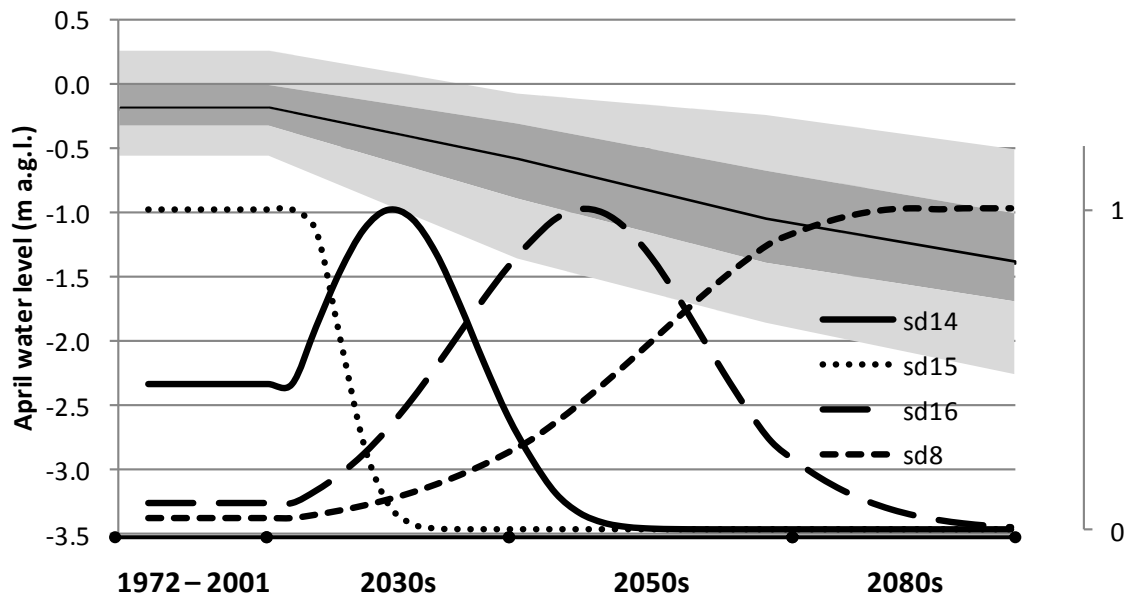
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109 Figure 5
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