



26 conditioned by drying soil and the lower temperatures. The lower temperatures are limiting only  
27 when the soil moisture is below 32% of the field capacity. Thus, we propose to keep the soil  
28 moisture around 50% during the fall as a key practices for mitigating the effect of seasonality  
29 and its intensification with the climate change, even more if added to management routine  
30 practices the soil and water conservation.

31 **Keywords:** tropical pasture seasonality, soil carbon dioxide emission, irrigation management  
32 strategy.

### 33 1. Introduction

34 Global climate change is challenging the sustainability of agricultural systems at all levels.  
35 Grasslands plays a major role in the global ecosystem and are commonly used to supply biomass  
36 to milk and meat production, and wool. The effects of seasonality on pasture ecosystems are  
37 mainly reflected during the driest and coldest periods of the year with a strong drop in pasture  
38 production, which can negatively affect the profitability and sustainability of the production  
39 systems that depend of grass (O'Mara, 2012, Keller et al., 2014). The reduced capacity of  
40 regeneration of the biomass before the arrival of the new rainy season in tropical areas is affected  
41 by the pre-eminence of temperature and rainfall factors, which makes it difficult to define  
42 resilience strategies. However, throughout the annual growth, winters are a period in which  
43 grasslands exhibit low growth rates and low resolution of biometric and growth changes, which  
44 in turn, hinders a deeper understanding of the environmental conditions for the delay of the  
45 spring regrowth of the pastures (Barbosa et al., 2007). In this context, studies on carbon soil  
46 efflux/exchange can help to detect the plant responses to environmental conditions in order to  
47 improve resilience strategies which may involve water management practices.

48 Among pastures, the *B. brizantha* is well spread in vast areas of tropical grasslands, savannas  
49 and forests with infertile acid soils of many countries in South America, India, Africa and  
50 Australia. Some of the main areas involved are the spear-grass zone of northeastern Australia,

51 the cerrado and Amazonian regions of Brazil and the llanos of Venezuela and Colombia (FAO,  
52 2005; Jank 2014).

53 Grassland ecosystems store most of their carbon in soils, where turnover is relatively slow,  
54 and in most grassland types, below-ground NPP is at least equal to or higher than aboveground  
55 production (IPCC, 2007). Both seasonality and irrigation unfeasibility of extensive pasture areas  
56 are a threat to pasture and farm profitability in these ecosystems. Several levels of overgrazing  
57 occur during the low grass period. This can lead to increased erosion of the soil and its  
58 degradation, as well as intensification of the energy partition for sensible heat, resulting in  
59 alterations of the hydrological cycle and the capacity of the soil to store carbon in these regions.  
60 These events further contribute either to a prolongation of dry season or to reduction of rainfall  
61 (IPCC, 2007). Future anthropogenic climate change and variability in weather events increase  
62 the need for mitigation in response to agricultural greenhouse gas emissions.

63 According the Distribution Mapping of World Grassland Types (Dixon et al., 2014) and the  
64 International Vegetation Classification (IVC) with dominant grassland types, South America  
65 contains 16 different types of grasslands with a large share of Brazil with three dominant types,  
66 being a) Brazilian-Parana Lowland Shrubland, Grassland and Savanna (2,035,627 km<sup>2</sup>), b)  
67 Brazilian-Parana Freshwater Marsh, Wet Meadow and Shrubland (170,501 km<sup>2</sup>), and c)  
68 Brazilian-Parana Montane Shrubland and Grassland (26,247 km<sup>2</sup>). Increasing our knowledge of  
69 biomass dynamics and carbon fluxes would be decisive in directing conservation policies and  
70 management practices.

71 In Brazil, the grassland used to production of beef cattle and milk is predominantly situated  
72 within the intertropical zone. In this zone the forage plants present vigorous growth during the  
73 warm and rainy season (October-March), contrasting with a drastic reduction or halt on the  
74 growth in the dry season (April to September) (Barbosa et al., 2007). This seasonal alternation in  
75 the growth rate of forage plants, “seasonality”, has been subject of several studies (Li and Snow,

76 2011; Beecher et al., 2013; Demanet et al., 2015; Wingler and Hennessy, 2016) and is highly  
77 correlated to productive performance of animals kept in pastures, affecting significantly dairy  
78 and meat production.

79 Reduced production of tropical forage plants during the driest periods of the year is one of  
80 the main issues in worldwide cattle breeding strategies. Pastures commonly reduce their  
81 production by near 75% in winter when compared to spring and summer seasons (Pedreira,  
82 1998; Dupas et al., 2010, Silva et al., 2012). The effect of seasonality on pasture is enhanced by  
83 solar radiation, temperature and precipitation rates, which also depends on the interaction among  
84 these climatic factors and the genetics of the cultivated species. Among several pastures in  
85 Brazil, the brachiaria genotypes occupies more than 70% of 172 million hectares used for  
86 grazing (Valle et al., 2001). Marandu cultivar is the most commonly used variety due to its  
87 tolerance to depleted soil fertility, insect resistance and elevated productivity (Valentini et al.,  
88 2008; Cruz et al., 2011). Consequently, agro-ecosystems where Marandu grows represent the  
89 most practical and economical alternative for ruminant feeding, by the lower production costs  
90 (Dias-Filho, 2014).

91 There is a consensus about the importance of grassland management to the carbon cycle and  
92 compelling reasons to consider this procedure as part of the holistic approach to carbon  
93 sequestration (Frank 2002; IPCC, 2007; Soussana et al., 2010; Ward et al., 2016; IPCC, 2014). In  
94 this respect, seasonality plays an important role (Skinner, 2007; Peichl et al., 2011) because  
95 Brachiaria species have very developed root system, which influence on both microorganisms'  
96 activity (Gopalakrishnan, et al., 2009) and carbon sequestration in soils (Andrade et al., 2008;  
97 Ramírez et al., 2009).

98 According to Davidson et al. (2002), the soil stores two or three times more carbon than the  
99 atmosphere, and the soil CO<sub>2</sub> efflux is one of the main components of global carbon cycle,  
100 accounting for about 50% of the carbon that composes the ecosystems (Wagai et al., 1998;

101 Roberts, 2000). The understanding of soil CO<sub>2</sub> efflux is the key to elucidate changes in the soil-  
102 plant-atmosphere system. Therefore, this comprehension is crucial for improving the  
103 mechanisms of the atmosphere-biosphere interaction in global and regional climate models (  
104 Zanchi et al., 2003).

105 The measurement of soil surface CO<sub>2</sub> efflux is the most widely used method to estimate soil  
106 respiration rate in ecosystems. This is associated with its temperature and moisture conditions,  
107 and dependent on the temporal and spatial variability of these parameters (Davidson et al., 2000;  
108 Rustad et al., 2000; Fang and Moncrieff, 2001). Understanding the CO<sub>2</sub> flow of soil and carbon  
109 stocks in grazing systems has been extremely important in assessing long-term trends in soil  
110 organic carbon content (Hopkins et al., 2009). On the other hand, the greatness of their  
111 measurements and sensitivity to small changes in the soil resulting from the plant-soil-  
112 atmosphere interaction, specifically CO<sub>2</sub> efflux, may allow a better characterization of the  
113 evaluation of the impact of seasonal variability on pastures, and assist in the definition of pasture  
114 management strategies throughout the coldest and driest period of the year to improve annual  
115 pasture production

116 In Brazil, studies analyzing CO<sub>2</sub> efflux in pastures were more restricted to the Cerrado and  
117 Amazon biomes (Von Randow et al., 2004; Ruhoff et al., 2009; Silva-Junior et al., 2013). Thus,  
118 the understanding of the relationship between soil moisture and CO<sub>2</sub> efflux especially in tropical  
119 ecosystems at either, short and long term is still highly limited. The majority of the studies have  
120 been focused on variations in CO<sub>2</sub> efflux in short periods of time (Grahammer et al., 1991; Liu et  
121 al., 2002), while others assessed the relations of efflux and humidity based on observations of  
122 seasonal variation in forests and temperate grasses with organic soils (Davidson et al., 1998, Luo  
123 et al., 1996; Jones, 2005), different from the tropical soils of the present study, characterized by  
124 low organic matter and high clay content.

125 The plant growth analysis enables the characterization of plant responses to climate change  
126 and weather variability (Hunt, 1990; Poorter and Garnier 1996) which can be applied to alleviate  
127 the effect of seasonal water deficit for plant growth. Thus, these analyses can characterize the  
128 physiological response related to growth and yield of brachiaria to distinct levels of irrigation  
129 (Waldron et al., 2002) and temperature responses (Medek et al. 2007). Comprehensive functions  
130 have been considered, such as Relative Growth Rate (RGR), Net Assimilation Rate (NAR) and  
131 the Leaf Area Ratio (LAR), from the dynamics of growth of the foliar area and the biomass of  
132 the plant, which are used to understand how the plants respond to climate/weather conditions  
133 (Hunt, 1990, Munns et al., 2017). Relative Growth Rate ( $RGR = LAR \times NAR$ ) is a prominent  
134 indicator of plant strategy with respect to productivity and its relationship with environmental  
135 stress and disturbance regime by the factors of the climate. This variable is closely related to the  
136 daily rate of photosynthesis per unit LA and time.

137 The joint evaluation of plant growth characteristics and the CO<sub>2</sub> efflux of the soil, could  
138 allow to discriminate the temporal hierarchy and the specific response of the brachiaria to the  
139 climatic factors and to propose management practices that increase the resilience before the  
140 current climatic variability and your tendency for the future. Therefore, the main objectives are  
141 (1) to measurement the soil respiration to understand the reasons they change over time; (2) to  
142 identify the contribution of each component of climate factors; (3) to clarify the relationships  
143 between soil respiration components and abiotic factors; (4) to establish a temporal hierarchy of  
144 low temperatures and soil moisture in the dynamics of biomass production and carbon flux  
145 throughout the growing seasons of the year.

## 146 **2. Materials and Methods**

### 147 **2.1. Experimental conditions and plant management**

148 Two experiments were carried out using *Brachiaria brizantha* cv Marandu, at the Agriculture  
149 Experimental Area (CCAEA), of the Agricultural Engineering Department of the Federal

150 University of Viçosa (UFV-Brazil). The first experiment was conducted under field conditions  
151 during the year 2014, to determine the physiological response of Brachiaria to climatic factors  
152 and the Net Carbon Exchange Rate (NCER) changes, due to seasonal climate variability. Taking  
153 into consideration the 2014 results insofar as water availability is concerned, the second  
154 experiment was carried out in drainage lysimeters, during 2016, aiming to elucidate changes in  
155 biomass production and soil CO<sub>2</sub> efflux linked to distinct levels of water availability in the  
156 winter/spring transition period.

### 157 **2.1.1 Field Experiments**

158 Field experiment was performed from January to December 2014, in an area of 10 m wide by  
159 20 m long, divided in four similar areas, characterized by a red-yellow Latosol, Dystrophic soil,  
160 with 70% of clay; 9% of silt; 12% of coarse sand; and 9% of fine sand. The pasture was sown on  
161 December 10, 2013 with 4-6 seeds per sowing point, with spacing of 0.25 x 0.25 m. Moreover,  
162 the spacing was adjusted to plant density among 20-25 plants m<sup>-2</sup>, which is a typical plant  
163 density used in tropical pastures. After emergency, the plants were allowed to grow until  
164 February 5, 2014 (57 days after sowing), when a “grazing cut equalization” (GCE) was carried  
165 out. This practice consists in cutting-off the aerial biomass at 5 cm above ground to stimulate  
166 tillering and the root expansion. After that, both pasture and experiment were considered  
167 established (Fig. 1 A).

168 Two harvest cuts occurred throughout the experiment. The first was performed in 21 May,  
169 and the second harvest in 12 November. Thus, the study was organized in two seasonal periods  
170 covering two-production cycles, summer-fall (05-Feb until 16-May/05) and winter-spring (06-  
171 Jun to 11-Dec), as function of the expected plant growth response to climate variability. In  
172 addition, the harvest time in both periods were defined as pre-flowering stages.

173 Soil acidity and fertility in both experimental areas were corrected from the results of  
174 physicochemical analysis as recommended by Ribeiro et al. (1999) and Barcelos et al. (2011),

175 the first adjusted to 70% base saturation index and the latter to attain 100, 80, 50 kg/ha of N-  
176  $P_2O_5$ - $K_2O$ , respectively. The phosphate was applied at the beginning of the experiment, while  
177 half of  $K_2O$  and N requirements were split between the experiment start and the first harvest.

### 178 **2.1.2 Lysimeter Experiment**

179 In the lysimeter experiment (year 2016), the influence of water availability on the growth of  
180 both biomass and soil  $CO_2$  efflux were evaluated. Therefore, four treatments of water availability  
181 based on accumulated brachiaria evapotranspiration (ETc) were applied with three replicates:  
182 T1=100% replacement (L1), T2=75% (L2), T3=50% (L3) and T4=25% (L4). Culture  
183 evapotranspiration was obtained as  $ETc = ETo \times Kc$ , with reference evapotranspiration (ETo) and  
184 different values of Kc based on cv. Marandu values which was 0.80, recommended in irrigation  
185 of growth pasture (Alencar et al., 2009), and 1.0 for the end phase of growth (Quintanilha et al.,  
186 2006).

187 The crop coefficient (Kc) is obtained by the relation between ETc under potential conditions  
188 and reference evapotranspiration (ETo) and varies with the stage of crop growth (Allen et al.,  
189 1998).

190 This lysimeter experiment was conducted from June 1 to November 18, 2016, covering two  
191 cycles of production, allowing for evaluation of temporal changes of soil  $CO_2$  efflux,  $CO_2$  leaf  
192 influx and biomass accumulation. During this period, brachiaria goes through two weather  
193 characteristics with dry and cold months (winter), and in spring when the temperature arose  
194 rapidly and the rain increased favoring conditions for brachiaria growth. Prior to the  
195 establishment of the experiment a standardization cut was achieved in all experimental units.  
196 Afterwards, fertilization and irrigation management have been carried out.

197 The soil in the lysimeters was red-yellow Latosol, Dystrophic with textural class of 70%  
198 clay; 9% silt; 12% coarse sand; and 9% fine sand. Two fertilizations were performed in the  
199 beginning of each production cycle. Aiming to meet the nutritional criteria, the L1 treatment was

200 adopted as reference for other treatments; therefore, 15, 3.5 and 18 kg ha<sup>-1</sup> of N, P<sub>2</sub>O<sub>5</sub> and K<sub>2</sub>O  
201 were applied, respectively, for each ton of dry mass produced in the previous harvest, as  
202 proposed by Vilela et al. (1998) and Barcellos et al. (2011).

203 The first evaluated cycle covered the end of autumn and the entire winter (01-Jun to 27-Set),  
204 with a total of 118 days. The second evaluated cycle was carried out on late spring, totalizing 51  
205 days between the first and the second harvest (28-Set to 18-Nov). The harvest occurred when the  
206 plants height was 35 cm.

## 207 **2.2. Climatic seasonality and meteorological data**

208 The World Meteorological Organization (WMO) defines climate as the statistical description  
209 of weather averaged over a period of time (30 years is usually) and describe a long –time  
210 atmosphere "behaves". Whilst 'weather' describes the physical state of the atmosphere at a  
211 particular place at a particular time, 'climate' can be defined as the probability of deviations from  
212 average values, including the probability of extreme values. The standard classification of the  
213 climatic zones is mainly based on the annual cycles of temperature and rainfall.

214 Daily meteorological data of 2014 was obtained from the National Institute of Meteorology  
215 (INMET) station No. 83642, located at 0.6 km from the experimental area. In the 2016-year  
216 experiment, data were collected by an automatic weather station, model Vantage Pro2 (DAVIS®)  
217 located in the experimental site.

218 Meteorological data of temperature, precipitation and evapotranspiration collected during  
219 both experiments are compared with climatological data series based on 34 observed years (1980  
220 to 2013, <https://utexas.app.box.com/v/xavier-et-al-ijoc-data>). This comparison is important to  
221 demonstrate the anomalous deviation of temperature and precipitation in the years 2014 and  
222 2016, both in the rainy season and in the dry season, compared to local climatological conditions.  
223 Our climatic characterization focus on the seasonal influence of meteorological  
224 factors/parameters on brachiaria growth and the CO<sub>2</sub> efflux. Special emphasis is given to the

225 months where changes of air and soil temperatures, precipitation and the evapotranspiration  
226 affected the brachiaria phenology and CO<sub>2</sub> efflux.

### 227 **2.3. Plant dry matter and growth**

228 Dry matter above ground (DMAG) was determined from sampling areas of 0.25m x 0.25m.  
229 The sample collection was carried out during nine times along summer/fall period and 11 times  
230 in winter/spring. Six samples were collected from each of four plots, and the results were  
231 expressed as accumulated dry matter in both space and time.

232 In 2016 experiment, the dry mass was determined from the biomass collected in a sample  
233 area of 0.25 x 0.25 m in four times at first production cycle and three times at second production  
234 cycle. The samples were collected from four repetitions.

235 To determine the total dry matter (stalks plus leaves), the ground biomass was cut-off at  
236 seven centimeters above the ground. In each sample, the leaf area, leaf mass and stems were  
237 determined. Leaf area was determined with a leaf area integrator LI-3100 (Li-COR, Lincoln, NE,  
238 USA). Samples were then dried until achieve constant weight in oven with forced air circulation  
239 at 70°C. Fresh and dry weight was determined in semi-analytical scale.

240 Dry matter, leaf area and ground area were used to calculate the leaf area index (LAI),  
241 relative growth rate (RGR), net assimilation rate (NAR), leaf area ratio (LAR), according to the  
242 Growth Analysis Methodology (Hunt, 1990, Munns et al., 2016).

### 243 **2.4. Soil CO<sub>2</sub> efflux and water capacity**

244 The efflux of CO<sub>2</sub> was measured along the growing season by 16 times in 2014, and seven  
245 times in 2016, between 8:00 and 10:00 AM. This process was conducted at two points for each  
246 plot, a total of eight points plus an additional central point at each lysimeter during 2016  
247 experiment.

248 The CO<sub>2</sub> soil efflux was measured through to net carbon exchange rate (NCER,  $\mu\text{mmol}$   
249  $\text{CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ ) measurements, with an infra-red gas analyzer (LC-Pro+, ADC BioScientific Ltd.,

250 Hoddesdon, United Kingdom) which was used coupled with a soil respiration chamber. The  
251 chamber was attached to stainless steel rings of 110 mm diameter by 70 mm height, inserted in  
252 the soil at a depth of 63 mm to avoid advective air fluxes. NCER readings were performed after  
253 an equilibration time of 10 to 15 minutes. The soil temperature was measured at the same time of  
254 NCER measurement, at 10 cm depth, and 5 cm away from soil cylinders.

255 The NCER short-term variability was analyzed by evaluating the physiological features of  
256 the biomass (LAI, RGR, Leaf photosynthesis), and daily and accumulated data of the  
257 meteorological factors (air temperatures and precipitation) to characterize the relationship with  
258 the plant responses through growing seasons. The relationship among variables was analyzed by  
259 Pearson correlation. Soil available water capacity (AWC) was determined from soil samples in  
260 12 times in the lysimeter experiment by the difference between volumetric water content in field  
261 capacity and the permanent wilting point (Allen et al., 1998; Borges et al., 2009, Silva et al.,  
262 2014).

## 263 **2.5. Leaf photosynthesis**

264 In 2016 experiment, the net CO<sub>2</sub> assimilation of grass leaves was measured between 10:00  
265 and 12:00 AM, concomitant with soil CO<sub>2</sub> efflux measurements in the median region of the +1  
266 leaf (younger and completely expanded leaf). An infrared gas analyzer (IRGA-LI-6400XT, LI-  
267 COR, Lincoln, NE, USA) configured to flux density of 1250  $\mu\text{mol m}^{-2} \text{s}^{-1}$  has been used, with  
268 atmospheric CO<sub>2</sub> inside the equipment chamber kept around 380  $\mu\text{mol CO}_2 \text{ mol}^{-1}$ , matching the  
269 average air concentration of CO<sub>2</sub> at the site, in Viçosa (Pereira-Flores, et al., 2016). The  
270 temperature inside IRGA-LI-6400XT block chamber was fixed in accordance with the  
271 environment temperature. The experimental data were analyzed by linear regression, Pearson  
272 correlation and descriptive statistics.

273

### 274 3. Results and Discussion

#### 275 3.1 Climate of 2014 and 2016 and brachiaria cardinal temperatures

276 During winter (June-July-August) can occur low temperatures very close to the minimum  
277 cardinal temperature for growth ( $T_b = 11.5\text{ }^\circ\text{C}$ ) in *Brachiaria brizanta* cv Maradu (Silva et al.,  
278 2012). The  $T_{\min}$  of June, July and August reached  $11.9\text{ }^\circ\text{C}$ ,  $12.0\text{ }^\circ\text{C}$  and  $11.2\text{ }^\circ\text{C}$ , respectively  
279 (Table1 in climatological serial data). In months before and after June-August period, the  $T_{\min}$   
280 varied from  $14$  to  $19\text{ }^\circ\text{C}$ , which represent non-limiting factors to the growth of cv. Marandu. It  
281 has been found that temperatures rising between  $12.5$  and  $16.5\text{ }^\circ\text{C}$ , increases the growth rate by  
282 about  $1.9$ - $5.9$  of biomass compared to the growth reported when temperature is  $11.5\text{ }^\circ\text{C}$  (Silva et  
283 al., 2012).

284 Annual rainfall average in the experimental location is  $1161.8\text{ mm}$ . The rainy season occurs  
285 between October and March (monthly rainfall between  $130$  and  $200\text{ mm}$ ), and a period of water  
286 restriction is experienced from April to September (monthly rainfall inferior to  $100\text{ mm}$ ). The  
287 lowest monthly rainfall occurred during the dry season of June, July and August, recording  $17.5$ ,  
288  $11.1$  and  $16.3\text{ mm}$ , respectively. Moreover, these months had also the highest relative ETo  
289 resulting in water deficit, with monthly evapotranspiration values exceeding the precipitation in  
290  $3.5$ ,  $6.3$   $5.4$  times (Table 1 in climatological serial data).

291 Thus, the experimental location has a prominent dry period and thermal limitation stage from  
292 June to August (mainly winter), for the growth of brachiaria and, at some extent, to other species.  
293 Prolonged drought conditions, as the rain deficit measured in successive years, can significantly  
294 reduce the water availability in the soil and lead to hydrological imbalances with significant  
295 higher evapotranspiration in relation to precipitation. The years 2014 and 2016 were  
296 characterized by lower minimum temperature and monthly rainfall compared to the climatology  
297 for both dry and rainy months (Table 1). This has resulted in significant deviation in annual  
298 precipitation accumulated and ETo/monthly precipitation.

299 During June and August,  $T_{\min}$  was extremely close to  $T_b$  of *B. brizantha* cv. Marandu. In  
300 these two years, the major discrepancy from climatology and the smallest  $T_{\min}$  occurred in  
301 August. In 2014, annual accumulated precipitation reached only 68.5% of the expected normal.  
302 The anomaly in precipitation reduced by 64, 82, 76, 87, 58, 67, 77 and 16% of the precipitation  
303 in the months of January, February (rainy months) and May, June, August, September (reduced  
304 rainfall period), October and December (rainy months), respectively. There was rainfall higher  
305 than the climatological mean only in October and December, and only in November and  
306 December the precipitation was higher than  $ET_o$  (Table 1). The behavior of the rains this year,  
307 allowed to measure intensified responses of the brachiaria, and to propose the experiment in  
308 lysimeters in order to improve our knowledge about the most critical period of the year, the  
309 transition between winter and spring.

310 In 2016, the annual rainfall has reached 99.3% of normal conditions and the reductions in  
311 monthly precipitation were 37, 32, 53, 33, 96, 28, 53 and 15% in the months of February, March,  
312 April, May, July, August, September and October, correspondingly. The greatest reductions in  
313 precipitation occurred in June (2014) and July (2016). Collected data of the relative  
314  $ET_o$ /precipitation rate in the months from May to October in 2014, and both July and August in  
315 2016, revealed that these year periods were those with the largest  $ET_o$ /precipitation unbalance.  
316 This implied also in greater soil draining in the year 2014 during the field experiment,  
317 particularly in the second production cycle.

318 The climatic evaluation showed that the minimum temperature ( $T_{\min}$ ) in 2014 and 2016  
319 approached the minimum cardinal temperature of the brachiaria, and the  $ET_o$  / precipitation ratio  
320 was anomalously disproportional in winter, in relation to the climatic pattern, which was  
321 favorable for our study.

322

## 323        **3.2 Field Experiment**

### 324        **3.2.1. Dry matter growth (DM) – 2014 experiment**

325        Dry matter above ground and leaf area index (Fig. 1 A, B) increased according to sigmoidal  
326        curves until the harvest in the two evaluated growing season, summer-fall and winter-spring.  
327        During the summer-fall period (05-Feb to 16-May/05), dry matter growth was higher than  
328        winter-spring (06-Jun to 11-Dec). The first harvest cut was carried out at 73 growing-days after  
329        grass cut equalization (GCE). In the winter-spring period, there were growth along the 209  
330        growing-days after the first harvest cut, which was separated by a short period of minimum plant  
331        growth (between July 28 and October 18), when a significant reduction in DM and LAI occurred  
332        (Fig. 1 B). After 18-Oct, a significant and continuous increase in DM and LAI are noted until the  
333        second harvest at 11-Dec (Fig. 1 A, B).

334        The maximum LAI values in the summer-autumn period were around 4.7, 59 days after  
335        GCE, and in the winter-summer period was 9.6 on day 209 after GCE. On the other hand, in the  
336        winter-spring period a double sigmoidal curve was observed, with a halt in LAI and biomass  
337        increase from July 28 to October 18 period of limiting low temperatures. This may be the  
338        greatest evidence that the lack of precipitation and consequent reduction of soil moisture are the  
339        determining factors in the dynamics of accumulation of the aerial biomass. In the subsequent  
340        period after the reduced LAI, these values increased almost linearly until the second harvest cut,  
341        coherently with the dry matter and culture growth rate. This fact was coincident with the year  
342        period where the highest precipitation and temperatures occurred (Table 1).

343        In the 2014 experiment, the amount of rain at the beginning of the rainy season (September-  
344        November) was lower than the climatological values (Table 1). The significant decrease in  
345        precipitation along with the high evapotranspiration recorded from May to October were crucial  
346        factors for the significant reduction of LAI and the growth rates (Fig. 1 B, C, D), which may  
347        have conditioned the observed delay in the regrowth of the brachiaria after 14-Aug. The growth

348 of biomass was only restored after 18-Oct, with increased precipitation and temperatures, two  
349 months after the end of low temperatures.

350 In consequence, the relative growth rate (RGR) and growth rate of the culture (GRC) of the  
351 summer-fall period were clearly higher than the winter-spring transition period. The RGR  
352 represents the efficiency of growth with respect to the initial biomass, while GRC expresses the  
353 productivity efficiency of the dry matter. Therefore, low values of these indices reflect the  
354 periods in which biomass growth was limited by climate.

355 Apparently, the low temperatures would have less importance than the water availability  
356 during the winter-spring and spring transition in the biomass production of the cultivar Marandu.  
357 In view of this hypothesis, the need to determine the temporal hierarchy of water deficit in the  
358 soil over low temperatures during the winter-spring transition period was identified.

### 359 **3.2.2. Air and soil temperatures influences**

360 Mean soil CO<sub>2</sub> efflux in summer-fall period was  $2.0 \pm 0.5 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$  and  $2.1 \pm 0.4$   
361  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$  in winter-spring period. Although there were prominent changes of efflux in  
362 each evaluated season, the values were consistent with CO<sub>2</sub> measurements described by Chang et  
363 al. (2013) and Dendy and Li (2010), who reported values between 2.3 and 6.9  $\mu\text{mol m}^{-2} \text{ s}^{-1}$ ,  
364 respectively. In addition, Varella et al. (2004) found high seasonal variation of CO<sub>2</sub> efflux in  
365 soils with native brachiaria in the Cerrado, and cultivated pasture, with mean values of 2.0  $\mu\text{mol}$   
366  $\text{m}^{-2} \text{ s}^{-1}$  in the dry season.

367 At the beginning of summer-fall period, the CO<sub>2</sub> efflux followed the minimum air  
368 temperature between 05-Feb and 26-Feb. From 26/02 onwards, there was a decoupling between  
369 temperature trend and NCER values, which increased up to 19/03. The rise in soil CO<sub>2</sub> efflux  
370 coincided with the exponential growth phase of dry matter recorded between 28-Feb to 19-Mar  
371 (Fig. 1 A). From 19-Mar, there was a decrease in NCER from 2.5 to 0.5  $\mu\text{mol.m}^{-2} \text{ s}^{-1}$  coinciding  
372 with the stationary phase of dry matter curve (Fig. 1 A), and the decrease of minimum air

373 temperature (Fig. 2). This was 30 days after the first harvest. After this period, the CO<sub>2</sub> efflux  
374 rose smoothly until a maximum efflux in 28-May (Fig. 2 B). From this date, the CO<sub>2</sub> efflux  
375 decreased continuously until 24-Jul. In this period, decoupling between low temperatures and  
376 CO<sub>2</sub> efflux values was evident. While the temperatures decreased, CO<sub>2</sub> efflux maintained  
377 minimum values until 24-Jul.

378 The lower levels of CO<sub>2</sub> efflux coincided with continuous biomass growth until 28-Set (Fig.  
379 1 A), despite the expressive decrease in temperature at 25-Jun, implying that the temperature  
380 drop did not hamper the Marandu growth. In addition, it could be inferred that other factors  
381 rather than temperature would hold the CO<sub>2</sub> efflux at its lowest level.

382 After 24-Jul, spikes of CO<sub>2</sub> efflux were observed in 3-Jul, 4-Set, 21-Oct, 23-Oct and 27-Nov.  
383 In the same period of efflux peaks observed in the winter, LAI and growth rates reduced  
384 significantly between 8-Aug and 18-Set, phenomena that cannot be explained by the variation in  
385 temperature.

386 During winter days, it is known that roots and lower shoots may have reduced respiratory  
387 rates due to low temperatures, which according to the temperature coefficient  $Q_{10}$  (Van't Hoff,  
388 1898, Taiz and Zeiger, 2010) can limit the metabolic rate of soil organisms and roots,  
389 consequently, the soil CO<sub>2</sub> efflux. In addition, low temperatures can also reduce the supply of  
390 energy for many physiological processes such as photosynthesis and transport of photosynthates  
391 to the roots and thus the respiratory rate of roots and microorganisms (Larcher, 2006, Taiz and  
392 Zeiger, 2010). According to Boone et al. (1998), the metabolic activity of roots strongly  
393 influences the sensitivity of soil respiration to temperature. Root respiration plus carbon  
394 oxidation in the rhizosphere contribute to the majority of CO<sub>2</sub> emitted by the soil. If cardinal low  
395 temperature limit is reached, the expected effect would be a drastic reduction of vegetative  
396 growth (Larcher, 2006).

397 In the period with the lowest temperatures, after 25-May, the CO<sub>2</sub> efflux remained almost  
398 constant in the lowest level already reached, with only small peaks and large peaks on 31-Jul  
399 during winter season. In the same period, significant and abrupt drops of air and soil  
400 temperatures occurred in 25-Jun, 16-Jul, 31-Jul and 7-Aug. However, on July 31 (Fig. 4c), a  
401 peak of CO<sub>2</sub> efflux occurred with the drops of temperatures. Analyzing the occurrence of  
402 precipitation from that date to the past, we find the recording of a 10 mm rainfall occurred on 25-  
403 Jul, six days before the occurrence of the CO<sub>2</sub> peak on 31-Jul. After 7-Aug, all the temperatures  
404 of the soil and air had increased, and from 28-Aug there was also a manifest growth of biomass,  
405 characterized by increases of LAI values (Fig. 1B). Leaf area index is an important indicator of  
406 land leaf cover and their productive potential through photosynthesis.

407 By the other hands, the decline in the biomass growth rate was not observed in the period of  
408 low temperatures. In fact, the near surface (data not showed) and soil (Fig. 3a) temperatures in  
409 three different periods of summer-fall, winter and spring-summer are poorly correlated with CO<sub>2</sub>  
410 efflux values. It is noted that low temperatures at the experimental location are not restrictive to  
411 increase the productivity of pastures, even with the occurrence of events of minimum  
412 temperatures below the Tb of the brachiaria (Fig 2a). This differs from the boreal region where  
413 low temperatures and the presence of snow strongly limit the production of grasses throughout  
414 the year (Davidson et al., 2000).

### 415 **3.2.3. Precipitation influence**

416 The results of the efflux peaks after rainfall events more temperature and CO<sub>2</sub> efflux  
417 evolution (Fig. 4 A) were separated into two graphs (Fig. 4 B, C). According to Fig. 4 B, there  
418 was accumulated rainfall of 77 mm in the summer/fall period during three days (from 6 to 8  
419 March). Three days after this rainfall event, there was a maximum of CO<sub>2</sub> efflux on 11-March. A  
420 second peak was recorded on 19-Mar, six days after a rainfall of 11.7 mm that has occurred on  
421 13-March. After March 19, a rainfall of less than 2 mm was recorded, and the values of

422 precipitation, temperature and CO<sub>2</sub> efflux decreased until the harvest cut on May 16 (Figure 4  
423 A). In ten weeks after harvesting, small increases in CO<sub>2</sub> flow were followed days after each  
424 occurrence of rainfall, even with minimum temperatures below 11 °C (Figure 4 A). The week of  
425 July 31 was an expressive increase in CO<sub>2</sub> efflux, after 11.1mm of rain on July 26, and other of  
426 different intensity after rains in August 19, 28 and September 2 (Fig. 4 C).

427 In hot and humid periods, similar CO<sub>2</sub> efflux behavior in response to rain events were also  
428 noticed, previously, in 11-March, and later, 02 and 23 of October with rains of 7.0, 9.3, and 12,3  
429 mm in 30 of September and October 20 and 26 respectively (Fig. 4 A). After October, other  
430 peaks with lower intensity and greater efflux response followed one another because soil  
431 moisture increased with or increased frequency of rainfall. Thus, small rainfall contributions  
432 were significant in the CO<sub>2</sub> efflux. The CO<sub>2</sub> efflux peaks after the rain events have also been  
433 verified by the correlation analysis (Fig. 3 D, F). Similar positive correlations between CO<sub>2</sub>  
434 efflux and water availability has been found previously (Pinto-Junior et al., 2009; Valentini et al.  
435 2008; Liu et al., 2002), also at low temperatures as in the grasslands of northern Mato Grosso  
436 State (Pinto-Junior et al., 2009). There were CO<sub>2</sub> efflux increases on 23/10 and 27/11 of about  
437 four times in relation to the period before 16/10. This expressive increase after 16/10 (October to  
438 December) could be attributed to the characteristic intensification of precipitation at this time of  
439 the year ( $r^2 = 0.51$ ,  $p < 0.05$ ; Fig 4 A), but also it can be associated to the observed increase of  
440 temperature and biomass which doubled in the last 55 days.

441 Changes in soil temperature can influence up to 80% of the temporal variations in CO<sub>2</sub> flow,  
442 with soil moisture adequate, however, as the soil dries, the influence of soil moisture  
443 predominates over temperature (Grahammer et al, 1991; Janssens et al., 2000; Smith et al.,2003).  
444 The results of our study were similar to the aforementioned results for periods with more  
445 elevated precipitation, between October and November of 2014. Regarding CO<sub>2</sub> efflux variation  
446 according to the studied variables, precipitation influenced significantly in CO<sub>2</sub> soil efflux during

447 the different seasons of the year, including periods of low temperatures of soil and air from May  
448 to August (Winter/Spring). Additionally, CO<sub>2</sub> efflux increases were verified from November, in  
449 agreement with enhanced precipitation, biomass accumulation (Fig 1 A, B) and higher soil and  
450 air temperatures (Fig 2 A).

451 The 2014 results showed the importance of the low temperatures in the production of  
452 biomass and soil respiration. However, the cessation of growth after the cold period and the  
453 registration of spikes of CO<sub>2</sub> efflux after rainfall events, led us to hypothesize that the humidity  
454 of the soil would be the most important factor limiting to early recovery of the post winter  
455 pasture production. Therefore, a second study was carried out in lysimeters to better understand  
456 and prioritize the influence of the seasonal variation of soil temperature and humidity in the  
457 winter-spring transition period, which was decisive for the evolution of CO<sub>2</sub> efflux and biomass  
458 production.

## 459 **3.2 Lysimeter Experiment**

### 460 **3.2.1. Dry matter growth (DM) – 2016 experiment**

461 Dry matter production and LAI were influenced by water availability levels in the first cycle  
462 much more than during the second cycle when the water restriction was reduced with rainfall  
463 (Fig 5 A, B). The production of biomass per plant during the first cycle (Fig 5 A), result in a  
464 production of 4900 kg ha<sup>-1</sup> to L1 treatment (control with 100% of ETo). The reduction in water  
465 availability affected pasture growth in other treatments, being 3229, 2444 and 1929 kg DM ha<sup>-1</sup>  
466 in L2, L3 and L4 treatments, respectively. These values correspond to a decrease of 18, 32 and  
467 43% in relation to L1 treatment. Growth reduction in treatments L3 and L4 occurred since  
468 August, while in L2 treatments the growth rate started to decrease only in September (Fig. 5 A),  
469 a few days before the rainy season arrives and intensifies. This point is important to support our  
470 proposal to prevent rapid soil desiccation by irrigations during the fall with slips that are

471 fractions of crop evapotranspiration, which can shorten the brachiaria production cycle,  
472 enhancing the number of harvests and productivity throughout the year.

#### 473 **3.2.4. CO<sub>2</sub> efflux of lysimeter soil and climate conditions**

474 Soil temperatures were lower in July and August varying between 15.3 and 18°C (Fig. 6 A).  
475 Temperatures have gradually risen from September to an average of 23.3 °C. But, no significant  
476 differences of soil temperature have been found among treatments. In fact, increased vegetation  
477 cover in all treatments reduced the solar radiation amount at the soil surface

478 Soil moisture remained higher during the entire experiment in L1 treatment (Fig. 6 B) with  
479 values above 60% of the available water capacity (AWC). In other treatments, the soil moisture  
480 gradually decreased due to evapotranspiration levels exceeding the water reposition. In the  
481 treatment L2, soil moisture reduced significantly around September, reaching 25% of AWC on  
482 September 17. On the other hand, in L3 and L4 assays, soil humidity was lower during almost  
483 the entire experiment, oscillating between 5 and 46% in the AWC of L3 treatment and between 2  
484 and 32% in the AWC of L3. From November 11, all treatments denoted soil moisture higher than  
485 50% of the AWC mainly due to rainfall occurrence in November.

486 These changes in AWC resulted in significant differences on soil CO<sub>2</sub> efflux (NCER) among  
487 treatments (Fig. 7 A). NCER showed higher rates in treatments with higher water availability.  
488 The NCER were 2.7, 2.4, 1.9 and 1.8  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$  for L1, L2, L3 and L4 treatments,  
489 respectively, corresponding to a reduction of 11, 30 and 34% of L2, L3 and L4 treatments in  
490 relation to L1. It is noteworthy that rain absence between 10/06 and 20/09 highlights the  
491 influence of soil moisture on CO<sub>2</sub> soil efflux, until before the onset of the rainy season.

492 Carbon dioxide efflux in treatments L3 and L4 was lower than L1 during the most of  
493 measurements (Fig. 7 A). However, the major discrepancies have occurred between August 25  
494 and September 17. This period coincided with the interval when L3 and L4 treatments  
495 experimented restriction in soil moisture, while soil temperature began to rise. In this phase,

496 minor biomass accumulation was also observed (Fig. 5 A). In turn, L1 and L2 CO<sub>2</sub> efflux were  
497 similar until mid-September, when soil moisture was still elevated. Nonetheless, when water  
498 levels in L2 treatment has fallen below 45% of AWC, a depletion on soil CO<sub>2</sub> efflux was  
499 observed.

500 In the second production cycle, starting in October, the CO<sub>2</sub> efflux increased in all  
501 treatments. Initially, this increment occurred due to the raise in soil temperature. Then, in  
502 November, the upsurge was fomented by soil moisture increase in every treatment resultant from  
503 augmented rainfall in November. In the last measurement of November, CO<sub>2</sub> efflux was above 4  
504  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$  in all assays.

505 We have found that soil temperature fluctuations also affect the CO<sub>2</sub> efflux. This could be  
506 perceived in treatments with greater water availability. The CO<sub>2</sub> efflux was lower during the  
507 Brazilian winter months, when soil temperature values decreased. The soil temperature and CO<sub>2</sub>  
508 efflux in treatment L1 were analogous (Fig. 6 A, Fig. 7 A) and the correlation between them was  
509 0.94 (Table 2).

510 In other treatments, correlations between soil temperature and CO<sub>2</sub> efflux were lower due to  
511 noticeable influence of lower water availability. Associations between CO<sub>2</sub> efflux and soil  
512 moisture were of 0.39; 0.12; 0.71 and 0.86 in L1, L2, L3 and L4 treatments, respectively. The  
513 correlation table (Table 2) showed that soil temperature was the main responsible for CO<sub>2</sub> efflux  
514 variation when the soil moisture was elevated.

515 The soil CO<sub>2</sub> emissions present complex nature; therefore, it is not possible to identify a  
516 single soil/environment attribute that explains, in isolation, its variation in time and space.  
517 Nevertheless, the great influence caused by moisture and temperature of soil exerted on CO<sub>2</sub>  
518 efflux is evident in this work. Furthermore, elevated moisture and temperatures offer proper  
519 conditions to CO<sub>2</sub> production, as they favor organic matter decomposition, roots respiration and

520 microbial respiration, thus increasing CO<sub>2</sub> emission from soil to atmosphere (Dias, 2006;  
521 Brandão, 2012).

522 Janssens et al. (2000) pointed out that soil temperature changes could influence on 80% of  
523 temporal variations of CO<sub>2</sub> flux, since adequate soil moisture is provided. Similarly, Smith et al.  
524 (2003) argued that CO<sub>2</sub> released by aerobic respiration is dependent on temperature, however,  
525 becomes reliant on moisture as soil dries.

526 Soil moisture tends to exhibit strong influence on efflux rates below or above extreme critical  
527 values of soil moisture (Silva et al., 2016). An elevated water content could prevent the diffusion  
528 of both O<sub>2</sub> and CO<sub>2</sub> in soil. In contrast, a limited water content in soil could inhibit soil  
529 microorganism activity and radicular respiration (Gonçalves; Pelegrine, 2015; Dias, 2006; Liu et  
530 al. 2002). In this work, critical soil moisture has been found close to 45% of the AWC, which  
531 fomented reduction of soil CO<sub>2</sub> efflux rates. Rosenthal et al. (1987) indicated values inferior to  
532 0.50 as factors responsible for significant reduction in plant growth and development.

533 Coser et al. (2008), has found increases of 24% in winter dry matter yield of elephant grass  
534 under nitrogen fertilization coupled with irrigation of 50% from evapotranspiration, in a region  
535 near to the site of this experiment. Similar findings were reported by Correa and Santos, 2006 in  
536 the state of Sao Paulo-Brazil with Tanzania grass, in which they verified that irrigation did not  
537 change the seasonality of production but promoted an increase in forage accumulation rate  
538 during spring and autumn. These authors also pointed out that the most pronounced effects of dry  
539 matter reduction were observed from September, especially in October, when there was no  
540 precipitation.

### 541 **3.2.5. CO<sub>2</sub> efflux of lysimeter soil and Net leaf assimilation of CO<sub>2</sub>**

542 Net assimilation rate of CO<sub>2</sub> in leaves (*A*), also known as leaf photosynthetic rate, was  
543 greater in treatments with more elevated water availability (Fig. 7 B), with values by about 28.3;  
544 25.4; 20.8 and 16.6 μmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup> for L1, L2, L3 and L4 treatments, respectively. This

545 corresponds to a reduction of 11, 27 and 42% in L2, L3 and L4 treatments in relation to the L1  
546 plot. It is noteworthy that these decreases were expected in plants subjected to water deficiency,  
547 which reflected strong stomatal adjustment to avoid excessive loss of water under stress  
548 conditions (Larcher, 2006; Romero and Botía, 2006).

549 In addition, the leaf net assimilation rate of CO<sub>2</sub> occurring in L3 and L4 treatments was  
550 inferior to L1 during the majority of measurements. Nonetheless, the sharpest reductions arose in  
551 September. In turn, the effects of lower water availability on L2 photosynthetic rates began in  
552 mid-September, when soil moisture was inferior to 45% of AWC, in line with data on soil CO<sub>2</sub>  
553 efflux rates (Fig. 6 B). Elevated soil temperature and soil moisture and (Fig. 6 A, B), favoring  
554 soil CO<sub>2</sub> efflux has also led to higher photosynthetic rates. The greater reduction in  
555 photosynthetic rates occurred on September 17, which coincides with driest soil period in L2, L3  
556 and L4 treatments. There were declines of 30, 88 and 99% of A values for L2, L3 and L4 assays,  
557 correspondingly.

558 The lower soil moisture between August and September in treatments with 50% or less of  
559 ETo compromised significantly the photosynthetic rates, biomass production and reduces soil  
560 CO<sub>2</sub> efflux. The partial irrigation could be an option as a management strategy to increase  
561 productivity.

562 CO<sub>2</sub> leaf net assimilation also showed elevated relation with CO<sub>2</sub> efflux (Fig. 7 A, B). The  
563 correlations between efflux data (NCER) and photosynthesis data (*A*) were 0.78; 0.75; 0.79 and  
564 0.85 in treatments L1, L2, L3 and L4. Several studies confirm the photosynthesis influence on  
565 soil efflux rates (Bahn et al., 2009; Kuzyakov and Gavrichkova, 2010). Photosynthesis affect the  
566 efflux of soil CO<sub>2</sub> through the substrate provision for respiration of both microorganisms and  
567 roots. Increasing evidences suggest that the supply of assimilates from photosynthetic active  
568 plant organs significantly modify root respiration (Luo and Zhou, 2006, Subke et al., 2009). In  
569 addition, photo-assimilates are directly exude through the roots and are rapidly used by

570 mycorrhizae and microorganisms inhabiting the rhizosphere and, thus, contributing to soil CO<sub>2</sub>  
571 formation (Merbach et al., 1999, Kuzyakov et al., 2003).

572 The contribution of CO<sub>2</sub> resultant from roots and rhizomicrobial respiration, known as CO<sub>2</sub>  
573 derived from root, is extremely elevated, reaching up to 90% of total soil CO<sub>2</sub> (Hanson et al.,  
574 2000). Several studies revealed that root CO<sub>2</sub> contribution decreased in periods of dormancy  
575 compared to the growing season (Dorr and Munichich, 1986; Rochette and Flanagan, 1997). This  
576 may be explained by the fact that these respiration types are closely related to plant assimilates  
577 supply into rhizosphere, therefore, directly dependent on photosynthesis.

578 Whichever factors that affect photosynthesis or the substrate supply to roots and rhizosphere  
579 microorganisms, such as solar radiation, leaf area index, water stress and nutritional status,  
580 represent an important determinant of root derived CO<sub>2</sub> efflux and therefore a contributing factor  
581 of soil CO<sub>2</sub> total efflux.

582 There is a highly variable time interval between carbon assimilation by photosynthesis and  
583 subsequent efflux of soil CO<sub>2</sub>. This depends on physiology, stage of plant growth and  
584 environmental conditions. In literature, delays of about 13 days for grasses and around 4-5 days  
585 for mature forest trees are found. Kuzyakov & Gavrichkova (2010) firmly discussed that soil  
586 CO<sub>2</sub> efflux models should incorporate parameters related to photosynthesis, as this variable  
587 represents one of the main drivers of carbon flows.

588 Our results combined with França da Cunha et al. (2012), strengthen the evidence that the  
589 lack of water in central and northern regions of Brazil may be the most important climatic factor  
590 for seasonal reduction of *B. brizantha* production. It is well known that seasonal variability  
591 accounts for 60% of the variation in annual production, while the interannual variability  
592 represents only 2.6% (Demanet et al., 2015). It is important to note that after the winter period,  
593 the availability of water in rivers for irrigation practices is minimal or null, whereas rivers and  
594 surface water sources are still abundant in the summer-fall period.

595        Considering irrigation of large areas of pasture, this process seems to be unfeasible, in spite  
596 of the occurrence of elevated evapotranspiration and significantly low precipitation in winter  
597 months, as shown in Table 1. However, discard the entire irrigation does not necessarily seem to  
598 be the most qualified strategy, considering the large reduction of low-cost food for livestock.  
599 Therefore, finding a half-term that guides on the amount and frequency of irrigation, when there  
600 is still water availability in the rivers, and whether that is economically viable in the annual  
601 balance of profitability of cattle production appear as interesting next steps. On the other hand, it  
602 will also be necessary to assess in what magnitude the irrigation or regional changes in the  
603 pluviometric regime can influence on annual carbon flow of these ecosystems with *B. brizantha*.  
604 Moreover, it would be important to elucidate whether the monoculture and current management  
605 techniques of Brachiaria pasture lands are taking into account productive sustainability.

606        Dantas et al., 2016, found similar results for the same cultivar of Brachiaria in São Paulo  
607 State-Brazil. These authors verified that the 100% replacement of the evapotranspiration or the  
608 maintenance in 50% of the water available in the soil, allows to effectively obtaining high  
609 production during the winter. Other results in irrigated conditions also showed increases of *B.*  
610 *brizantha* cv. Marandu of 1.65, 1.70 and 2.9 times in May, August and September, respectively,  
611 compared to the lowest production in April (Silva and Silva Junior, 2009). This can thus allow  
612 the rapid recovery and pasture production after the coldest (July) and dried period of the year  
613 (July to October).

614        Our results show that only the equivalent of 50% of the evapotranspiration could be needed  
615 since the irrigations are done during the Fall when there is still enough water in the reservoirs.  
616 Maintaining the amount of water available in the soil slightly above of 35% can be a generalized  
617 criterion to increase the productivity of *B. brizantha* cv Marandu during autumn and winter, and  
618 consequently increase annual productivity.

619 The experiment in lysimeter allowed observing the pre-eminence of the water deficit over  
620 temperature not crecimiento da brachiaria. The interaction of the two factors, deepening of the  
621 dry and low temperatures strongly limited the increase of biomass during the winter, mainly in  
622 treatments with high water deficiency. This effect can also be inferred that both factors interfere  
623 with photosynthesis and the metabolism of plants by different mechanisms, resulting in lower  
624 growth rates, and with the effect of drying much longer and more intense than low temperatures

#### 625 **4. Conclusions**

626 There was differential influence of climatic factors, regarding time of year, on the soil  
627 respiration cultivated with *Brachiaria brizantha* cv. Marandu. Even at low and moderate  
628 temperatures, precipitation was crucial to increase the soil CO<sub>2</sub> efflux. Moreover, the growth rate  
629 of dry matter was influenced by rainfall events. Thus, a potential irrigation management in the  
630 period after August 15, when minimum temperatures increased, could influence the growth curve  
631 enabling more elevated growth rates. This would reduce the cycles of harvest and enhance forage  
632 productivity during the year either for grazing or silage.

633 There was a differential influence of climatic factors on the respiration of the soil cultivated  
634 with *Brachiaria brizantha* cv. Marandu depending on the year period. The most important  
635 climatic factor for NCER variation was precipitation taking into account the preeminent  
636 correlation with efflux, mainly in the driest period of the year (May to August). Furthermore, soil  
637 temperature also impact on CO<sub>2</sub> efflux variation, since adequate moisture was presented in the  
638 soil.

639 Although precipitation exhibited more significant correlation with CO<sub>2</sub> efflux, the isolation  
640 of a single determining factor in CO<sub>2</sub> efflux variation is not correct, considering that all elements  
641 act in conjunction with Marandu cultivar development in soil respiration.

642 The controlled variation of soil moisture at low temperatures (Lysimeter experiment) allowed  
643 for a more precise characterization of the response of the biomass and CO<sub>2</sub> efflux from the soil,

644 and consequently, to define a management proposal to mitigate seasonality on pasture  
645 production. The results of the lysimeter test contributed to determine the soil moisture hierarchy  
646 on the low temperatures during the winter, and these findings were fundamental to define the  
647 management strategy aiming at the mitigation of the seasonality on the production of grass.

648 Based on the results of 2014 and the new findings in the evaluation of lysimeter in 2016, we  
649 can define a strategy for the mitigation of seasonality in the availability of food for livestock. We  
650 consider it necessary to irrigate during the fall, when water is still available for irrigation in  
651 rivers and reservoirs. This would avoid the continuous drying of the soil and a faster reaction of  
652 the plants with the beginning of the first precipitations of the new water cycle accompanied by  
653 higher temperatures. This strategy, although it cannot be generalized, would reduce the effects of  
654 the winter period on the annual *Brachiaria* harvest cycles, which would contribute to improving  
655 the productive management of the pastures.

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663

## 664 **5. References**

665 Alencar, C.A.B., Cunha, F.F., Martins, C.E., Cóser, A.C., Rocha, W.S.D., Araújo, R.A.S.  
666 Irrigação de pastagem: atualidade e recomendações para uso e manejo. R. Bras. Zootec., v.38,  
667 p.98-108, 2009.

668

669 Allen, R.G., Pereira, L.S., Raes, D., Smith, M. (1998) Crop evapotranspiration —guidelines  
670 for computing crop water requirements. Roma FAO. 300p. il (FAO. Irrigation and Drainage.  
671 Paper 56).

672 ANUALPEC (2008) ‘Anuário da pecuária brasileira’ (Informa Economics FNP: São Paulo)

673 ABIEC (2013) Estatísticas: balanço da pecuária. 2013. Associação Brasileira das Indústrias  
674 Exportadoras de Carnes. Available at: <http://www.abiec.com.br/texto.asp?id=8>

675 AMS, 2015 American Meteorological Society: Climatology. Glossary of Meteorology.  
676 [Available from <http://glossary.ametsoc.org/wiki/climatology>.] Accessed date Jun 2015.

677 Andrade, H.J., Brook, R., Ibrahim, M., 2008. Growth, production and carbon sequestration of  
678 silvopastoral systems with native timber species in the dry lowlands of Costa Rica. Plant Soil,  
679 308, 11-22.

680

681 Bahn, M., Kutsch, W., Heinemeyer, A., 2009. Synthesis: emerging issues and challenges for  
682 an integrated understanding of soil carbon dynamics. Soil carbon dynamics. An integrated  
683 methodology. Cambridge, UK: Cambridge University Press, in press.

684 Barbosa, R.A., Nascimento Júnior, D., Euclides, V.P.B.; Silva, S.C. da, Zimmer, A.H.,  
685 Torres Júnior, R.A.A., 2007. Tanzanian grass subjected to combinations between intensity and  
686 grazing frequency. Pesquisa Agropecuária Brasileira, v.42, p.329-340.

687 Barcellos, A.O., Balbino, L.C., Stone L.F.,(2011). Crop-livestock-forestry integration-  
688 Brasília, DF: Embrapa 130p.

689 Beecher, M., Hennessy, D., Boland, T.M., McEvoy, M., O’Donovan, M. Lewis. E. (2013).  
690 The variation in morphology of perennial ryegrass cultivars throughout the grazing season and  
691 effects on organic matter digestibility. Grass and Forage Science. 70(1) p.19-29.

692

- 693 Boone, R.D., Nadelhoffer, K.J., Canary, J.D., Kaye, J.P., 1998. Roots exert a strong influence  
694 on the temperature sensitivity of soil respiration. *Nature* 396, 570-572.
- 695 Borges, T.A., Oliveira, F.A., Silva, E.M., Goedert, W.J., 2009. Evaluation of physico-  
696 hydraulic parameters of Red Latosol under grazing and under Cerrado. *Ver. Bras. de Eng. Agríc.*  
697 *Ambient.* 13, 18-25
- 698 Bueno, M.R., Teodoro, R.E.F., Alvarenga de, C.B., Gonçalves, M.V., 2009 . Determination  
699 of the crop coefficient for tanzania grass. *Bioscience Journal*, v. 25, n. 5, p.29-35.
- 700 Correa, L.A. and Santos P.M. Manejo e utilização de plantas forrageiras dos gêneros  
701 *Panicum*, *Brachiaria* e *Cynodon*. Embrapa Pecuária Sudeste. Sao Carlos, SP. 2003. 36p.
- 702 Chambers, J. Q., Tribuzy, E. S., Toledo, L., Chispim, B. F. O., Higuchi, N., Santos, J.,  
703 Araujo, A. C., Kruijt, B., Nobre, A. D., Trumbore, S. E. (2002). Respiration from a tropical  
704 forest ecosystem partitioning of sources and low carbon use efficiency. *Ecological Application*,  
705 14:S72-S88.
- 706 Chang, Z., Liu, X, Feng, Q., Zhang, X., 2013 Temporal variation of soil CO<sub>2</sub> efflux on  
707 sloping pasture of Heihe River basin and effects of temperature and soil moisture. *J Geol.*  
708 *Geosci.* 2, 111. doi:10.4172/2329-6755.1000111
- 709 Cruz, P.G., Santos, P.M., Pezzopane, J.R.M., Oliveira, P.A., Araujo, L.C., 2011. Empirical  
710 models to estimate the dry matter accumulation of marandu grass with agrometeorological  
711 variables / Modelos empíricos para estimar o acúmulo de matéria seca de capim-marandu com  
712 variáveis agrometeorológicas. *Pesq. Agropec. Bras.*, 46, 675-681.
- 713 Davidson, B.A., Belk, E., Boone, R.D., 1998. Soil water content and temperature as  
714 independent or confound factors controlling soil respiration in a temperature mixed hardwood  
715 forest. *Global Change Biol.* 4, 217-227.

- 716 Dantas, G.F., Faria, R.T., Santos, G.O., Dalri, A.B., Palaretti, L.F., 2016. Produtividade e  
717 qualidade da brachiaria irrigada no outono/inverno . Eng. Agríc., 36, 469-481.  
718 <http://dx.doi.org/10.1590/1809-4430-Eng.Agric.v36n3p469-481/2016>
- 719 Davidson, E. A.; Savage, K.; Verchot, L. V.; Navarro, R., 2002. Minimizing artifacts and  
720 biases in chamber-based measurements of soil respiration. Agr. Forest Meteorol. 113, 21-37.
- 721 Davidson, E.A., Verchot, L.V., Cattáneo, J.H., Ackerman I., Carvalho, J.E.M., 2000. Effects  
722 of soil water content on soil respiration in forests and cattle pastures of eastern Amazonia.  
723 Biogeochemistry, 48, 53-69.
- 724 Demanet, R., Mora, M.L., Herrera, M.Á., Miranda, H., Barea, J.M., 2015. Seasonal variation  
725 of the productivity and quality of permanent pastures in Andisols of temperate regions. J. Soil  
726 Sci. Plant Nutr. 15, 111-128. <https://dx.doi.org/10.4067/S0718-95162015005000010>
- 727 Dendy, J., Li, Y., 2010. Soil CO<sub>2</sub> efflux in pastures and secondary forests on Hawaii Island.  
728 Pacific Agric. Nat. Resources 2, 9-11.
- 729 Dias, P.F., Souto, S.M., Correia, M.E.F.C., Rocha, G.P., Moreira, J.F., Rodrigues, K.M.,  
730 Franco, A.A., 2006. Tree nitrogen fixers and macrofauna of soil in Digitaria hybrid grassland.  
731 Pesquisa Agropecuária Brasileira, v. 41, n. 06, p. 1015-1021,
- 732 Dias-Filho, M. B., 2014. Recovery of degraded pastures in the Amazon: challenges,  
733 opportunities and perspectives. In: SAMBUICHI, R. H. R. et al. (Org.). Agri-environmental  
734 policies and sustainability: challenges, .opportunities and lessons learned. Brasília, DF: Ipea,p.  
735 149-169
- 736 Dixon, A.P., Faber-Langendoen, D., Josse, C., Morrison, J., Loucks, C.J., 2014. Distribution  
737 mapping of world grassland types. J. Biogeogr. 41, 2003-2019.
- 738 Dörr, H., Münnich K.O., 1986. Annual variations of the <sup>14</sup>C content in soil CO<sub>2</sub>  
739 Radiocarbon, 28 , pp. 338-345.
- 740

- 741 Dupas Dupas, E., Buzetti, S., Sarto, A.L., Hernandez, F.B.T., Bergamaschine, A.F. 2010.  
742 Dry matter yield and nutritional value of Marandu grass under nitrogen fertilization and  
743 irrigation in cerrado in São Paulo. *Revista Brasileira de Zootecnia*, Viçosa, MG, v.39, n.12,  
744 p.2598-2603.
- 745 FAO 2005. *Grasslands of the World . Plant Production and Protection Series No. 34*. Edited  
746 by J.M. Suttie, S.G. Reynolds and C. Batello. <http://www.fao.org/docrep/008/y8344e/y8344e00.htm>
- 747 Fang, C., Moncrieff, J.B., 2001. The dependence of soil CO<sub>2</sub> efflux on temperature. *Soil*  
748 *Biol. Biochem.*33, 155-165.
- 749 França da Cunha, F., Mota, R.M., Alencar, B.C. A., Oliveira, R.A., Cóser, A.C., Martins,  
750 C.E., Cecon, P.R., Araújo, S.R.A., 2012. Productivity of *Brachiaria brizantha* cv. Xaraés in  
751 different managements and doses of fertilization, rest periods and times of the year. *Idesia*  
752 (Arica), 30, 75-82. <https://dx.doi.org/10.4067/S0718-34292012000100009>
- 753 Frank, A.B., 2002 Carbon dioxide fluxes over a grazed prairie and seeded pasture in the  
754 Northern Great Plains. *Environ. Pollut.* 116, 397-403.
- 755 Gopalakrishnan, S., Watanabe, T., Pearse, S.J., Ito, O., Hossain, Z.A.K.M., Subbarao, G.V.,  
756 2009. Biological nitrification inhibition by *Brachiaria humidicola* roots varies with soil type and  
757 inhibits nitrifying bacteria, but not other major soil microorganisms. *Soil Sci. Plant Nutr.* 55,  
758 725-733 <https://doi.org/10.1111/j.1747-0765.2009.00398.x>
- 759 Grahammer, K., Jawson, M. D., Skopp, J., 1991. Day and night soil respiration from a  
760 grassland. *Soil Biol. Biochem.* 23, 77-81.
- 761 Hanson, P. J., Edwards, N. T., Garten, C. T., and Andrews, J. A., 2000. Separating root and  
762 soil microbial contributions to soil respiration: A review of methods and observations,  
763 *Biogeochemistry*, 48, 115–146, doi:10.1023/A:1006244819642, .
- 764 Hopkins, D.W., Waite, I.S., McNicols, J.W., Poulton, P.R., Macdonald, A.J., Donnell,  
765 A.G.O., 2009. Soil organic carbon contents in long-term experimental grassland plots in the UK

766 (Palace Leas and Park Grass) have not changed consistently in recent decades. *Global Change*  
767 *Biology*, 15, 1739–1754, doi: 10.1111/j.1365-2486.2008.01809.x .

768 Hunt, R., 1990. *Basic Growth Analysis: Plant Growth Analysis for Beginners*. Unwin  
769 Hyman, London, 1990, 112p.

770 IPCC - Intergovernmental panel on climate change, 2014. *Carbon and Other Biogeochemical*  
771 *Cycles*. Available from: <[https://www.ipcc.ch/pdf/assessment](https://www.ipcc.ch/pdf/assessment-report/ar5/wg1/WG1AR5_Chapter06_FINAL.pdf)  
772 [report/ar5/wg1/WG1AR5\\_Chapter06\\_FINAL.pdf](https://www.ipcc.ch/pdf/assessment-report/ar5/wg1/WG1AR5_Chapter06_FINAL.pdf)> Accessed in Oct 24, 2015.

773 IPCC 2007, *Climate change 2007: The physical science basis*. Tech. rep., Intergovernmental  
774 Panel on Climate Change, 18, Paris, summary for Policymakers. Contribution of Working Group  
775 I to the Fourth Assessment Report.

776 Jank, L., Barrios, S.C., Valle, C.B., Simeão, R.M., Alves, G.F. 2014. A The value of  
777 improved pastures to Brazilian beef production. *Crop & Pasture Science*.  
778 <http://dx.doi.org/10.1071/CP13319>

779 Janssens, I.A., Kowalski, A.S., Longdoz, B., Ceulemans, R., 2000 Assessing forest soil CO<sub>2</sub>  
780 efflux: An in situ comparison of four techniques. *Tree Physiol.* 20, 23-32.

781 Jones, S.K., R.M. Rees, U.M. Skiba, and B.C. Ball. 2005. Greenhouse gas emissions from a  
782 managed grassland. *Global and Planetary Change* 47:201-211.

783 Kang, S., Sueyoung, D., Lee, D., Lee, D., Jin, V. L., Kimball, J. S., 2003. Topographic and  
784 climatic controls on soil respiration in six temperature mixedhardwood forest slops,  
785 Korea. *Global Change Biology*, 9: 1427-1437.

786 Keller, D., Baisden, W.T., Timar, L., Mullan, B., Clark, A. Grassland production under  
787 global change scenarios for New Zealand pastoral agriculture. *Geosci. Model Dev.*, 7, 2359–  
788 2391, 2014. doi:10.5194/gmd-7-2359-2014.

789 Kuzyakov, Y., Gavrichkova, O., 2010. REVIEW: Time lag between photosynthesis and  
790 carbon dioxide efflux from soil: a review of mechanisms and controls. *Glob Change Biol.* 16:  
791 3386–3406.

792 Kuzyakov, Y., Leinweber, P., Sapronov, D., Eckhardt, K.U., 2003. Qualitative assessment of  
793 rhizodeposits in non-sterile soil by analytical pyrolysis *Journal of Plant Nutrition and Soil*  
794 *Science*, 166 , pp. 719-723

795 Larcher, W. (2006) *Ecofisiologia vegetal*. 3 ed. São Carlos: Rima. 550p

796 Li, F.Y., Snow, V.O., Holzworth, D.P., Johnson, I.R., 2010. Integration of a pasture model in  
797 APSIM. IN: Dove H, Colvenor Ra eds. Food security from sustainable agriculture. Proceedings  
798 of 15th Australian Agronomy Conference. Loncoln, New Zealand, 15-18 November 2010  
799 (doi:www.regional.org.au/asa/2010/farming-systems/simulation-decision-support/6993  
800 lifyrevis.htm).

801 Liu, X., Wan, S., Su, B., Hui, D., Luo, Y., 2002. Response of soil CO<sub>2</sub> efflux to water  
802 manipulation in a tallgrass prairie ecosystem. *Plant Soil* 240, 213-223.

803 Luo, Y., Jackson, R. B., Field, C.B., Mooney, H.A., 1996. Elevated CO<sub>2</sub> increases  
804 belowground respiration in California grasslands. *Oecologia* 108, 130-137.

805 Luo, Y. and Zhou, X. *Soil respiration and the environment*. San Diego, Elsevier, 2006. 316p.

806 Medek, D.E., Ball, M.C., Schortemeyer, M., 2007. Relative contributions of leaf area ratio  
807 and net assimilation rate to change in growth rate depend on growth temperature: comparative  
808 analysis of sub-Antarctic and alpine grasses. *New Phytologist*, 175, 290–300.

809 Merbach, W., Mirus, E., Knof, G., Remus, R., Ruppel, S., Russow, R., Gransee, A., Schulze,  
810 J., 1999. Release of carbon and nitrogen compounds by plant roots and their possible ecological  
811 importance. *J Plant Nutr. Soil Sci* 162:373–383

812 Munns, R., Schmidt, S., Beveridge, C. *Plant in action*. Edition 2. Australian Society of Plant  
813 Scientists, New Zealand Society of Plant Biologists, and New Zealand Institute of Agricultural

814 and Horticultural Science 2010–2016. <http://plantsinaction.science.uq.edu.au/content/about>  
815 (Access in March 2017).

816 O'Mara, F.P. The role of grasslands in food security and climate change. *Ann Bot.* 2012  
817 Nov; 110(6): 1263–1270. doi: 10.1093/aob/mcs209

818 Pedreira, C.G.S., Nussio, L.G, Da Silva, S.C., 1998 .Soil-climatic conditions for the  
819 production of *Cynodon* sp. In Simpósio sobre manejo da pastagem, 15., Piracicaba. Anais.  
820 Piracicaba: FEALQ, 1998. p85-113

821 Peichl, M., Leahy, P., Kiely, G., 2011. Six-year stable annual uptake of carbon dioxide in  
822 intensively managed humid temperate grassland Ecosystems, 14 pp. 112-126.

823 Pereira-Flores, M.E., Justino, F., Ruiz-Vera, U.M., Stordal, F., Melo, A.A.M., Rodrigues, R.  
824 de A., 2016. Response of soybean yield components and allocation of dry matter to increased  
825 temperature and CO<sub>2</sub> concentration. *Australian Journal of Crop Science.*  
826 Doi:10.21475/ajcs.2016.10.06.p7310. ISSN: 1835-2707.AJCS 10(6): 808-818.

827 Pinto-Junior, O. B., Sanches, L., Dalmon, A.C., Nogueira, J.S., 2009. Efflux of CO<sub>2</sub> from the  
828 soil in transition forest Amazonia Cerrado and pasture area. *Acta Amazonica*, 39, (4), 813-821.

829 Poorter, H., Garnier, E., 1996. Plant growth analysis: evaluation of experimental design and  
830 computational methods. *Journal of Experimental Botany* 47, 1343–1351.

831 Quintanilha, S.C., Hernandez, F.B.T., Vanzela, L.S., Lima, R.C., Upas, E., Buzetti, S.  
832 Respostas do capim Mombaça e Brizanta a irrigação. In: Congresso Anual de Iniciação  
833 Científica, 18., 2006, Jaboticabal. Anais... Jaboticabal, UNESP, 2006.

834 Raich, J.W., Schlesinger, W.H.,1992. The global carbon dioxide flux in soil respiration and  
835 its relationship to vegetation and climate. *Tellus B* 44, 81-99. doi:10.1034/j.1600-0889.1992.t01-  
836 1-00001.x

837

- 838       Ramírez, B.L., Ramírez, H.F., Suárez, J.C., 2009. Captura de carbono y desarrollo radicular  
839 de sistemas de uso del suelo en la Amazonia Colombiana. *Livestock Research for Rural*  
840 *Development*. 21, Article #91. Retrieved March 21, 2017, from  
841 <http://www.lrrd.org/lrrd21/6/rami21091.htm>
- 842       Ribeiro, A.C., Guimarães, P.T.G., Alvarez V.V.H., 1999. Recommendation for or use of  
843 corrective and fertilizers in Minas Gerais: 5. Viçosa: Comissão de Fertilidade do Solo do Estado  
844 de Minas Gerais, 359 p.
- 845       Roberts, J. M. Effects of temperature on soil respiration: a brief overview. Wallingford:  
846 Center for Ecology and Hydrology, 2000. 45p.
- 847       Rochette, P., Flanagan, L.B., 1997. Quantifying rhizosphere respiration in a corn crop under  
848 field conditions *Soil Science Society of America Journal*, 61,pp.466-474
- 849       Romero, P., Botía, P., 2006. Daily and seasonal patterns of leaf water relations and gas  
850 exchange of regulated deficit-irrigated almond trees under semiarid conditions. *Environmental*  
851 *and Experimental Botany*, v. 56, p. 158–173.
- 852       Rosenthal, W.D., Arkin, G.F., Shouse, P.L., 1987. Water deficit effects on transpiration and  
853 leaf growth. *Agron. J.*, 79:1019-1026.
- 854       Ruhoff, A. L., Saldanha, C. B., Collischonn, W., Uvo, C.B., Rocha, H.R.; Cabral, O.M.R.,  
855 2009. Multivariate analysis of evapotranspiração process in areas of closed and sugar cane.  
856 *Revista Brasileira de Recursos Hídricos*, 14(4), 137-146.
- 857       Rustad, L.E., Huntington, T.G., Boone, R.D., 2000. Controls on soil respiration: Implications  
858 for climate change. *Biogeochemistry* 48, 1-6.
- 859       Santos, S.M., Leal, L.M., Silva, M.M. Characterization of dry and rainy season for the city of  
860 Viçosa-MG. (1998). *Anais 1980-2006. Congressos Brasileiros de Meteorologia*.  
861 <http://www.cbmet.com/cbm-files/12-772ff471a91aa964c11bb28ec5826576.pdf>

862 Silva-júnior, J. A., Costa, A.C.L.; Azevedo, P.V., Costa, R.F., Metcalfe, D.B., Gonçalves,  
863 P.H.L., Braga, A.P., 2013. Soil CO<sub>2</sub> fluxes in the national forest of Caxiuanã, Pará, during the  
864 experiment ESECAFLOR/LBA. *Revista Brasileira de Meteorologia* Ribeira, 28( 1), 85-94.

865  
866 Silva, E.A., Silva, W.J., Barreto, A.C., Oliveira, J., Barbosa, A., Paes, J.M.V., Ruas, J.R.M.,  
867 Queiroz, D.S. (2012). Dry matter yield, thermal sum and base temperatures in irrigated tropical  
868 forage plants. *Revista Brasileira de Zootecnia*, 41(3), 574-582. [https://dx.doi.org/10.1590/S1516-](https://dx.doi.org/10.1590/S1516-35982012000300014)  
869 [35982012000300014](https://dx.doi.org/10.1590/S1516-35982012000300014)

870 Silva, B.M., Silva, E.A., Oliveira, G.C., Ferreira, M.M., Serafim, M.E., 2014. Plant-  
871 available soil water capacity: estimation methods and implications. *R. Bras. Ci. Solo*, 38:464-475

872 Silva, E.A., Silva, W.J., Barreto, A.C., Oliveira Junior, A.B., Paes, J.M.V., Ruas, J.R.M.,  
873 Queiroz, D.S. Dry matter yield, termal sum and base temperatures in irrigated tropical forage  
874 plants. *Revista Brasileira de Zootecnia*, v.41, n.3, p.574-582, 2012.  
875 <http://dx.doi.org/10.1590/S1516-35982012000300014>

876 Silva, W.J., Silva Júnior, L.C. Avaliação do efeito da soma térmica no crescimento e  
877 desenvolvimento de forrageiras tropicais. XVI Congresso Brasileiro de Agrometeorologia – 22 a  
878 25 de Setembro de 2009. Belo Horizonte – MG.  
879 [http://www.sbagro.org.br/anais\\_congresso\\_2009/cba2009/124.pdf](http://www.sbagro.org.br/anais_congresso_2009/cba2009/124.pdf)

880 Silva, C.M., Vasconcelos, S.S., Mourão Júnir, M., Bispo, C.J. C., Kato, O.R., Silva Junior  
881 A.C. da., 2016. Temporal variation of soil CO<sub>2</sub> efflux in agroforestry systems with oil palm May  
882 28 ; 46( 1 ): 1-12. Available in: [http://www.scielo.br/scielo.php?script=sci\\_arttext&pid=S0044-](http://www.scielo.br/scielo.php?script=sci_arttext&pid=S0044-59672016000100001&lng=pt)  
883 [59672016000100001&lng=pt](http://www.scielo.br/scielo.php?script=sci_arttext&pid=S0044-59672016000100001&lng=pt). <http://dx.doi.org/10.1590/1809-4392201500193>.(Accessed in  
884 July 2018)

885 Skinner, R.H., Corson, M.S., Rotz, C.A., 2007. Evaluating warm-season grass production in  
886 temperate-region pastures: a simulation approach. *Agricultural Systems* 93: 252-268.

- 887 Smith, K.A., Ball, T., Conen, F., Dobbie, K.E., Massheder, J., Rey, A., 2003. Exchange of  
888 greenhouse gases between soil and atmosphere: Interactions of soil physical factors and  
889 biological processes. *Eur J. Soil Sci.* 54, 779-791. [https://doi.org/10.1046/j.1351-](https://doi.org/10.1046/j.1351-0754.2003.0567.x)  
890 0754.2003.0567.x
- 891 Soussana, J. F., Tallec, T., and Blanfort, V., 2010. Mitigating the greenhouse gas balance of  
892 ruminant production systems through carbon sequestration in grasslands, *Animal*, 4, 334-350,  
893 doi:10.1017/S1751731109990784.
- 894 Souza, P.T., Banys, V.L.; Dias, M. et al. 2016. Tillering of *Brachiaria brizantha* cv. Marandu  
895 inoculated with *Azospirillum Brasilense* or fertilized with nitrogen. In: International Rangeland  
896 Congress, Saskatton. Proceeding. Saskatton: [s.n.] p.1-2
- 897 Subke, J.A., Vallack, H.W., Magnusson, T., Kell, S.G., Metcalfe, D.B., Högberg P., Ineson,  
898 P., 2009. Short-term dynamics of abiotic and biotic soil  $^{13}\text{CO}_2$  effluxes after in situ  $^{13}\text{CO}_2$   
899 pulse labeling of a boreal pine forest *New Phytol.*, 183 , pp. 349-357
- 900 Taiz, L. and Zeiger, E. (2010) *Plant Physiology*, Fifth edition. Sinauer Associates.  
901 Sunderland, MA. (In press), 690p.
- 902 Tonato, F. Determination of productive and qualitative parameters of *Cynodon* spp.  
903 depending on climatic variables. Piracicaba, 2003. 82p. Dissertação (Mestrado)- Escola Superior  
904 de Agricultura "Liz de Queiroz", Universidade de São Paulo.
- 905 Valentini, C. M. A., Gonçalves, A. J., Pelegrine, G. DE F. R., 2015. Breathing of soil as a  
906 bioindicator in degraded areas. *Revista Internacional de Ciências*, 5 dezembro . p. 127-143 DOI:  
907 10.12957/ric.2015.19581
- 908 Valentini, C.M.A., Espinosa, M.M., de Paulo, S.R., 2008. Estimate of  $\text{CO}_2$  efflux of soil, of a  
909 transition forest in northwest of Mato Grosso state, using multiple regression. *Cerne* 14, 9-16.

910 Valle C.B., Euclides V.P.B., Macedo, M.C.M. Selecting new Brachiaria for Brazilian  
911 pastures. In: International Grassland Congress, 19, 2001, São Pedro. Proceedings. Piracicaba:  
912 FEALQ, 2001.

913 Van't Hoff, J. H. (1898). Lectures on theoretical and physical chemistry. In Chemical  
914 Dynamics Part I(pp. 224–229). London: Edward Arnold.  
915 <https://archive.org/details/theoreticollectu02hoffrich>

916 Vanhala, P., 2002. Seasonal variation in the soil respiration rate in coniferous forest soils Soil  
917 Biol. Biochem., 34, pp. 1375-1379

918 Varella, R.F., Bustamante, M.M.C., Pinto, A.S., Kisselle, K.W., Santos, R.V., Burke, R.A.,  
919 Zepp, R.G.,

920 Viana, L.T., 2004. Soil fluxes of CO<sub>2</sub>, CO, NO, and N<sub>2</sub>O from an old pasture and from  
921 native savanna in Brazil. Ecol Appl., 14, S221-S231.

922 Vilela, L., Soares, W. V., Sousa, D. M. G. de; Macedo, M. C. M., 1998. Liming and  
923 fertilization for pasture in the cerrado region. Planaltina. Embrapa Cerrados. 16p

924 Von Randow, C., Manzi, A.O. Kruijt, B. Oliveira, P.J., Zanchi, F.B., Silva, R.L., Hodnett,  
925 M.G., Gash, J.H.C., Elbers, J.A., Waterloo, M.J., Cardoso, F.L., Kabat, P., 2004. Comparative  
926 measurements and seasonal variations in energy and carbon exchange over forest and pasture in  
927 South West Amazonia. Theor. Appl. Climatol., 78, 5-26. [https://doi.org/10.1007/s00704-004-](https://doi.org/10.1007/s00704-004-0041-z)  
928 0041-z

929 Wagai, R., Brye, K.R., Gower, S.T., Norman, J.M., Bundy, L.G., 1998. Land use and  
930 environmental factors influencing soil surface CO<sub>2</sub> flux and microbial biomass in natural and  
931 managed ecosystems in southern Wisconsin. Soil Biol. Biochem. 30, 1501-1509.

932 Waldron, B.L.,; Asay, K.H., Jensen, K.B., 2002. Stability and yield of cool-season pasture  
933 grass species grown at five irrigation levels. Crop science, v. 42, n. 3, p. 890-896.

934 Ward, S. E., Smart, S. M., Quirk, H., Tallowin, J. R. B., Mortimer, S. R., Shiel, R. S., Wilby,  
935 A. and Bardgett, R. D., 2016. Legacy effects of grassland management on soil carbon to depth.  
936 *Glob. Change Biol.* 22, 2929-2938. doi:10.1111/gcb.13246

937 Watson, D. J., 1947. Comparative physiological studies on the growth of field crops.  
938 Variation in net assimilation rate and leaf area between species and varieties, and within and  
939 between years. *Ann. Bot.-London* 11. 41-76.

940 Wingler, A., and Hennessy, D., 2016. Limitation of Grassland Productivity by Low  
941 Temperature and Seasonality of Growth. *Front. Plant Sci.* 7:1130. doi: 10.3389/fpls.2016.01130

942 Xu, J., Wu, L.S., Chen, W.P., Chang A.C., 2008. Simultaneous determination of  
943 pharmaceuticals, endocrine disrupting compounds and hormone in soils by gas chromatography–  
944 mass spectrometry *J. Chromatogr. A*, 1202, pp. 189-195

945 Zanchi, F.B., Rocha, H.R., Kruijt, B., Cardoso, F. L., Deus, J.A., Aguiar, L.J.G., 2003.  
946 Measurement of soil CO<sub>2</sub> efflux: monitoring with automatic cameras on forest and pasture in  
947 Rondônia. In: VI Congresso de Ecologia do Brasil, Fortaleza. Anais. Fortaleza-CE. p. 631-632.

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**Table 1:** Viçosa climate series of 1980-2013, 2014 and 2016 year growing seasons

Climate factor	Jan	Feb	Mar	Apr	Mai	Jun	Jul	Aug	Sep	Oct	Nov	Dec	
	Summer			Fall			Winter			Spring			S
Climatological serial data													
T <sub>max</sub> °C	28.6	29.6	28.6	27.3	25.2	24.2	24.2	25.5	26.2	27.2	27.3	27.9	
T <sub>min</sub> °C	19.0	18.8	18.4	16.8	14	11.9	11.2	12	14.2	16.4	17.7	18.6	
Pp (mm)	198.8	129.6	156.2	64.4	37.3	17.5	11.1	16.3	56.5	101.4	174.4	198.3	
ETo (mm)	122.8	119.6	111.1	89.6	72.1	61.1	69.8	88.0	98.3	114.1	111.5	116.5	
ETo/pp	0.6	0.9	0.7	1.4	1.9	3.5	6.3	5.4	1.7	1.1	0.6	0.6	
Meteorological data – 2014													
T <sub>max</sub> °C	30.9 <sup>+</sup>	30.9 <sup>+</sup>	28.1 <sup>-</sup>	27.3	25.8 <sup>+</sup>	24.9 <sup>+</sup>	23.8 <sup>-</sup>	26.1 <sup>+</sup>	28.3 <sup>+</sup>	28.8 <sup>+</sup>	28.5 <sup>+</sup>	29.3 <sup>+</sup>	
T <sub>min</sub> °C	23.4 <sup>+</sup>	18.3 <sup>-</sup>	18.2 <sup>-</sup>	17.1 <sup>+</sup>	13.4 <sup>-</sup>	12.7 <sup>+</sup>	12.8 <sup>+</sup>	11.6 <sup>-</sup>	13.9 <sup>-</sup>	15.4 <sup>-</sup>	18.0 <sup>+</sup>	18.8 <sup>+</sup>	
Pp (mm)	72.2 <sup>-</sup>	23.8 <sup>-</sup>	182.8 <sup>+</sup>	84.4 <sup>+</sup>	8.4 <sup>-</sup>	2.2 <sup>-</sup>	11.1	6.9 <sup>-</sup>	18.8 <sup>-</sup>	23.6 <sup>-</sup>	203.0 <sup>+</sup>	167.2 <sup>-</sup>	
ETo (mm)	131.7 <sup>+</sup>	126.9 <sup>+</sup>	102.1 <sup>-</sup>	85.3 <sup>-</sup>	93.3 <sup>+</sup>	66.9 <sup>+</sup>	72.7 <sup>+</sup>	92.9 <sup>+</sup>	114.6 <sup>+</sup>	129.4 <sup>+</sup>	111.9 <sup>+</sup>	125.5 <sup>+</sup>	
ETo/pp	1.82 <sup>+</sup>	5.33 <sup>+</sup>	0.56 <sup>-</sup>	1.01 <sup>-</sup>	11.10 <sup>+</sup>	30.39 <sup>+</sup>	6.55 <sup>+</sup>	13.47 <sup>+</sup>	6.10 <sup>+</sup>	5.49 <sup>+</sup>	0.55 <sup>-</sup>	0.75 <sup>+</sup>	
Meteorological data – 2016													
T <sub>max</sub> °C	28.0 <sup>-</sup>	31.0 <sup>+</sup>	29.1 <sup>+</sup>	29.5 <sup>+</sup>	25.5 <sup>+</sup>	22.7 <sup>-</sup>	25.0 <sup>+</sup>	26.3 <sup>+</sup>	27.7 <sup>+</sup>	26.8 <sup>-</sup>	26.9 <sup>-</sup>	28.4 <sup>+</sup>	
T <sub>min</sub> °C	19.0	18.7 <sup>-</sup>	18.5 <sup>+</sup>	16.0 <sup>-</sup>	14.2 <sup>+</sup>	12.1 <sup>+</sup>	10.9 <sup>-</sup>	10.8 <sup>-</sup>	14.1 <sup>-</sup>	16.7 <sup>+</sup>	17.4 <sup>-</sup>	17.9 <sup>-</sup>	
Pp (mm)	352.6 <sup>+</sup>	81.2 <sup>-</sup>	106.6 <sup>-</sup>	30.2 <sup>-</sup>	24.8 <sup>-</sup>	72.6 <sup>+</sup>	0.4 <sup>-</sup>	11.8 <sup>-</sup>	26.8 <sup>-</sup>	85.8 <sup>-</sup>	125.0 <sup>-</sup>	235.8 <sup>+</sup>	
ETo (mm)	99.1 <sup>-</sup>	118.5 <sup>-</sup>	97.5 <sup>-</sup>	95.4 <sup>+</sup>	68.5 <sup>-</sup>	52.0 <sup>-</sup>	68.9 <sup>-</sup>	91.0 <sup>+</sup>	96.8 <sup>-</sup>	92.3 <sup>-</sup>	90.9 <sup>-</sup>	113.2 <sup>-</sup>	
ETo/pp	0.3 <sup>-</sup>	1.5 <sup>+</sup>	0.9 <sup>+</sup>	3.2 <sup>+</sup>	2.8 <sup>+</sup>	0.7 <sup>-</sup>	172.3 <sup>+</sup>	7.7 <sup>+</sup>	3.6 <sup>+</sup>	1.1	0.7 <sup>+</sup>	0.5 <sup>-</sup>	

960 Values followed by (+) and (-) signifies superior and inferior values to Normal Climatological  
 961 conditions (NC), respectively.

962

963

964 Table 2. Pearson linear correlation of CO<sub>2</sub> efflux values in the soil (NCER) regarding soil  
 965 temperature and moisture in treatments with 100% (L1), 75% (L2), 50% (L3) and 25% (L4) of  
 966 evapotranspiration reposition of accumulated crop (ETc ac.).

Treatment	Soil temperature	Soil moisture
L1	0.84**	-0.39
L2	0.64*	0.12
L3	0.56*	0.71**
L4	0.60*	0.86**

968 \*\*Significant correlation at 1% level 1% (p<0.01)

969 \* Significant correlation at 5% level (p<0.05)

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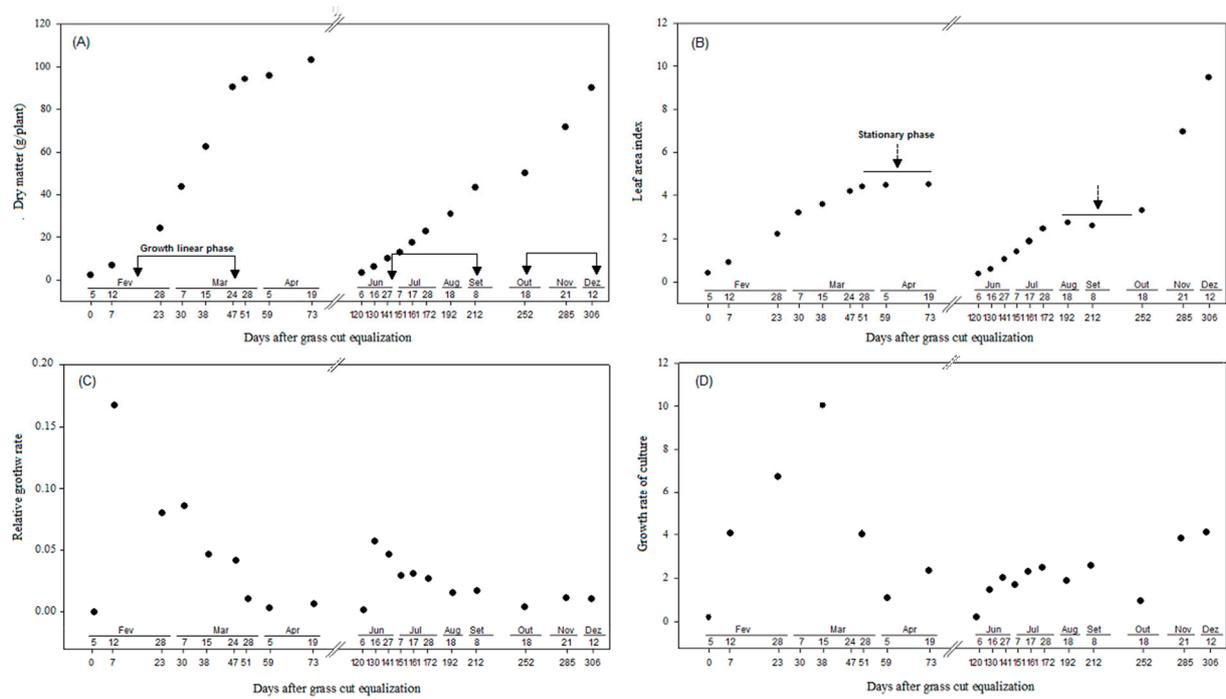
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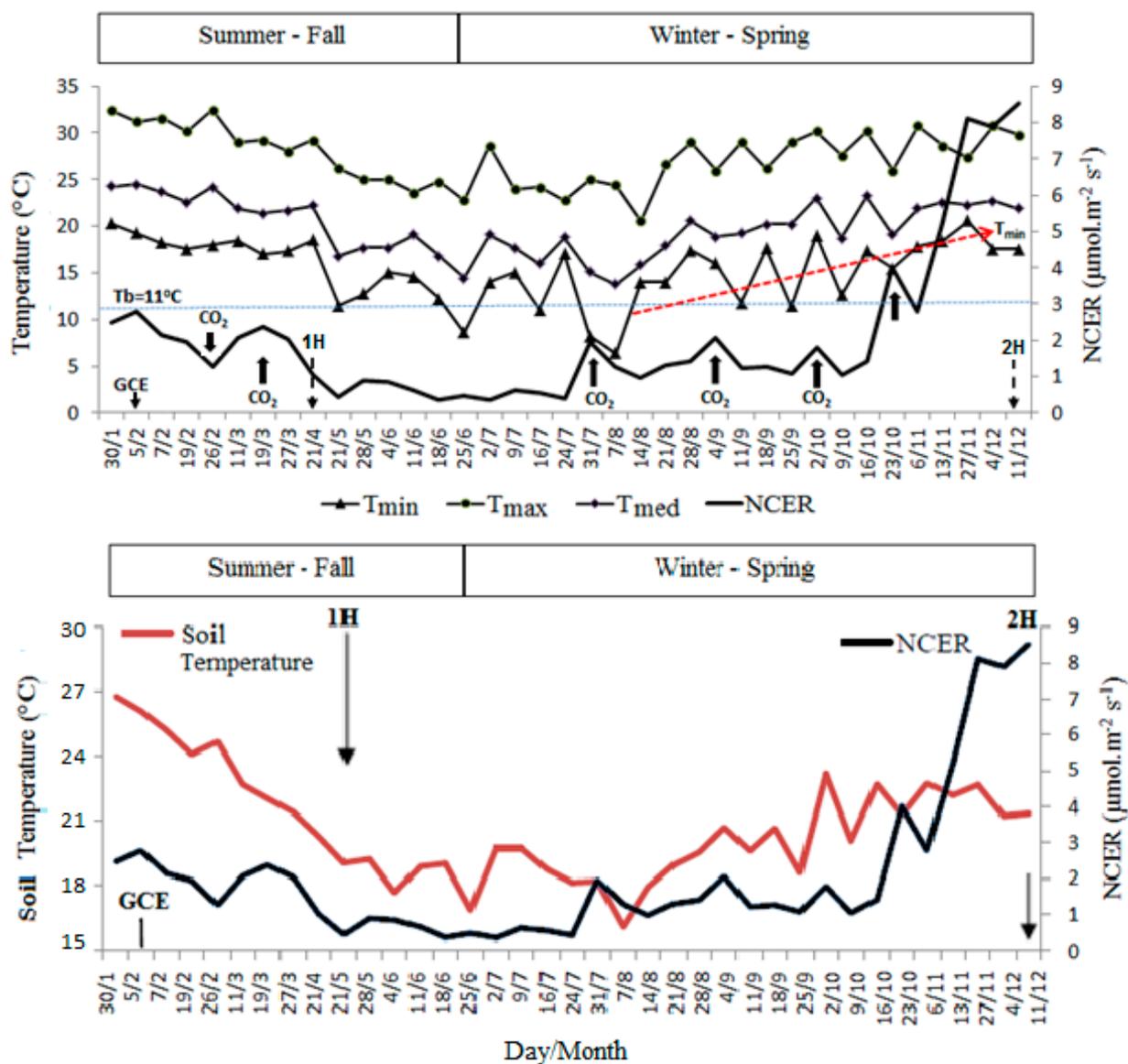
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981  
 982 **Figure 1.** Dry matter (A), Leaf area index (B), Relative (C) and Culture (D) growth rates of  
 983 *Brachiaria brizantha* cv Marandu, in two growth seasons of 2014. Solid arrow shows the initial  
 984 and final linear phase on each curve. Dotted arrows represent the onset or stationary phase of  
 985 curves.

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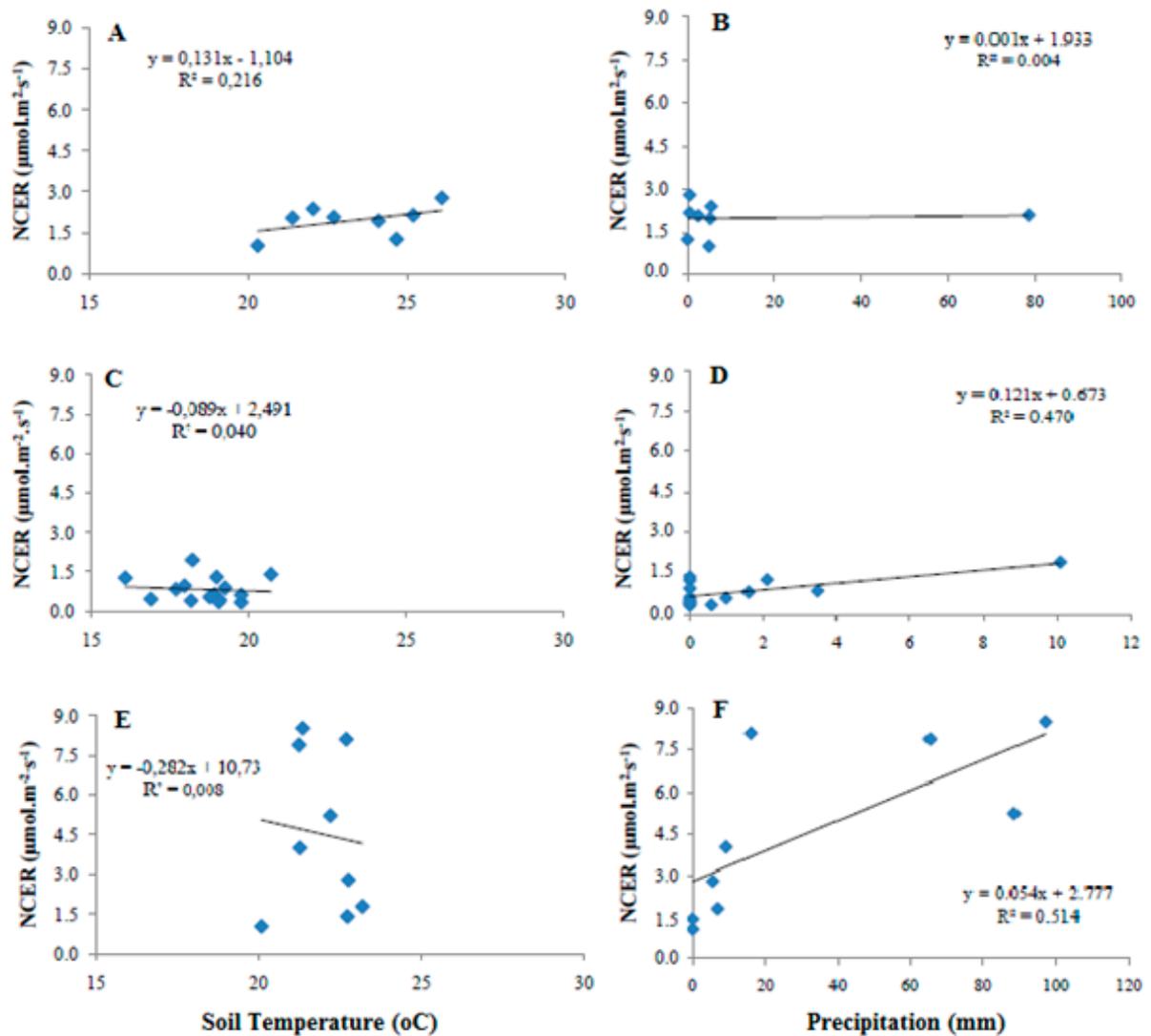


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989 **Figure 2.** Temperatures of air (A), soil (B) and NCER evolution on both evaluated growth  
 990 seasons. The black arrows denote change points in biomass management (GCE=grass cut  
 991 equalization, 1H=First harvest, 2H=Second harvest). The dotted red line shows the increase of  
 992 the minimum temperatures from 14 / Aug and the fine dotted blue line shows the minimum  
 993 cardinal temperature ( $T_b$ ) for the cv. Marandu (A).

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997 **Figure 3.** Correlation ( $P < 0.05$ ) between NCER ( $\text{CO}_2$  efflux) with Soil Temperature (A, C, E)  
 998 and between NCER and Precipitation (B, D, E) at February to April (A, B), May to August (C, D,  
 999 dry period) and October to December (E, F, rainy period).

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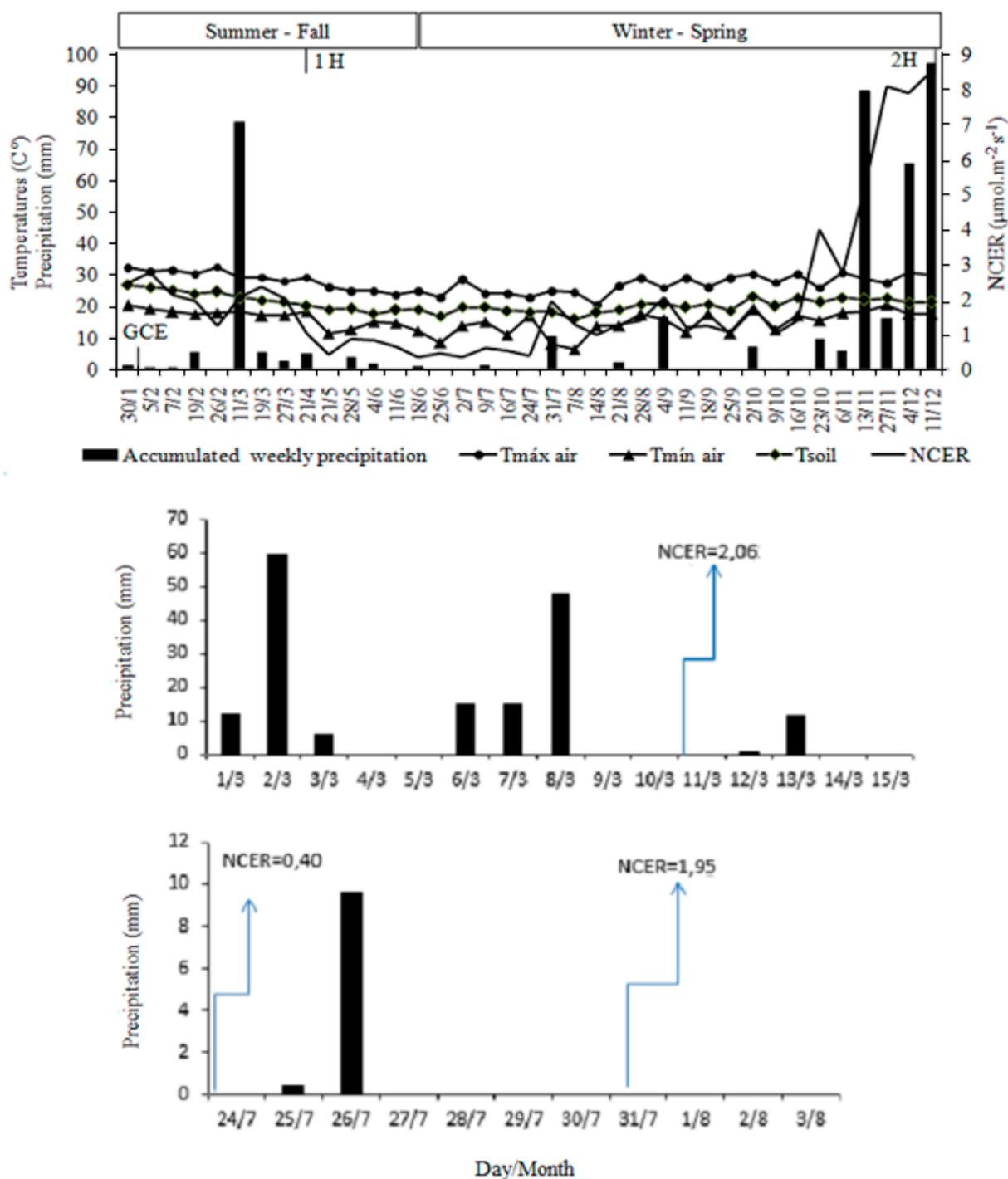
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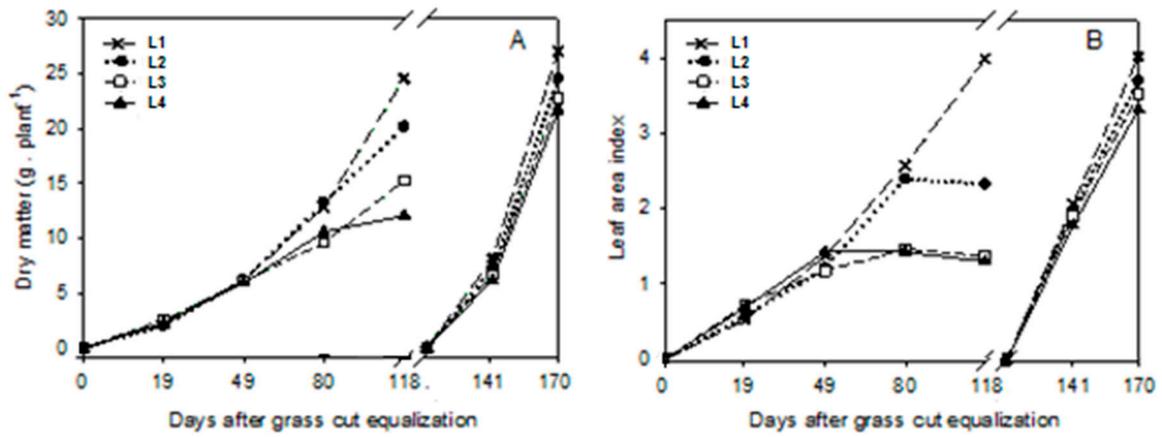
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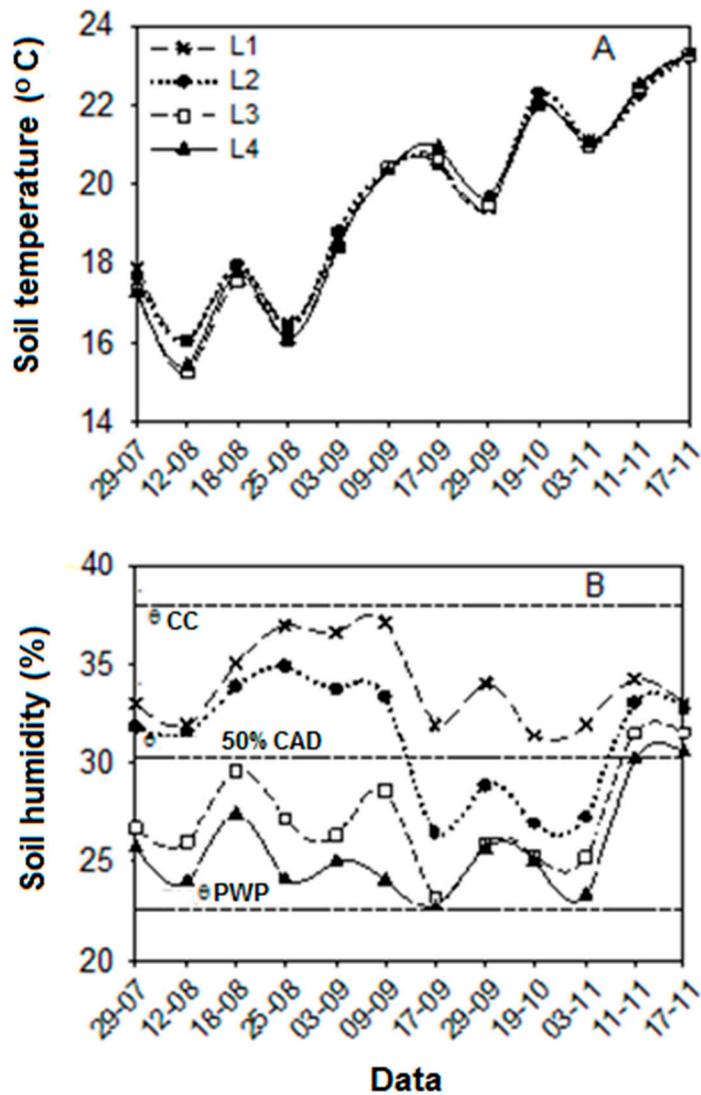
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1007 **Figure 4:** NCER (CO<sub>2</sub> efflux) evolution compared to precipitation, air and soil temperatures on  
 1008 growth seasons (up). Response time to NCER after rain event in contrast season, summer  
 1009 (medium) and winter (down). Blue arrows point out the day where ascension of CO<sub>2</sub> after rain  
 1010 events was found.

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1012  
 1013  
 1014 **Figure 5.** Dry matter accumulation (A) and leaf area index (B) from the standardization of  
 1015 *Brachiaria brizantha*, cv Marandu until the end of experimental period in treatments with 100%  
 1016 (L1), 75% (L2), 50% (L3) and 25% (L4) of evapotranspiration reposition of accumulated crop  
 1017 (ETc ac.).  
 1018  
 1019



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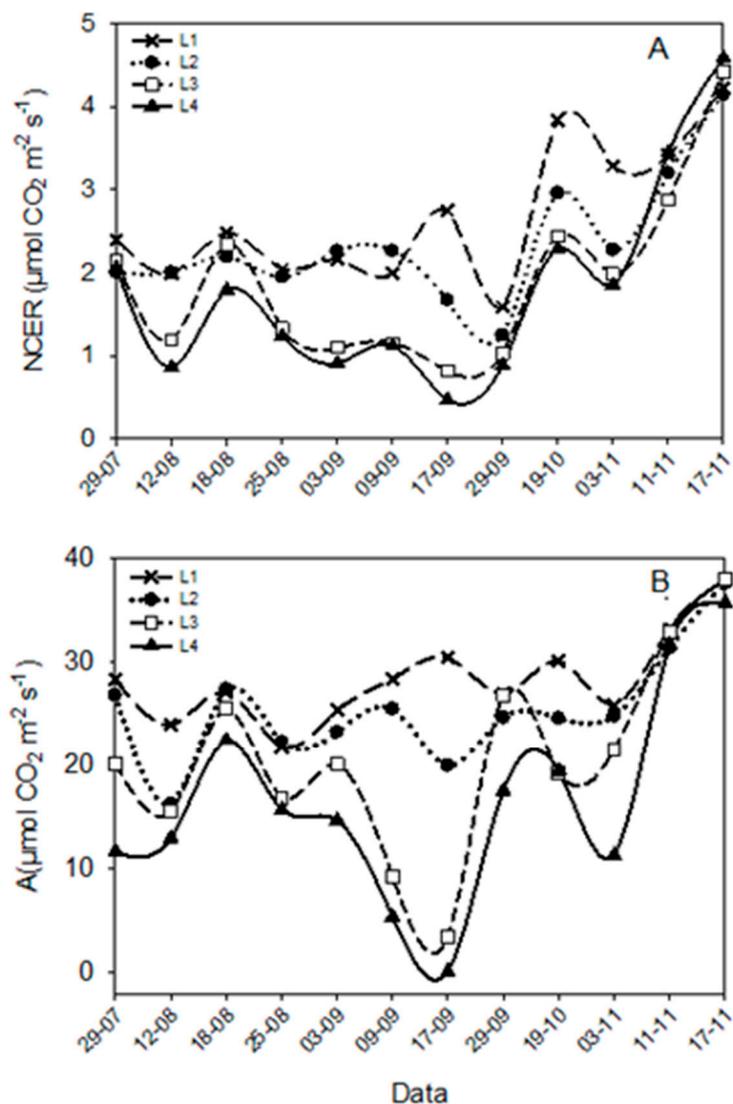
1021 **Figure 6.** Soil temperature and moisture in the 0-5 cm layer of areas containing *Brachiaria*  
 1022 *brizantha*, cv Marandu, submitted to treatments with 100% (L1), 75% (L2), 50% (L3) and 25%  
 1023 (L4) of accumulated culture evapotranspiration reposition (ETc ac.) (A).  $\theta_{CC}$  = soil volumetric  
 1024 moisture in field capacity;  $\theta_{50\% CAD}$  = 50% of available water capacity (AWC);  $\theta_{PWP}$  = soil  
 1025 volumetric moisture at the permanent wilting point (B).

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1031 **Figure 7.** Soil CO<sub>2</sub> efflux (NCER) in areas with *Brachiaria brizantha*, cv Marandu, submitted to  
 1032 treatments with 100% (L1), 75% (L2), 50% (L3) and 25% (L4) of accumulated culture  
 1033 evapotranspiration reposition (ETc ac.) (A). Leaf net assimilation of CO<sub>2</sub> (A) in areas of  
 1034 *Brachiaria brizantha*, cv Marandu, submitted to treatments with 100% (L1), 75% (L2), 50%  
 1035 (L3) and 25% (L4) of evapotranspiration reposition of accumulated crop (ETc ac.) (B).