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The Ground-Based Absolute Radiometric Calibration of Landsat 8 OLI

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Abstract: This paper presents the vicarious calibration results of Landsat 8 OLI that were obtained using the reflectance-based approach at test sites in Nevada, California, Arizona, and South Dakota, USA. Additional data were obtained using the Radiometric Calibration Test Site, which is a suite of instruments located at Railroad Valley, Nevada, USA. The results for the top-of-atmosphere spectral radiance show an average difference of -2.7 , -0.8 , 1.5 , 2.0 , 0.0 , 3.6 , 5.8 , and 0.7% in OLI bands 1–8 as compared to an average of all of the ground-based measurements. The top-of-atmosphere spectral reflectance shows an average difference of 1.6 , 1.3 , 2.0 , 1.9 , 0.9 , 2.1 , 3.1 , and 2.1% from the ground-based measurements. Except for OLI band 7, the spectral radiance results are generally within $\pm 5\%$ of the design specifications, and the reflectance results are generally within $\pm 3\%$ of the design specifications. The results from the data collected during the tandem Landsat 7 and 8 flight in March 2013 indicate that ETM+ and OLI agree to each other to within $\pm 2\%$ in similar bands in top-of-atmosphere spectral radiance, and to within $\pm 4\%$ in top-of-atmosphere spectral reflectance.

Keywords: Landsat 8; OLI; radiometric calibration; surface reflectance

1. Introduction

The Landsat Data Continuity Mission (LDCM) was launched on 11 February 2013 from Vandenberg Air Force Base, and it was renamed Landsat 8 after the transition from the National Aeronautics and Space Administration (NASA) to the United States Geological Survey (USGS). The platform contains two instruments: the Operational Land Imager (OLI), and the Thermal Infrared Sensor (TIRS), which complement one another in spectral coverage. OLI is a visible and near infrared (VNIR) multispectral sensor that operates from 400–2500 nm, and TIRS is a two-band thermal sensor that operates from 10.6–12.5 μm [1,2]. The in-flight absolute radiometric calibration of OLI is the focus of this work.

Landsat 8 is the latest platform in the 40-year Landsat series of satellites. The VNIR bands retain many of the characteristics of previous Landsat sensors such as a 185-km swath width, a 30-m ground instantaneous field of view (GIFOV) for the multispectral bands, and a 15-m GIFOV for the panchromatic band (Table 1). OLI was developed by the Ball Aerospace and Technology Corporation in Boulder, Colorado, USA. Significant changes were made in the electro-optical design as compared to the Enhanced Thematic Mapper Plus (ETM+) on Landsat 7. One prominent change was the transition from a whiskbroom configuration (ETM+) to a pushbroom configuration, which was successfully demonstrated by the Earth Observing One (EO-1) Advanced Land Imager (ALI) [3,4]. The pushbroom configuration allows OLI to have a signal-to-noise ratio (SNR) that is eight times greater than ETM+ when viewing typical radiance levels. OLI has 12-bit radiometric resolution, which is 16 times greater than the eight-bit data from ETM+. The higher SNR coupled with the higher bit level means that OLI is able to resolve smaller radiometric differences than ETM+. Although the pushbroom configuration has a much higher SNR, the radiometric calibration is more complex due to the approximately 7000 detectors in each of the multispectral bands.

Table 1. The specifications for the spectral bands of ETM+ and OLI used in this work. The center wavelength is determined by band averaging the relative spectral response (RSR) for each channel. Note that ETM+ band 6 is thermal, and it is not included here.

OLI		ETM+	
Band	Center Wavelength (nm)	Band	Center Wavelength (nm)
1	443.0	--	--
2	482.6	1	478.7
3	561.3	2	561.0
4	654.6	3	661.4
5	864.6	4	834.6
6	1609.1	5	1650.3
7	2201.3	7	2208.2
8 (Pan)	591.7	8 (Pan)	720.1

The top-of-atmosphere (TOA) spectral radiance and TOA spectral reflectance are two products available from OLI, and the calibration traceability follows two separate paths [4,5]. The preflight

radiometric calibration and SI traceability of OLI was completed using a 27.9-cm spherical integrating source (SIS), which was calibrated at the Facility for Spectroradiometric Calibration (FASCAL) at the National Institute of Standards and Technology (NIST). The measurements at NIST were also independently verified using transfer radiometers from NASA, NIST, the University of Arizona, and also with the Ball Standard Radiometer (BSR), which is spectrally controlled using filters that match the RSRs of OLI. When the SIS was returned to Ball Aerospace, it was measured again using Ball's radiometer and its own internal monitor. The radiance calibration was then transferred to the large SIS, nicknamed the Death Star Source (DSS), using Ball's Commercial-Off-The-Shelf (COTS) Transfer Radiometer (CXR). Upon completion of the calibration transfer, the DSS output was measured using the transfer radiometers that were previously used to measure the SIS. OLI was calibrated in a thermal vacuum chamber using the DSS by viewing the DSS through a window whose spectral transmission was previously measured in the laboratory. A heliostat facility at Ball Aerospace in Boulder was also used as a transfer-to-orbit calibration standard [6,7]. The reflectance calibration of OLI was performed using two sets of measurements. A transfer calibration panel was measured at the NIST Spectral Tri-function Automated Reference Reflectometer (STARR) facility, and the bidirectional reflectance factor (BRF) of the flight diffusers was measured using the goniometric facility at the Remote Sensing Group (RSG), which is part of the College of Optical Sciences at the University of Arizona. In addition to the absolute radiometric calibration, other prelaunch tests of OLI included linearity, stray light, and ghosting. An extensive description of the OLI instrument and its preflight calibration is described in [5,8].

Vicarious radiometric calibration provides an independent bridge that links the preflight and post-launch calibrations. This work describes the post-launch absolute radiometric calibration of OLI using a combination of techniques that include the reflectance-based approach, the Radiometric Calibration Test Site (RadCaTS), and an intercalibration with Landsat 7 ETM+ during the tandem flight period when both satellites were in similar orbits for three days. Section 2 describes the methodology of the reflectance-based approach, which includes the surface reflectance and atmospheric measurements, as well as a description of the radiative transfer process that is used to convert the surface measurements to a top-of-atmosphere quantity (e.g., spectral radiance or reflectance). Section 2 also describes the RadCaTS instrumentation and methodology, which is modelled on the reflectance-based approach. Section 3 describes the field work that occurred during the 14-month period reported here, and Section 4 presents the results for each of the three techniques. Section 5 provides a discussion on the uncertainty of the reflectance-based approach and presents a preliminary analysis of the uncertainty in RadCaTS. Finally, Section 6 provides a summary of the results.

2. Methodology

2.1. Reflectance-Based Approach to Vicarious Calibration

One technique used to perform the absolute radiometric calibration of airborne and spaceborne systems is the reflectance-based approach. It has been in use for over 25 years, and is a mature, well-understood procedure [9–12]. Measurements of the surface and atmosphere are made during a given overpass, and the results are used in a radiative transfer code to determine the TOA quantity for the sensor under test. In this work, the TOA spectral radiance and TOA reflectance are both reported since

they are standard products in the radiometrically- and geometrically-corrected L1T data that are available from USGS. It should be noted that it is important to compare the ground results to each of these products because their onboard calibration is determined using different pathways. The TOA spectral radiance is determined using the preflight calibration work that took place at the Ball Aerospace and Technology Corporation, while the TOA reflectance is determined by using the onboard solar diffusers.

Five test sites located in Nevada, California, Arizona, and South Dakota, USA were used in this work. Two sites that are routinely used by RSG are Railroad Valley, Nevada, and Ivanpah Playa, California. Alkali Lake, Nevada, is used as a supplementary test site when access to one of the other two is restricted. Railroad Valley has been used by RSG since 1996 for such sensors as Landsat 5 Thematic Mapper (TM), Landsat 7 ETM+, the Moderate Resolution Imaging Spectroradiometer (MODIS) onboard Terra and Aqua, the Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER), the Multiangle Imaging SpectroRadiometer (MISR), and the constellation of RapidEye satellites [11,13–15]. It is 15×15 km in size, and it is at an altitude of 1435 m. *In situ* data collected over the past 15 years have revealed that even through the surface reflectance changes throughout the year, it can generally be described as having an annual cyclical pattern. Ivanpah Playa has also been in use by RSG since 1996, and it is slightly smaller than Railroad Valley with a useable area of 2×10 km. It is at an altitude of 800 m, and has a hard surface with a flat topography. The Committee on Earth Observation Satellites (CEOS) has recognized Railroad Valley and Ivanpah Playa as instrumented radiometric calibration sites. Alkali Lake, Nevada, is the third site that is used by RSG in this work. It is used less frequently than Railroad Valley and Ivanpah, mostly when weather conditions are not favorable at the other two sites. The useable area at Alkali Lake is approximately 3×4 km, and it is at an altitude of 1467 m. Red Lake, Arizona, is the fourth site used by RSG for this work. It was originally used during a joint field campaign with the University of Arizona and NASA Goddard Space Flight Center (GSFC) to collect data during the three-day tandem flight of Landsat 7 and Landsat 8. During this three-day period, the two platforms were in similar orbits, and were able to view the same site within minutes of each other at similar viewing geometries. Railroad Valley, Ivanpah Playa, and Alkali Lake were not on the orbital track, so after a preliminary analysis it was determined that Red Lake would be suitable as a new site, based on its spatial and spectral characteristics. It is located 50 km north of Kingman, Arizona, with a useable area of approximately 6×6 km at an altitude of 845 m.

South Dakota State University (SDSU) typically uses a site for *in situ* measurements that is located in Brookings, South Dakota. It is a 100×250 -m area located at an altitude of 505 m, and it is surrounded by a larger grass area that is 300×500 m. The grass on the actual calibration site is routinely maintained during the spring, summer, and fall months so that changes in the bidirectional reflectance distribution function (BRDF) due to the structure of the grass is minimized. The SDSU test site was on the path of Landsat 7 and Landsat 8 during the three-day tandem flying period in March 2013, and data were successfully collected by SDSU for both sensors. Landsat 8 OLI true-color images of the five test sites used in this work are shown in Figure 1, where all of the sites are on the approximate same scale. In addition, Figure 2 shows the surface BRDF of each site for typical summer solar illumination angles as well as the BRDF of the site normalized to a solar zenith angle of zero.

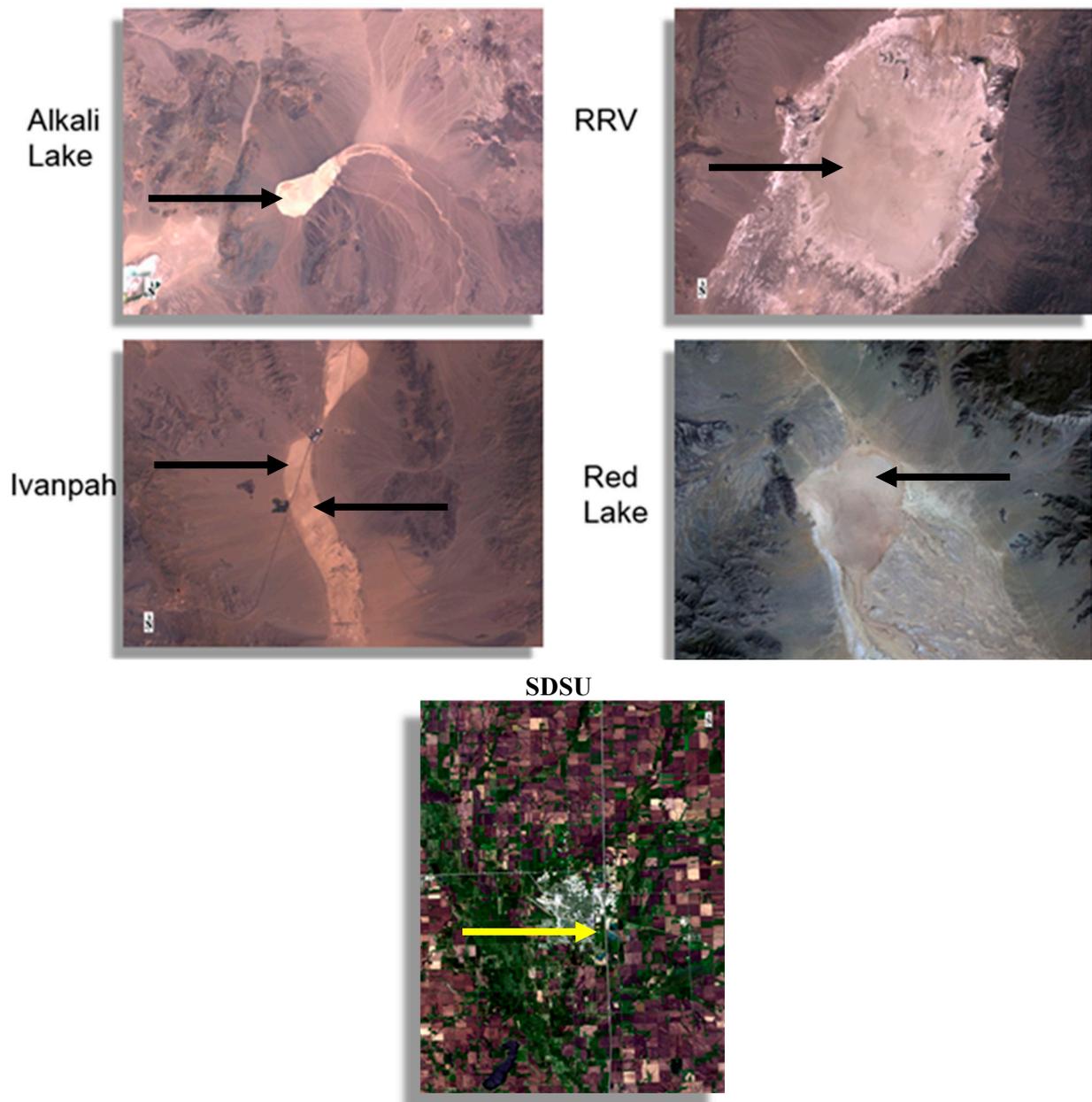


Figure 1. The test sites used to calibrate Landsat 8 OLI. All of the sites are on approximately the same scale, and the true-color images are created using bands 2–4 of OLI. The arrows show the general area(s) where surface reflectance measurements were made at each test site.

2.1.1. Surface Reflectance

The surface BRF is measured using a portable hyperspectral spectroradiometer (350–2500 nm) that is carried on the site in a predetermined pattern. The foreoptic is mounted to a boom arm that keeps it away from the user, thereby ensuring that the surface being measured is free of shadows. The main unit is carried in a backpack, and the computer is carried on platform in front of the user (Figure 3). Measurements of a Spectralon reference panel are made at predetermined points throughout the site collection, and the BRF is determined by taking the ratio of output digital numbers (DNs) over the two targets.

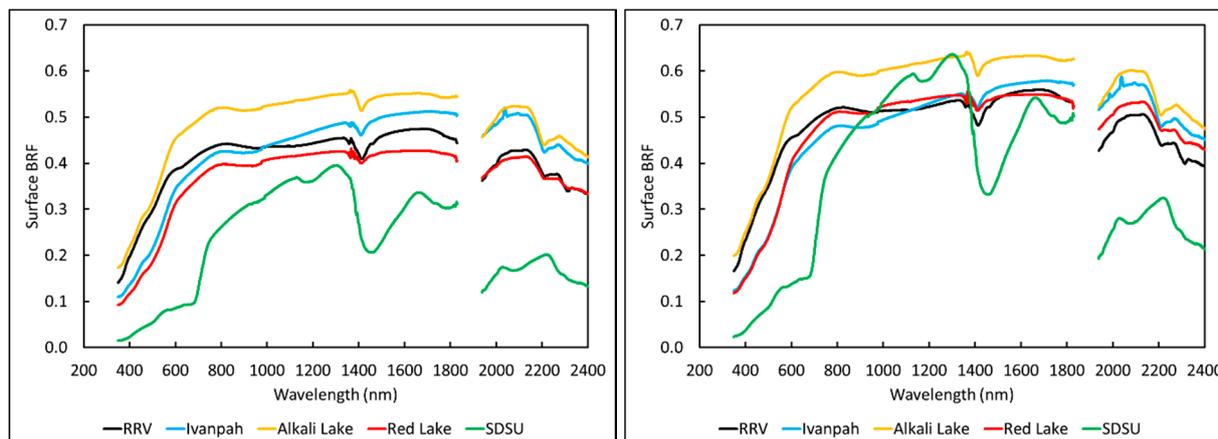


Figure 2. The hyperspectral surface BRF of the various test sites during typical summer illumination angles (left). The results are obtained from portable spectroradiometers that are carried across each site. The same results are shown with the solar zenith angle effect removed (right). Noisy data due to atmospheric absorption, which creates a low SNR in the portable spectrometers, have been removed for clarity.



Figure 3. Surface BRF collection at Ivanpah Playa (left) and the SDSU site (right).

The Spectralon field panels used by RSG and SDSU are routinely measured in RSG's laboratory so that the BRF is known as a function of angle and wavelength, and the panels are always used in the same orientation in the field as well as when being measured in the laboratory [16–19]. The typical BRF for a panel routinely used in the field is shown in Figure 4 for three wavelengths. The final BRF used in the radiative transfer code is determined by spatially averaging over the test area and band averaging for each of the bands of the sensor of interest. The size and orientation of the sampled area depends on the architecture of the satellite sensor and the size of the site. In the case of a pushbroom sensor such as OLI, the long side of the site is perpendicular to the spacecraft's motion. In the case of a whiskbroom sensor such as Landsat 7 ETM+, the long side of the site is oriented in the along-track direction of the satellite.

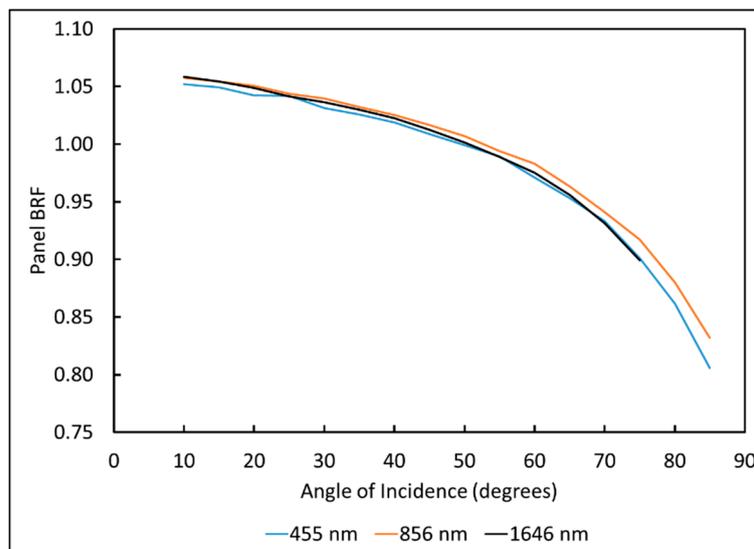


Figure 4. An example of the RSG results for the BRF of a Spectralon panel as a function of angle at three wavelengths. There is minimal spread in the BRF as a function of wavelength, which indicates that this panel has not yellowed due to ultraviolet (UV) exposure.

The size of the site is chosen to include a reasonable sample of the detectors. This is more practical in the case of the 16 multispectral detectors of ETM+, but the large number of detectors in OLI means that only a small sample of the complete detector package can be measured. Based on the heritage of previous measurements for pushbroom sensors such as ASTER, the sites at Ivanpah Playa, Alkali Lake, and Red Lake were chosen to be 120×300 m. The Railroad Valley site has semi-permanent tracks that were established for the calibration of large-footprint sensors such as MODIS, so the site used for OLI is a 90×500 -m portion of the 1-km² site. The collection area at the SDSU site is 100×250 m.

2.1.2. Atmospheric Measurements

Atmospheric measurements are made throughout the day with the goal of determining the atmospheric transmission and radiance for each of the bands in the sensor under test during overpass. The automated solar radiometer (ASR) is the primary atmospheric instrument used in the reflectance-based approach [20] by both RSG and SDSU. An ASR tracks the sun throughout the day and measures the incoming solar irradiance extinction due to atmospheric absorption and scattering. The ASR data are used to determine the total atmospheric spectral optical depths, which are then broken down into subcomponents such as the molecular, aerosol, ozone, and water vapor optical depths [21,22]. In the case of the desert sites used by RSG, it is assumed that the aerosols can be modeled by a power law distribution, which means that the Angstrom exponent can be used to define the aerosol size distribution. Columnar water vapor is determined using a modified-Langley approach [23,24], and the columnar ozone amount from the Ozone Monitoring Instrument (OMI) is used for this work. Ancillary data such as temperature and pressure are measured using equipment that is brought to the site.

At the SDSU site, an ASR is used with a Langley analysis to derive instantaneous optical depth values. The measurements from an Analytical Spectral Devices (ASD) spectrometer and a NIST-traceable reflectance standard are combined to determine the surface reflectance of the field. These values along with standard meteorological measurements and sensor view geometry are used as input into the

MODerate resolution atmospheric TRANsmission (MODTRAN) radiative transfer code. The initial model is propagated from the sun through the atmosphere to ground level, and is used to predict: (1) spectral radiance at the reflectance standard, (2) spectral radiance at the grass target, and (3) the diffuse-to-global irradiance ratio of the sky. The results are compared to the radiance measurements made with the ASD, and the diffuse-to-global ratio is compared to a Yankee Environmental Systems shadow band radiometer. The comparison identifies any anomalies in the model predictions. If any are identified, the model is updated and rerun until a model fit to reality is made. Once the model is matched to the ground measurements, it is propagated to the top of the atmosphere. The data are then band averaged using the RSRs of the sensor under test [25].

2.1.3. Top-of-Atmosphere Spectral Radiance and Reflectance

In general terms, the atmospheric and surface BRF measurements are used as inputs into the MODTRAN radiative transfer code [26]. The *in situ* surface measurements are averaged so that BRF is determined for the entire collection area at 1-nm intervals ranging from 350–2500 nm. The atmospheric inputs include the solar and view geometries, the molecular optical depth, the aerosol optical depth, columnar ozone, columnar water vapor, the Angstrom exponent, and the atmospheric index of refraction. This work uses the ChKur (combined Chance and Kurucz) exoatmospheric solar irradiance, which is distributed with MODTRAN 5. The output is the band-averaged TOA spectral radiance, which is also converted to TOA reflectance for comparison with the two OLI products available from USGS. The main advantage to using TOA reflectance is that the effects due to solar spectrum mismatch are eliminated.

2.2. The Radiometric Calibration Test Site (RadCaTS) at Railroad Valley, Nevada

RadCaTS was developed by RSG to supplement the *in situ* data that are collected during routine field campaigns. It is based on the reflectance-based approach, in that measurements of the atmosphere and surface BRF are made during the time of interest. In the case of RadCaTS, the surface BRF measurements are made using four ground-viewing radiometers (GVRs), and the atmospheric measurements are made using a Cimel sun photometer, which is part of the global Aerosol Robotic Network (AERONET) [27].

2.2.1. Surface Reflectance

The surface BRF measurement in RadCaTS is based on the reflectance-based approach, but it uses four nadir-viewing radiometers that have been absolutely radiometrically calibrated, as opposed to taking the ratio of the DN_s measured over the ground to a reference panel (Figure 5). The current GVRs have eight spectral channels that operate from 400 to 1550 nm (Table 2), and the spectral selection of each channel is controlled using interference filters. The temperature of the critical electronics, including the focal plane, is controlled using a thermo-electric (TE) cooler during daylight operation. Each GVR is an autonomous standalone unit that is powered using a 12-V solar-charged system, and each unit includes its own data logger [28,29]. The field of view is controlled using precise apertures, and there are no powered optics, which is a concept similar to the ASRs that were developed at the University of Arizona.



Figure 5. A ground-viewing radiometer at RadCaTS. The instrument head is 1.5 m above the surface. At this height, each of the eight detectors is viewing the same approximate area on the ground.

Table 2. The interference filter specifications for the eight channels of a GVR. The bandwidth is specified as the full width at half maximum (FWHM).

Band	Center Wavelength (nm)	Bandwidth (nm)
1	400	
2	450	
3	500	
4	550	
5	650	20
6	850	
7	1000	
8	1550	

The surface BRF, ρ , at each of the eight GVR channels is determined by:

$$\rho = \frac{\pi C_{GVR} V_{GVR}}{\frac{E_0}{d^2} \tau_A \cos \theta + E_{SKY}} \quad (1)$$

where C_{GVR} is the radiometric calibration coefficient of the GVR ($\text{W m}^{-2} \text{sr}^{-1} \mu\text{m}^{-1} \cdot \text{V}^{-1}$), V_{GVR} is the output voltage (V), E_0 is the exoatmospheric spectral solar irradiance when the Earth-Sun distance is one astronomical unit (AU) ($\text{W m}^{-2} \cdot \mu\text{m}^{-1}$), d is the Earth-Sun distance (AU), τ_A is the direct solar beam transmission (unitless), θ is the solar zenith angle, and E_{sky} is the diffuse spectral sky irradiance ($\text{W m}^{-2} \cdot \mu\text{m}^{-1}$). AERONET Cimel sun photometer data are used as input into MODTRAN to determine the direct solar beam transmission and the diffuse sky irradiance. The process of determining the hyperspectral surface BRF for the time of interest is as follows:

1. The band-averaged surface BRF is determined for each channel of every GVR using Equation (1).
2. The band-averaged BRF of the entire 1-km² site in each of the eight GVR channels is determined by averaging the individual GVR measurements for each channel.
3. A reference spectra of hyperspectral data is chosen from a library of Railroad Valley BRF spectra.
4. The chosen reference spectra is fit to the average BRF using a least squares fit, which produces a hyperspectral BRF for the date and time of interest.

The reference hyperspectral library is created from 12 years of *in situ* measurements of the 1-km² area of Railroad Valley that is used to calibrate large-footprint sensors such as MODIS, MISR, the Visible Infrared Imaging Radiometer Suite (VIIRS) and the Advanced Very High Resolution Radiometer (AVHRR) [30–32]. The reference data are binned into monthly averages, which is reasonable because the surface BRF at Railroad Valley is largely cyclical from year to year. An example of multispectral measurements being used to scale reference data to produce a hyperspectral data set for the time of interest is shown in Figure 6. There are well-known conditions when the monthly averages are not indicative of the site, and this includes times when the surface is wet after a rain event, and also the presence of efflorescent salts that emerge when the surface is returning to a dry and stable state. In these cases, the individual measurements of the GVRs may have a bias in BRF if they are viewing a portion of the ground that is atypical of the overall surface. There is also the possibility that snow is present on the site during the winter. In this case, *in situ* surface data that were collected in January 2013 when Railroad Valley was covered with spatially-uniform snow are used as the reference BRF that is scaled by the GVR measurements. The BRF of snow is typically very high in the VNIR region of the spectrum, but very low in the shortwave infrared (SWIR), which affects OLI bands 6 and 7.

2.2.2. Atmospheric Measurements

The atmospheric measurements at RadCaTS are made using a Cimel sun photometer, which collects direct solar irradiance and sky radiance data throughout the day. The data at RadCaTS are uploaded to NASA GSFC, processed, and made available on the Internet. The data used for RadCaTS processing include the aerosol optical depth at 550 nm, the Angstrom exponent, the columnar water vapor amount, and the OMI-derived columnar ozone amount. The values are used as input into MODTRAN to determine the spectral transmission of the solar beam, as well as the diffuse sky irradiance. Radiative transfer codes typically overestimate the diffuse component of the total downwelling spectral irradiance, so field measurements were made in 2011 to determine the accuracy of MODTRAN at Railroad Valley. Results from this work include a diffuse irradiance correction factor that is currently in use by RSG [33]

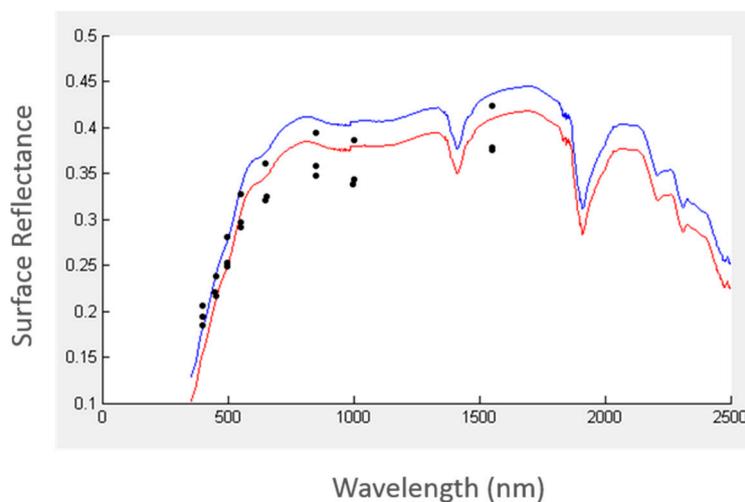


Figure 6. An example showing how the multispectral surface BRF results from three GVRs are used to scale reference hyperspectral BRF data. The reference data (blue) are shifted up or down using the average of the multispectral GVR data (black points). Each of the points at a given wavelength is from one of the three GVRs used during this specific date. The red curve shows the resulting hyperspectral data that are used in the radiative transfer code for a specific date and time of interest.

2.2.3. Top-of-Atmosphere Spectral Radiance and Reflectance

The output from MODTRAN for the RadCaTS data is similar to that of the *in situ* measurements. The measurements from each GVR are averaged to create one hyperspectral BRF data set for the entire 1-km² site for the date and time of interest. MODTRAN is used to process the hyperspectral surface BRF and atmospheric measurements into band-averaged TOA quantities, which are then compared to the imagery obtained from USGS.

The TOA spectral radiance, L_λ , is given by:

$$L_\lambda = \frac{\rho_\lambda}{\pi} \left(\frac{E_{0\lambda}}{d^2} \tau_{A\lambda} \cos \theta + E_{SKY\lambda} \right) \quad (2)$$

where ρ_λ is the hyperspectral BRF determined using the GVRs, $E_{0\lambda}$ is the exoatmospheric spectral solar irradiance when the Earth-Sun distance is one AU ($\text{W} \cdot \text{m}^{-2} \cdot \mu\text{m}^{-1}$), d is the Earth-Sun distance (AU), $\tau_{A\lambda}$ is the direct solar beam transmission (unitless), θ is the solar zenith angle, and $E_{sky\lambda}$ is the diffuse spectral sky irradiance ($\text{W} \cdot \text{m}^{-2} \cdot \mu\text{m}^{-1}$). The resulting hyperspectral TOA spectral radiance is band averaged to each channel of the sensor under test (e.g., OLI or ETM+).

3. Data

3.1. Reflectance-Based Approach and Cross Calibration with ETM+

The ground-based data collection began after the completion of the check-out phase, which was thirty days after launch. The test sites used were Railroad Valley, Nevada, Ivanpah Playa, California, Alkali Lake, Nevada, and Red Lake, Arizona. In the early stages of the field collects, RSG was joined by members of NASA GSFC, who were collecting data for OLI as well as for the calibration of the NASA

GSFC LiDAR Hyperspectral and Thermal (G-LiHT) airborne instruments [34]. The field team from SDSU was concurrently collecting data at the SDSU field site. A summary of the conditions during the field campaigns is shown in Tables 3 and 4, which also include the information for the tandem flight period. RSG and GSFC were at Red Lake, Arizona, on 29 March 2013, and SDSU was at the SDSU site on 30 March 2013. The ground-based measurements can be thought of as a transfer radiometer and SIS. That is, the test sites act as stable radiometric sources during each overpass, which occurred within several minutes of one another. The cross comparison is performed by comparing ETM+ and OLI to the ground-based measurements.

Table 3. A summary of the *in situ* field campaign conditions for the University of Arizona RSG (UA) and GSFC data collection.

Date (yyyy-mm-dd)	2013-03-19	2013-03-24	2013-03-29	2013-04-20	2013-04-29	2013-10-15	2013-11-14	2014-02-04	2014-03-08	2014-05-09
Site (Team)	Railroad Valley (UA)	Ivanpah Playa (UA, GSFC)	Red Lake (UA, GSFC)	Railroad Valley (UA)	Ivanpah Playa (UA)	Red Lake (UA)	Railroad Valley (UA)	Red Lake (UA)	Red Lake (UA)	Railroad Valley (UA)
Collection area (m)	90 × 500	120 × 300	120 × 300	90 × 500	120 × 300	120 × 300	90 × 500	120 × 300	120 × 300	90 × 500
Time (UTC)	18:19:54	18:16:29	18:05:34 (L7) 18:12:12 (L8)	18:22:47	18:17:26	18:11:03	18:22:47	18:10:33	18:09:55	18:20:38
Solar zenith angle (degrees)	44.0	40.0	38.9 (L7) 38.0 (L8)	32.0	27.7	46.6	58.9	56.3	46.3	27.1
Solar azimuth angle (degrees)	146.4	142.8	140.0 (L7) 142.3 (L8)	141.4	133.7	159.7	162.1	151.2	145.6	135.2
Aerosol Optical Depth (550 nm)	0.048	0.068	0.068	0.085	0.122	0.046	0.021	0.032	0.025	0.089
Ozone (DU)	298	296	306	309	323	286	326	333	311	356
Water vapor (cm)	1.19	0.77	1.64	1.10	1.24	0.77	0.58	0.82	0.60	1.21
Temperature (°C)	14.1	16.4	25.1	15.5	33.9	18.8	10.3	9.0	18.4	20.5
Pressure (mb)	859.3	926.7	922.8	859.5	919.0	919.8	859.2	920.0	925.6	853.1

Table 3. Cont.

Date (yyyy-mm-dd)	2013-03-19	2013-03-24	2013-03-29	2013-04-20	2013-04-29	2013-10-15	2013-11-14	2014-02-04	2014-03-08	2014-05-09
Satellite Information										
View zenith angle (degrees)	4.6	4.3	5.4 (L7) 2.7 (L8)	0.8	3.7	5.9	0.3	5.1	5.5	0.3
View azimuth angle (degrees)	102.6	103.1	101.9 (L7) 100.7 (L8)	95.0	96.9	101.7	122.6	100.5	95.1	148.9

Table 4. A summary of the *in situ* field campaign conditions for the SDSU data collection.

Date (yyyy-mm-dd)	2013-03-30	2013-06-10	2013-07-12	2013-08-13	2013-08-29	2013-09-30	2013-10-16
Site (Team)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)	SDSU Test Site (SDSU)
Collection area (m)	100 × 250	100 × 250	100 × 250	100 × 250	100 × 250	100 × 250	100 × 250
Time (UTC)	17:07:14 (L7) 17:09:43 (L8)	17:13:23	17:13:22	17:13:23	17:13:25	17:13:15	17:13:14
Solar zenith angle (degrees)	43.9	26.0	27.8	34.2	38.8	49.5	55.2
Solar azimuth angle (degrees)	151.3	139.2	137.8	145.0	150.0	159.1	162.2
Aerosol Optical Depth (550 nm)	0.158	0.060	0.183	0.195	0.067	0.016	0.077
Ozone (DU)	324	287	290	291	289	284	279
Water vapor (cm)	0.53	2.76	1.76	2.93	2.64	2.01	3.02
Temperature (°C)	7.2	18.9	27.8	21.7	31.0	22.5	8.9
Pressure (mb)	951.9	947.9	948.9	956.7	945.5	936.0	954.6
Satellite Information							
View zenith angle (degrees)	0.1 (L7) 5.0 (L8)	0.1	0.3	0.3	0.2	0.3	0.2
View azimuth angle (degrees)	278.9 (L7) 278.9 (L8)	278.9	278.9	278.9	278.9	278.9	278.9

3.2. RadCaTS

A summary of the conditions at each of the successful RadCaTS collects is shown in Table 5, which shows the solar and satellite viewing geometry as well as the key atmospheric components that are used in MODTRAN. The collection dates should be similar to those of the *in situ* measurements made using on-site personnel, but two dates in March and April 2013 were considered outliers due to spatial

sampling issues with the GVRs. In each of these cases, the efflorescent salt on the site changed the spatial homogeneity such that the GVR measurements were not truly indicative of the overall BR_F of the 1-km² site. *In situ* measurements using on-site personnel have much higher spatial sampling, so spatial non-uniformity effects are greatly reduced. There were 26 Landsat 8 overpasses at RadCaTS during the period March 2013–May 2014, and seven collects were successful. Two of the seven were excluded because of spatial uniformity issues in March and April 2013. The main cause of the unsuccessful collects was the presence of clouds at the site. This was determined using a combination of sources such as the output voltage of the GVRs, the AERONET aerosol optical depth and Angstrom parameter, and the OLI cirrus band.

Table 5. A summary of the conditions for each of the RadCaTS collections in 2013. The atmospheric results are determined using an AERONET Cimel sun photometer. The data from 19 March 2013 and 20 April 2013 are considered outliers due to the reduced spatial uniformity of the site.

Date (yyyy-mm-dd)	2013-05-22	2013-06-07	2013-06-23	2013-08-10	2013-11-14
Site	Railroad Valley				
Collection area (m)	1000 × 1000	1000 × 1000	1000 × 1000	1000 × 1000	1000 × 1000
Time (UTC)	18:23:06	18:23:03	18:22:54	18:22:49	18:22:47
Solar zenith angle (degrees)	24.4	23.0	23.1	29.8	58.9
Solar azimuth angle (degrees)	132.1	127.4	124.8	135.1	162.1
Aerosol Optical Depth (550 nm)	0.062	0.051	0.044	0.042	0.019
OMI Ozone (DU)	335	300	323	311	286
Water vapor (cm)	0.75	0.84	0.87	0.83	0.41
Temperature (°C)	25.0	32.5	27.8	27.8	8.8
Pressure (mb)	845	857	851	858	860
Satellite Information					
View zenith angle (degrees)	0.2	0.3	0.7	0.9	0.3
View azimuth angle (degrees)	85.2	76.0	100.0	91.7	122.6

4. Results

4.1. Reflectance-Based Results

After 14 months on orbit, there were nineteen ground-based data sets collected using the reflectance-based approach (UA: 10 data sets, SDSU: seven data sets, GSFC: two data sets). They originate from five sites, and the results for both the TOA spectral radiance and reflectance are shown in Figure 7. The results are shown as a percentage difference between the ground-based results and OLI:

$$\text{Percent difference} = (\text{Measured} - \text{OLI})/\text{OLI} \quad (3)$$

where *Measured* are the ground-based data. In both the TOA spectral radiance and reflectance cases, the results from both the desert sites and vegetated site agree with OLI to within the uncertainties of the methods, which are on the order of 2.5%–3.0%. There is a divergence between the desert and vegetated sites in the results for OLI bands 1 and 2, which are the coastal aerosol (443 nm) and blue (483 nm)

bands. The source of this discrepancy is most likely due to a combination of very low surface reflectance (Figure 2) and atmospheric effects, which play a larger role when the surface reflectance is less than 0.2. At wavelengths greater than the vegetation red edge (700 nm), the surface reflectance of the vegetated site increases rapidly to values similar to the desert sites. In the multispectral bands 3–7, the results between the two types of sites and OLI agree to within the uncertainties of the methodology. In the bottom row of Figure 7, the data from the three field teams are averaged into one final data set for both TOA spectral radiance and reflectance. In the case of both TOA spectral radiance and reflectance, OLI is generally in agreement with the ground-based results to within the design specifications (5% and 3%, respectively). The only band that is just out of specification is 7 (2.2 μm), which typically has a very low signal at both the desert and vegetated sites.

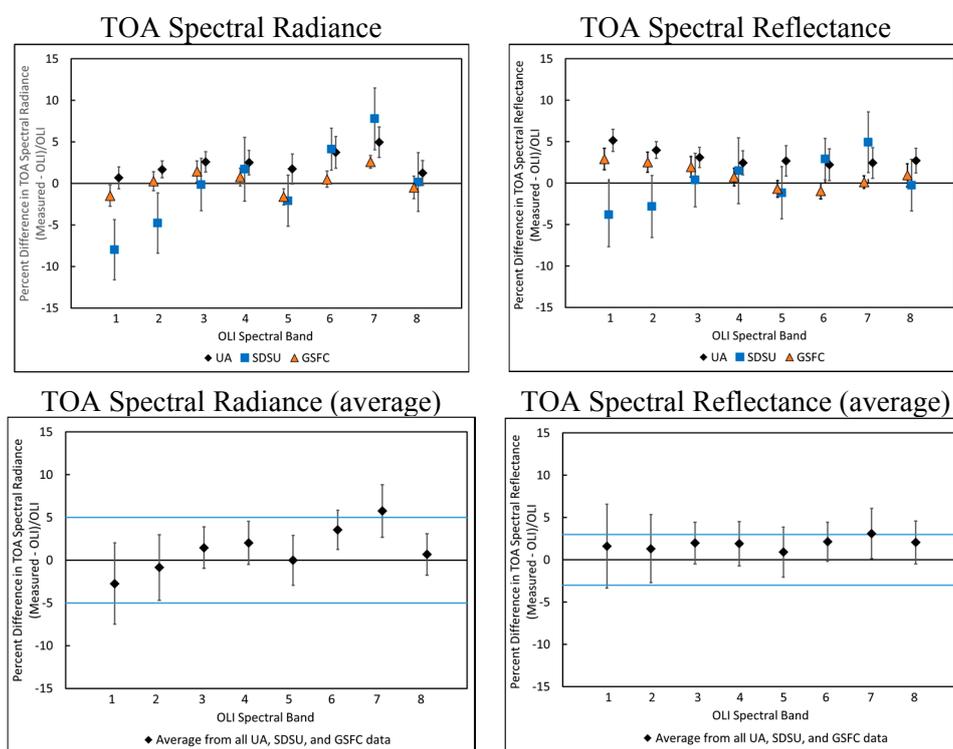


Figure 7. A summary of *in situ* results from the University of Arizona RSG (UA), SDSU, and GSFC for TOA spectral radiance (top left) and reflectance (top right) for Landsat 8 OLI. The uncertainty bars are the 1σ standard deviation of the measurements by each team. Note that band 9 (cirrus) is excluded from the results due to an extremely low SNR in the ground measurements. The bottom row shows the results for TOA spectral radiance (bottom left) and reflectance (bottom right) when the ground-based results from all three teams are consolidated into one data set and then compared to OLI. The uncertainty bars in the average graphs are the 1σ standard deviation of the average from all three teams.

Figures 8 and 9 show the temporal results of TOA spectral radiance and TOA reflectance, respectively, for the fourteen-month duration of this work. The RSG data gap between Days Since Launch (DSL) 77 and 246 was caused by poor weather conditions during the attempted field campaigns.

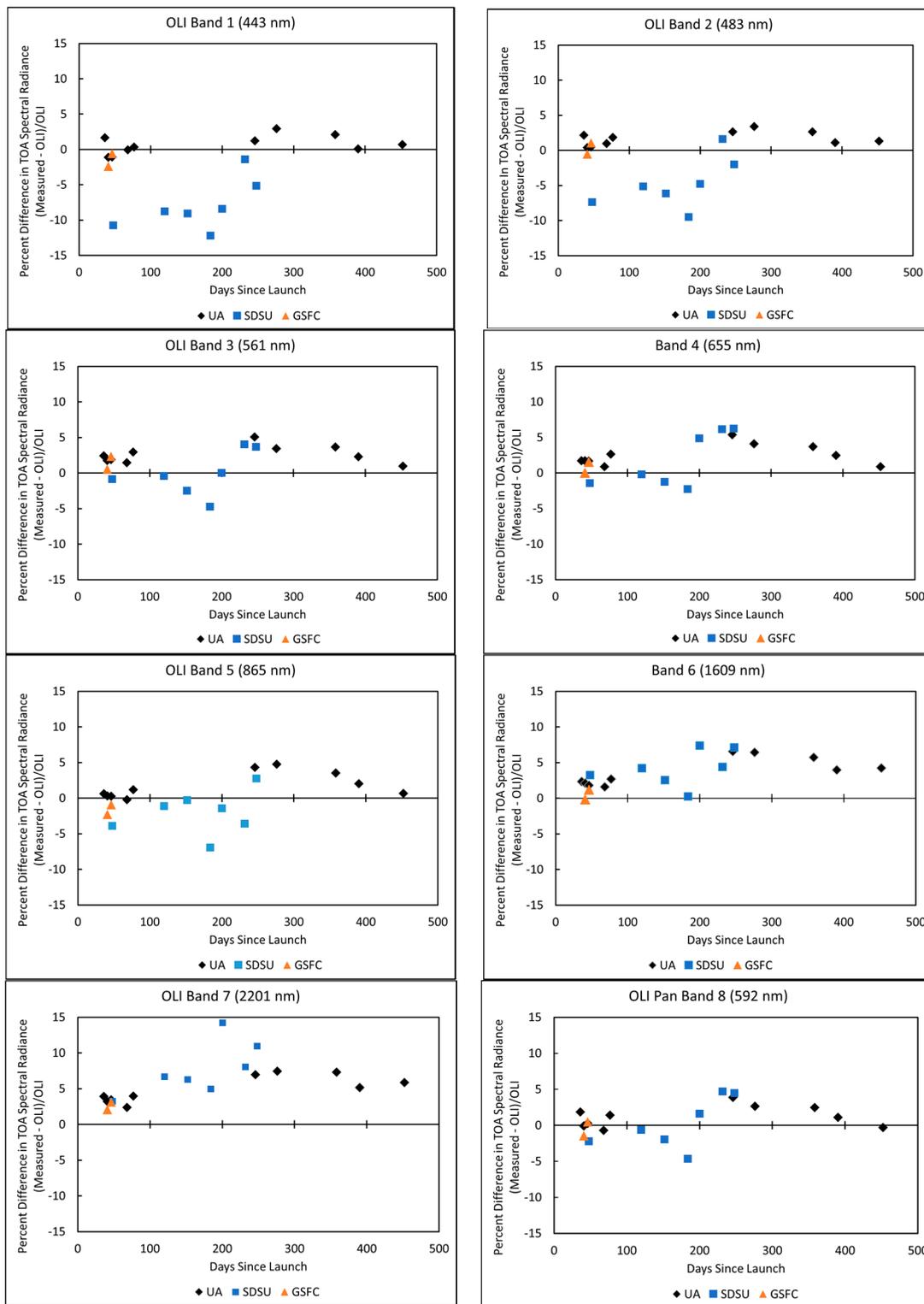


Figure 8. The temporal TOA spectral radiance results from the *in situ* measurements for Landsat 8 OLI. The percentage difference is determined using $(\text{Measured} - \text{OLI})/\text{OLI}$.

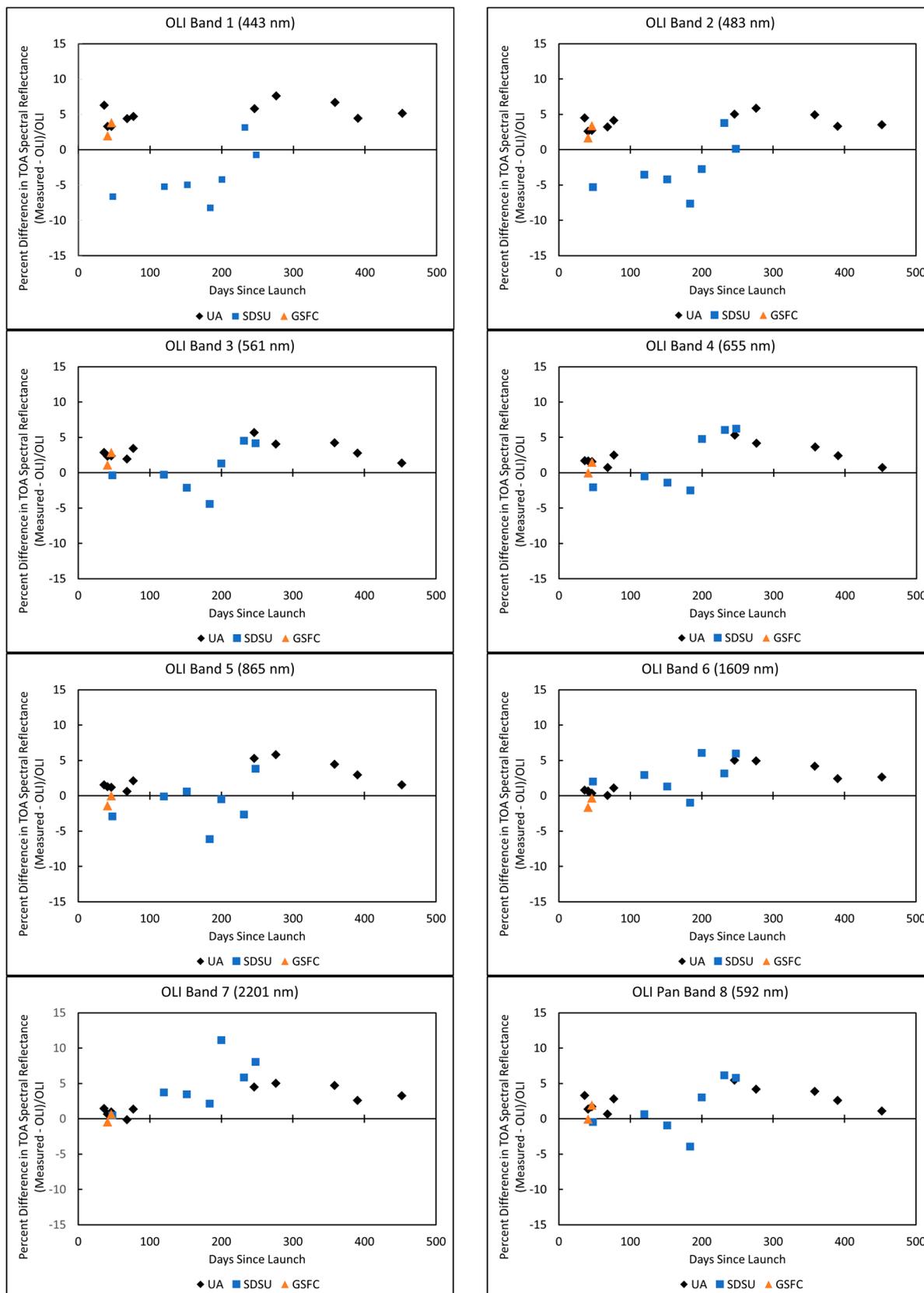


Figure 9. The temporal TOA spectral reflectance results from the *in situ* measurements for Landsat 8 OLI. The percentage difference is determined using $(\text{Measured} - \text{OLI})/\text{OLI}$.

The SDSU data gap after DSL 250 is due to winter conditions and the high probability of snow on the site. The temporal results in both the TOA spectral radiance and reflectance cases do not show a discernable trend in degradation in any of the channels at this time.

4.2. RadCaTS Results

The RadCaTS results are from data collected between March 2013 and May 2014, and originate from five successful collects. Figure 10 shows the average of the complete data set, and also includes the University of Arizona RSG results from the reflectance-based approach. This is a logical comparison, because the reflectance-based approach is considered to be the gold standard of RSG's measurements, and the goal of RadCaTS is to have accuracy and precision on par with that of the reflectance-based approach. The uncertainty bars in Figure 10 are the 1σ standard deviation of the measurements. The results for both OLI products are similar to the reflectance-based results, but there is a reasonably constant bias between the on-site-personnel-based measurements and RadCaTS in the TOA spectral radiance as a function of wavelength. The bias in the TOA spectral radiance results do not have a pattern as a function of wavelength.

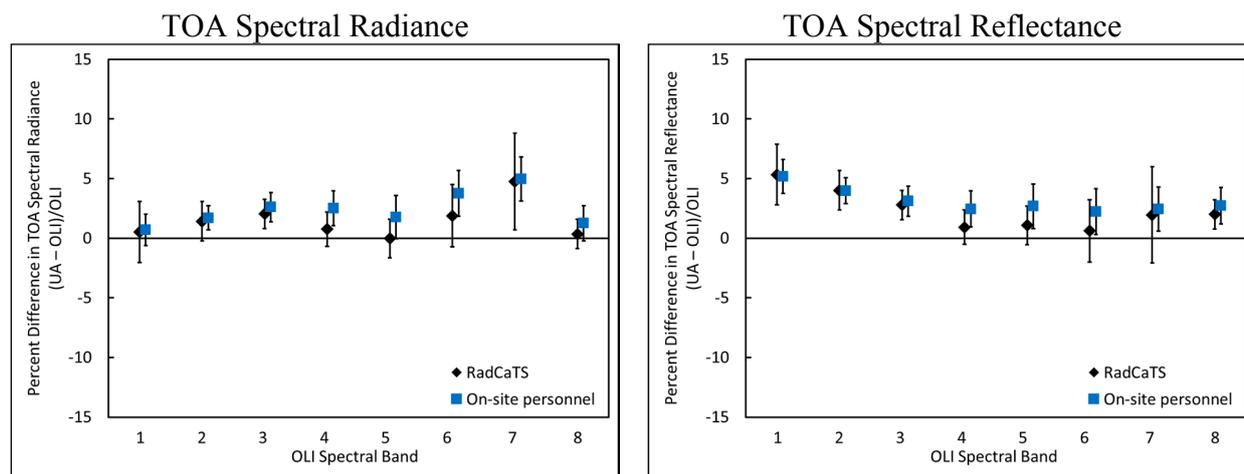


Figure 10. The average TOA spectral radiance and reflectance results from RadCaTS for Landsat 8 OLI. The RadCaTS results are shown with the *in situ* data collected by RSG at Railroad Valley for comparison.

4.3. Landsat 7 ETM+ and Landsat 8 OLI Cross Comparison Results

The tandem flight in March 2013 produced one data set at each of two sites. UA and GSFC collected data at Red Lake Arizona, while SDSU personnel collected data at the SDSU site. The results are shown in Figure 11 for the TOA spectral radiance and TOA spectral reflectance. In each of the two graphs, the percent difference between the Landsat sensor and the average of the ground-based measurements is shown. Similar ETM+ and OLI spectral bands are combined for comparison. In the case of TOA spectral radiance, the largest difference between ETM+ and OLI is 1.8%, which occurs at the extreme spectral regions (blue and SWIR). The percent difference in the middle bands is typically less than 1%. In the case of TOA spectral reflectance, the largest difference between ETM+ and OLI is in the SWIR bands (3%–4%). The ETM+ and OLI TOA spectral reflectance results are within 1% in the VNIR bands.

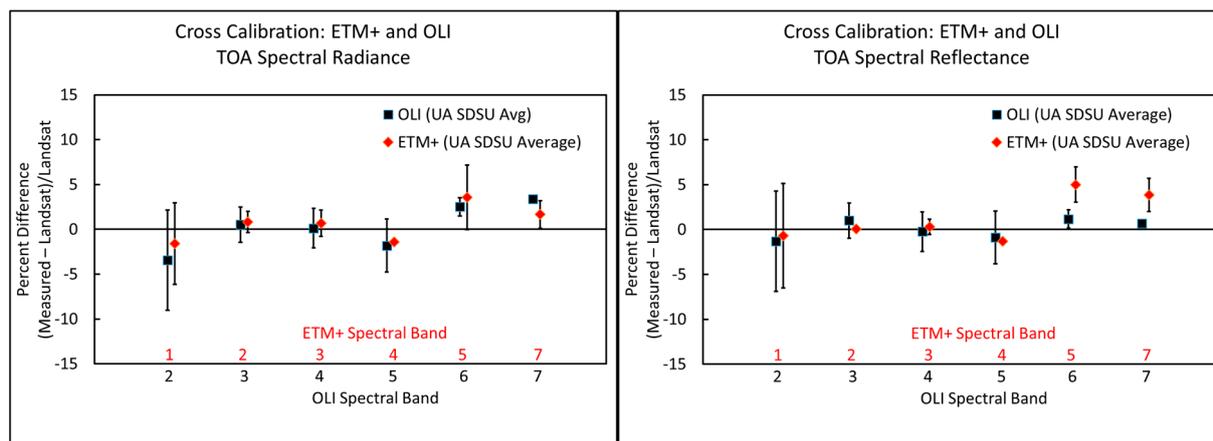


Figure 11. The cross calibration of Landsat 7 ETM+ and Landsat 8 OLI during the three-day tandem flight period in March 2013. The percent difference between the sensor and the ground-based measurements is determined as $(\text{Measured} - \text{Landsat})/\text{Landsat}$, and the uncertainty bars are the 1σ standard deviation of the measurements.

5. Uncertainty Analysis

5.1. Reflectance-Based Approach

The uncertainty of the reflectance-based approach has been analyzed by RSG in support of work with Landsat 5 TM, Satellite Pour l'Observation de la Terre (SPOT) Haute Resolution Visible (HRV), and Landsat 7 ETM+. Early work showed that the reflectance-based approach had an absolute radiometric uncertainty of $\pm 5\%$, and estimated that improvements should result in an uncertainty of $\pm 3\%$ in the middle of the visible portion of the spectrum [12,22,35]. More recently, the uncertainty was re-evaluated to include updates to the surface reflectance retrieval, which is typically the primary source of uncertainty in the reflectance-based approach [36]. The uncertainty of the reflectance-based approach was determined to be 2.5%. RSG has since updated the laboratory methodology for characterizing the field reference panels, and this has decreased the uncertainty of the surface BRF retrieval. In this work, the uncertainty analysis in [36] is updated with the current surface BRF uncertainty, and the specified uncertainty of MODTRAN 5, which is $\pm 2\%$ in radiance [26].

An important component in the reflectance-based approach is the determination of the test site surface BRF, and the uncertainty is dominated by the laboratory determination of the reference panel BRF. The previous method involved pressing polytetrafluoroethylene (halon) into a reference sample using NIST-specified directions [17]. This was then used as a laboratory reference to which the field panels would be compared to determine their BRF as a function of wavelength and angle. More recently, RSG sends a Spectralon panel to NIST to be characterized, and upon its return, it is used as the new laboratory standard to which the field panels are compared. The main advantage to this process is the decrease in uncertainty in the determination of the BRF for a field panel. In previous work, the uncertainty in reflectance panel BRF calibration was estimated to be $\pm 2\%$ (1σ). The most current laboratory panel measurement uncertainty estimates have been updated to be $\pm 1.0\%$ – 1.4% throughout most of the VNIR and SWIR regions [19]. The reflectance-based uncertainty table in [36] has been updated with the new panel uncertainty values, and the new results are shown in Table 6. Assuming an average panel BRF

measurement uncertainty of $\pm 1.3\%$ over the VNIR and SWIR, the current uncertainty in the TOA spectral radiance is $\pm 2.6\%$ in the middle of the visible spectral region. The value of “---” is shown in the absorption computation for the 2005 analysis because it addressed the fact that wavelengths could be chosen where there is a negligible effect on the retrieved TOA spectral radiance due to water vapor or ozone. The same value of “---” is shown in the non-lambertian surface column, and it is justified by previous research that has shown that the non-lambertian nature of Railroad Valley is negligible in the determination of TOA spectral radiance for near-nadir view angles and typical aerosol conditions.

Table 6. The uncertainty of the reflectance-based approach, which includes updated values for the laboratory calibration of Spectralon reference panels. The table is based on earlier results presented in [35] and is valid for clear-sky conditions in the middle of the visible spectral region.

Source of Uncertainty	Source	TOA Spectral Radiance
Ground reflectance measurement		1.6%
Reference panel calibration (BRF)	1.3%	
Site measurement errors	1.0%	
Optical depth measurements	0.5%	<0.1%
Extinction optical depth	0.5%	
Absorption computations		---
Column ozone (OMI)	3.0%	
Column water vapor	5.0%	
Choice of aerosol complex index	100%	0.5%
Choice of aerosol size distribution		0.3%
Junge parameter	0.3%	
Non-lambertian surface	10%	---
Other		
Non-polarization vs. polarization		0.1%
Inherent accuracy of MODTRAN 5		2.0%
Uncertainty in solar zenith angle		0.2%
Total RSS uncertainty		2.6%

5.2. RadCaTS

The RadCaTS methodology is modeled on the reflectance-based approach, and the determination of surface BRF is also the primary influence in the uncertainty of TOA quantities. In the RadCaTS processing scheme, the surface BRF, ρ , is determined using Equation (1). The uncertainty, $\Delta\rho$, is given by:

$$\Delta\rho = \sqrt{\left(\frac{\partial\rho}{\partial C_{GVR}} \Delta C_{GVR}\right)^2 + \left(\frac{\partial\rho}{\partial V_{GVR}} \Delta V_{GVR}\right)^2 + \left(\frac{\partial\rho}{\partial E_0} \Delta E_0\right)^2 + \left(\frac{\partial\rho}{\partial \tau_A} \Delta \tau_A\right)^2 + \left(\frac{\partial\rho}{\partial \theta} \Delta \theta\right)^2 + \left(\frac{\partial\rho}{\partial E_{SKY}} \Delta E_{SKY}\right)^2} \quad (4)$$

where ΔC_{GVR} is the uncertainty in the GVR radiometric calibration coefficient, ΔV_{GVR} is the uncertainty in the output voltage, ΔE_0 is the uncertainty in the exoatmospheric solar irradiance, $\Delta \tau_A$ is the uncertainty in the atmospheric transmission, $\Delta \theta$ is the uncertainty of the solar zenith angle, and ΔE_{SKY} is the uncertainty in the diffuse sky irradiance. Similarly, the uncertainty in TOA spectral radiance, ΔL , determined in Equation (2) is:

$$\Delta L = \sqrt{\left(\frac{\partial L}{\partial \rho} \Delta \rho\right)^2 + \left(\frac{\partial L}{\partial E_0} \Delta E_0\right)^2 + \left(\frac{\partial L}{\partial \tau_A} \Delta \tau_A\right)^2 + \left(\frac{\partial L}{\partial \theta} \Delta \theta\right)^2 + \left(\frac{\partial L}{\partial E_{SKY}} \Delta E_{SKY}\right)^2} \quad (5)$$

The main sources of uncertainty originate from:

- (1) Radiometric calibration of the GVRs.
- (2) The radiative transfer code uncertainty, which includes:
 - (a) Uncertainty in the solar irradiance model.
 - (b) Atmospheric transmission.
 - (c) Determination of the diffuse sky irradiance.
- (3) Surface BRDF retrieval, which includes uncertainties introduced by scaling the hyperspectral reference BRDF data with point measurements from the GVRs.

(1) Radiometric calibration of the GVRs

The main components of the GVR calibration are the determination of the radiometric calibration coefficient and the SNR of each channel, both of which are measured in RSG's calibration laboratory prior to deployment [29]. The uncertainty in the radiometric calibration coefficient, ΔC_{GVR} , is estimated to be $\pm 2\%$ for a new radiometer, and future work will determine how this changes with time. The SNR values, which are inversely proportional to the uncertainty in the output voltage, ΔV_{GVR} , range from 1725–7105, depending on the channel and GVR.

(2a) Uncertainty in the solar irradiance model

The uncertainty in the solar model, E_0 , was estimated by referencing the Thuillier database of solar spectral irradiance, which was derived from spectrometer measurements made on the Space Shuttle during the three ATmospheric Laboratory for Applications and Science (ATLAS) missions [37,38]. One major advantage to using these measurements is that they were made above the Earth's atmosphere, thereby minimizing uncertainties that are introduced by atmospheric correction algorithms. The Thuillier database uncertainty was also cited because it provides a well-documented description of the expected uncertainties in the measurements, and also because it is recommended by CEOS for use as a standard exoatmospheric irradiance for the purpose of satellite sensor calibration. In summary, the uncertainty in exoatmospheric solar irradiance, $\Delta E_0/E_0$, is estimated to be $\pm 1.5\%$ at 450 nm, $\pm 0.9\%$ at 650 nm, $\pm 1.1\%$ at 850 nm, and $\pm 0.8\%$ at 1550 nm (all values are the 1σ standard deviation).

(2b) Uncertainty in the atmospheric transmission

RadCaTS uses the MODTRAN radiative transfer code to determine the atmospheric transmission. The atmospheric inputs to MODTRAN for the reflectance-based approach are based on measurements from an automated solar radiometer [20]. In order to understand the uncertainty in the MODTRAN output, the uncertainty in the ASR was analyzed. This began with an analysis of the Langley method, which is used to calibrate the ASR in each of its ten channels. The zero-airmass intercept uncertainty for each of the ten channels was $\pm 1.5\%$ – 3.0% . From this, the uncertainty in the various atmospheric optical depths (e.g., total optical depth, molecular, ozone, and aerosol) was determined. The actual uncertainty analysis on the atmospheric transmission was determined by perturbing the inputs to the code based on the uncertainty in the ASR and the uncertainty in the on-site ancillary instruments that are used to

measure temperature and pressure, for example. The ozone input is determined using OMI data, which are stated to have an uncertainty of $\pm 3\%$ [39]. The results from the study showed that the uncertainty in atmospheric transmission, $\Delta\tau_A$, ranged from $\pm 1.1\%$ in the blue region of the spectrum down to $\pm 0.3\%$ in the SWIR region of the spectrum.

RadCaTS uses a Cimel sun photometer to perform atmospheric measurements. The instrument is routinely calibrated at NASA GSFC, and the uncertainty in aerosol optical depth during clear-sky conditions is stated to be ± 0.010 for wavelengths greater than 440 nm, and ± 0.020 for wavelengths less than 440 nm [27]. An analysis of the ASRs indicate that the uncertainty of the retrieved aerosol optical depth at 550 nm is ± 0.006 on average for typical clear-sky conditions. Even though the uncertainty in the aerosol optical depth that is determined using a Cimel is slightly larger than that from an ASR, this work assumes that the uncertainty in atmospheric transmittance is of the same order of magnitude. Further analysis is underway to compare the atmospheric properties determined using the Cimel and ASR, and their effect on the TOA spectral radiance uncertainty.

(2c) Uncertainty in the diffuse sky irradiance

The diffuse sky irradiance, E_{SKY} , is perhaps one of the more difficult quantities to determine accurately using radiative transfer codes. In 2011, RSG personnel directly measured E_{SKY} during an experiment to determine the accuracy of the MODTRAN 5.2 output under typical clear-sky conditions at Railroad Valley [33]. A hyperspectral spectroradiometer was used to measure the surface reflectance at the same time as a GVR. It was found that the MODTRAN prediction of E_{SKY} at Railroad Valley was as much as 50% higher in the VNIR than measurements made using a calibrated spectroradiometer. This discrepancy can change the surface reflectance retrieval by as much as 10%, most notably in the blue region where the ratio of diffuse-to-global solar irradiance is greatest. An E_{SKY} correction factor that accounted for the large difference between MODTRAN and the measured values was created and is currently in use. After the correction is applied, the uncertainty in E_{SKY} is estimated to be $\pm 5\%$ at this time. The correction factor is currently applicable to conditions that are similar to those when the measurements were made, that is, low aerosol optical depth ($\delta_{550\text{ nm}} < 0.10$) and cloud free. Additional field work will be completed in the future to further understand and enhance the correction factor for E_{SKY} .

(3) Uncertainty in surface BRF retrieval

The four GVRs at RadCaTS are effectively making point measurements of a 1-km² region. Two studies were undertaken to determine the accuracy of this configuration, and to see how well the results compare to the typical manned vicarious calibration method, where a portable spectroradiometer is carried 4.8 km throughout the site [40,41]. Results from both of these studies indicate that because of Railroad Valley's homogeneity, four radiometers are adequate to sample the 1-km² site to within $\pm 2\%$ of the average value. In terms of accuracy, the surface reflectance determined using four RadCaTS GVRs agreed to within $\pm 1\%$ of those made using traditional portable spectroradiometers.

The overall uncertainty in the surface BRF retrieval for RadCaTS consists of two parts: the inherent uncertainty of an individual GVR measurement, and the uncertainty introduced by fitting the hyperspectral BRF reference data to the BRF determined by the GVRs. That is, the resulting hyperspectral BRF data will differ slightly from the average of the GVR measurements for a given date

and time (Figure 6). The uncertainty created by shifting the hyperspectral reference data was determined by measuring the magnitude of the scaling effect for twenty three dates in 2013 when RadCaTS was in use. The scaling factor was analyzed at 18:30 and 20:50 UTC, which is approximately symmetric in time around solar noon. The analysis was also performed to determine if there is a morning-afternoon bias since most of the *in situ* reference BRF data were collected during morning overpasses (e.g., Terra MODIS). Table 7 shows the overall uncertainty in BRF, which is the RSS of the two individual uncertainties.

Table 7. The RadCaTS surface BRF uncertainty (1σ) for four of the GVR spectral bands for typical clear-sky conditions and a solar zenith angle of 45° .

Source of Uncertainty	Surface BRF Uncertainty			
	GVR Band			
	450 nm	650 nm	850 nm	1550 nm
Individual GVR measurement	2.7%	2.4%	2.5%	2.2%
Scaling hyperspectral reference data	2.6%	1.7%	2.8%	3.3%
Total BRF uncertainty	3.7%	2.9%	3.8%	4.0%

RadCaTS Uncertainty Results

The uncertainty analysis results assume that there are clear-sky conditions, and the solar zenith angle is 45° , which was chosen as a representative value for the overpass times of sun-synchronous Earth-observing sensors at Railroad Valley. When extremes in solar zenith angle are used (e.g., 15° and 60°), the uncertainty in surface BRF and TOA spectral radiance varies by $\pm 0.5\%$. Typical spectrally-dependent values for E_0 , τ_A , and E_{SKY} were used in this analysis.

When the results of the individual inputs are used in (5), the uncertainty in the TOA spectral radiance determined using RadCaTS is shown in Table 8 for four of the eight GVR bands. The overall uncertainty is generally dominated by the uncertainty in the BRF retrieval, but E_{SKY} plays a greater role in the blue end of the spectrum.

Table 8. The RadCaTS TOA spectral radiance uncertainty (1σ) for four of the GVR spectral channels for typical clear-sky conditions and a solar zenith angle of 45° .

RadCaTS TOA Spectral Radiance Uncertainty			
GVR Band			
450 nm	650 nm	850 nm	1550 nm
4.1%	3.1%	4.1%	4.1%

6. Conclusions

The results from an extensive campaign of ground-based field measurements during the first year of operation have indicated that the Landsat 8 OLI instrument is performing well, and it has remained stable throughout this period. The ground-based data collected for this work were compared to the L1T imagery that are available from USGS through Earth Explorer. When all of the reflectance-based field data are averaged together, there is agreement with OLI in each spectral band to within the design specifications for TOA spectral radiance (5%), and TOA spectral reflectance (3%), except for band 7, which is slightly higher in both cases. The temporal results have not shown an appreciable pattern of degradation in the

OLI instrument. The results from the automated RadCaTS facility at Railroad Valley have shown results similar to those obtained by ground personnel, having slightly better agreement with OLI than *in situ* measurements made by on-site personnel at Railroad Valley. The number of RadCaTS data sets was smaller than one would expect throughout the year, and this was mainly due to the abnormally high number of dates with poor weather conditions throughout the year.

Ground-based calibration work will continue using on-site personnel from the University of Arizona and South Dakota State University. In addition, RadCaTS will also be used to monitor OLI throughout its mission lifetime. The excellent radiometric performance and stability of OLI will assist vicarious calibration teams to refine their methodologies and field collection techniques, which includes the automated RadCaTS facility.

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Author Contributions

All of the authors were involved various aspects of the data collection and/or analysis during the period of this work. Jeff Czaplá-Myers and Nikolaus Anderson were involved in the *in situ* collection and analysis of field data, which included simultaneous field work with Kurt Thome and Joel McCorkel, who provided the GSFC results. Stuart Biggar was part of the preflight calibration team that made measurements at Ball Aerospace and NIST. Larry Leigh and David Aaron were involved in the SDSU field campaigns, and provided the SDSU results for this paper. Dennis Helder and Nischal Mishra provided expertise in the analysis of the field data.

Conflicts of Interest

The authors declare no conflict of interest.

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